

In-situ observations of young contrails – overview and selected results from the CONCERT campaign

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Abstract. Lineshaped contrails were detected with the research aircraft Falcon during the CONCERT – CONtrail and Cirrus ExpeRimenT – campaign in October/November 2008. The Falcon was equipped with a set of instruments to measure the particle size distribution, shape, extinction and chemical composition as well as trace gas mixing ratios of sulfur dioxide (SO₂), reactive nitrogen and halogen species (NO, NO_y, HNO₃, HONO, HCl), ozone (O₃) and carbon monoxide (CO). During 12 mission flights over Europe, numerous contrails, cirrus clouds and a volcanic aerosol layer were probed at altitudes between 8.5 and 11.6 km and at temperatures above 213 K. 22 contrails from 11 different aircraft were observed near and below ice saturation. The observed NO mixing ratios, ice crystal and soot number densities are compared to a process based contrail model. On 19 November 2008 the contrail from a CRJ-2 aircraft was penetrated in 10.1 km altitude at a temperature of 221 K. The contrail had mean ice crystal number densities of 125 cm⁻³ with effective radii r_{eff} of 2.6 μm. The presence of particles with $r > 50$ μm in the less than 2 min old contrail suggests that natural cirrus crystals were entrained in the contrail. Mean HONO/NO (HONO/NO_y) ratios of 0.037 (0.024) and the fuel sulfur con-

version efficiency to H₂SO₄ ($\epsilon_{S\downarrow}$) of 2.9 % observed in the CRJ-2 contrail are in the range of previous measurements in the gaseous aircraft exhaust. On 31 October 2010 aviation NO emissions could have contributed by more than 40% to the regional scale NO levels in the mid-latitude lowest stratosphere. The CONCERT observations help to better quantify the climate impact from contrails and will be used to investigate the chemical processing of trace gases on contrails.

1 Introduction

While aircraft induced cloudiness potentially has the largest aviation impact on climate, the exact magnitude of its contribution to the anthropogenic radiative forcing remains highly uncertain (Lee et al., 2009). Aircraft induced cloudiness plays a role for the climate by reflecting solar radiation and trapping outgoing terrestrial longwave radiation. Latter effect is expected to dominate for contrails and thin cirrus, which results in a net positive radiative forcing (Penner et al., 1999).

Contrails form when the relative humidity in the young exhaust increases due to mixing of the hot and humid exhaust with the colder and less humid ambient air (Schumann, 1996; Kärcher, 1996). When saturation with respect to water is reached, water condenses on soot and sulfate aerosol and



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forms liquid droplets. Thereby the aerosol originates from the combustion process or was entrained in the exhaust with the ambient air (Jensen et al., 1998). When the exhaust further cools down through mixing with ambient air, ice saturation may be reached and ice nucleates in the liquid droplets (Kärcher and Yu, 2009). Hence contrail formation and properties of young contrails may depend on the combustion process and on properties of the contrail-forming aircraft (Schumann, 1996; Schumann et al., 2000; Sussmann and Gierens, 1999). When the ambient air remains supersaturated with respect to ice, the contrail ice crystals can grow by condensation of entrained water vapor (Schröder et al., 2000). Contrail dimensions and contrail coverage increase in ice supersaturated air due to particle sedimentation and contrail spreading in an atmospheric environment with wind shear, while contrail extinction and ice crystal concentrations decrease due to dilution (Freudenthaler et al., 1995).

Observations of the microphysical properties of young contrails are rare. Mean effective radii of contrail ice crystals derived from in situ data show values below $2\text{ }\mu\text{m}$ initially (Heymsfield et al., 1998; Baumgardner and Gandrud, 1998), increasing due to condensation to values of up to $5\text{ }\mu\text{m}$ at 30 min contrail age (Schröder et al., 2000). Initial ice crystal concentrations larger than 1000 cm^{-3} have been detected in a few seconds old contrails (Heymsfield et al., 1998) decreasing by dilution to concentrations of less than 100 cm^{-3} within the first minutes (Poellot et al., 1999; Febvre et al., 2009). In persistent contrails with ages of few hours, ice crystal concentrations of less than 20 cm^{-3} have been observed (Lawson et al., 1998; Schäuble et al., 2009). Contrail widths of 100–300 m have been derived from in situ measurements (Poellot et al., 1999) for less than 15 minutes old contrails increasing to 1–3 km for less than 30 min old contrails (Freudenthaler et al., 1995).

Little is known on chemical contrail properties and the uptake of HNO_3 in contrails. Gaseous HNO_3 has been detected in aircraft exhaust at cruise (Arnold et al., 1992; Tremmel et al., 1998). High levels of HNO_3 are expected in the young aircraft exhaust, which can be rapidly taken up by aircraft aerosol and contrails (Kärcher, 1996). At temperatures below 205 K, the gas phase HNO_3 is significantly reduced due to uptake in contrail ice crystals as detected for the contrail of a WB-57 at 14 km altitude (Popp et al., 2004). Nitric acid in the ice crystals at these temperatures might be present in the form of NAT (Gao et al., 2004; Voigt et al., 2008). Schäuble et al. (2009) show first quantitative measurements of the HNO_3 content in contrails. They report enhanced ice phase HNO_3 fractions in persistent contrails compared to natural cirrus clouds (Voigt et al., 2006, 2007; Kärcher and Voigt, 2006; Krämer et al., 2008). When the age of the contrail is known, the contrail ice crystals can serve as atmospheric laboratory for the study of the temporal evolution of trace gas uptake in ice crystals (Schäuble et al., 2009; Kärcher et al., 2009b).

Given the sparsity of in-situ measurements of young contrails, measurements of chemical and optical contrail properties are of importance to better quantify their chemical and radiative impact on the atmosphere. Hence we report here on a new set of contrail observations above Europe. The measurements have been performed with the DLR research aircraft Falcon during the CONCERT (CON-trail and Cirrus ExperimenT) campaign from 22 October to 20 November 2008. The first two weeks of the campaign (CONCERT-CHEMISTRY) were focussed on heterogeneous chemistry on contrails, cirrus and aerosol, while the second part (CONCERT-2-CONTRAIL) concentrated on microphysical and optical contrail properties. During the CONCERT campaign, cirrus clouds were encountered during 10 mission and 2 instrument test flights for almost 2 hours. 22 contrails from 11 different aircraft were sampled in total for 1.7 hours at temperatures from 213 to 229 K (27 and 28 October and 2, 17, 19 and 20 November 2008). Here we use the contrail observations for model validation. Further, we investigate in detail the microphysical properties from a CRJ-2 contrail and analyze the chemical processing of SO_2 and NO in the engine and the young contrail. Large scale aircraft NO emissions were predicted from global model simulations and measured during two mission flights on 31 October 2008.

In addition, a sulfate aerosol layer was detected in the lowermost stratosphere at altitudes between 8.6 and 11.6 km (28 and 31 October 2008). The layer most likely originated from the eruption of the Aleutian volcano Mt. Kasatochi on 7/8 August 2008, injecting 1.5 Mt SO_2 into the high latitude stratosphere. We discuss the chemical processing of sulfur and halogens species in the aged volcanic plume and derive the e-folding lifetime of SO_2 in the northern latitude lowest stratosphere.

2 Aircraft instrumentation

During the CONCERT campaign, the DLR research aircraft Falcon was equipped with a set of instruments to detect microphysical particle properties and trace gas composition in the UTLS.

2.1 The Forward Scattering Spectrometer Probe FSSP300

Particle size distribution and number density were detected with a forward scattering spectrometer probe (Schröder et al., 2000). The FSSP300 detects light scattered by particles in the diameter range 0.45–17.7 μm . The particles were grouped into size channels according to T-matrix calculations by Borrmann et al. (2000) assuming aspherical particles with an aspect ratio of 1:2 composed of pure ice with a refractive index of 1.31. Adjustments of the mean channel size were made to account for instrumental manufacturing differences. Further, FSSP channels 8+9, 10+11, 12+13+14+15, 16+17,

18+19, 20+21+22+23+24, 25+26, 29+30 were grouped to increase the counting statistics particularly in channels with a small channel widths. This data evaluation method proved to achieve best agreement between the FSSP data and the particle scattering phase function detected with the Polar Nephelometer.

2.2 The 2D-C Probe, the Polar Nephelometer (PN) and the Cloud Particle Imager (CPI)

The particle size distribution of large particles (100 μm to 1 mm) was measured with a 2D-C probe, the particle shape (2.3 μm pixel size) with a cloud particle imager (CPI) and the scattering phase function of cloud particles (3 μm to 1 mm) with a Polar Nephelometer (PN). The method of data processing, the reliability of the instruments and the uncertainties of the derived microphysical and optical parameters have been described in detail by Gayet et al. (2009).

The 2D-C probe provides information on crystal size and shape for the size range 100 to 800 μm . The method of data processing used in this study has been described in detail by Febvre et al. (2009). The “reconstructed” method using partial images, was considered for the calculations and the sampling surfaces were derived according to Heymsfield and Parrish (1978). In order to improve the statistical significance of low particle concentrations, a 5-second running mean was applied. Irregular ice particles were the most predominant crystals sampled during the CONCERT experiment. Therefore, the bulk parameters were calculated assuming the surface-equivalent diameter relationships (Heymsfield, 1972; Locatelli and Hobbs, 1974). As the sensitivity of the probe to small particles decreases with airspeed (Lawson et al., 2006), particles smaller than about 100 μm may not be detectable with the 2D-C at a Falcon airspeed of 200 m/s.

The CPI records very high resolution (2.3 μm) digital images of cloud particles as they pass through the sample tube of the instrument. It casts an image of the particle on a solid-state CCD camera by freezing the motion of the particle using a 25 ns pulsed, high-power laser diode. Upstream lasers precisely define the depth-of-field so that at least one particle in the image is almost always in the focus. This eliminates out-of-focus sizing errors that have plagued the conventional 2D imaging probes. The method of data processing leads to derive the particle size distribution, shape and extinction coefficient.

The PN is designed to measure the optical and microphysical parameters of clouds containing either water droplets or ice crystals or a mixture of both over a size range from a few micrometers to about 800 microns diameter. The probe measures the scattering phase function of an ensemble of cloud particles which intersect a collimated laser beam near the focal point of a paraboloidal mirror. The light scattered from polar angles from $\pm 3.49^\circ$ to $\pm 169^\circ$ is reflected onto a circular array of 54 photodiodes. The signal processing electronics and computer storage can provide one measurement of the

scattering phase function every 10 ms. The optical properties (extinction coefficient, asymmetry factor, backscattering coefficient) are derived from the measured scattering phase functions. The particle size spectra and subsequent derived quantities such bulk and size parameters are retrieved using an inversion technique. The particle phase discrimination (water droplet/ice particles) can be derived from the shape of the scattering phase function.

2.3 The Cloud Imaging Probe (CIP)

During the last flight the 2D-C has been exchanged with a Cloud Imaging Probe (CIP), which is a new version of the 2D-C. The CIP measures the size and shape of particles passing through its collimated laser beam, from 125 μm to 1600 μm with a resolution of 25 μm . It is capable to detect particle concentrations up to 100 cm^{-3} at airspeeds up to 200 m/s. It uses a fast 64-element photodiode array to generate 2-dimensional images of the particles. The uncertainty in the CIP particle number concentration is mainly determined by the uncertainty in the sample volume of 20%. At low particle concentrations also counting statistics have to be taken into account. The uncertainty in the particle size decreases with particle size and is $\pm 25\text{ }\mu\text{m}$ for particles $> 125\text{ }\mu\text{m}$ in diameter (De Reus et al., 2009).

2.4 The Spectral Modular Airborne Radiation measurement sysTem (SMART)-Albedometer

Spectral upwelling radiances and irradiances were measured with the Spectral Modular Airborne Radiation measurement sysTem (SMART)-Albedometer (Wendisch et al., 2001). During the CONCERT campaign, an optical inlet with a viewing angle of 1.5° designed to measure radiances as well as an irradiance optical inlet (half-dome shape) were each connected to two plain-grated spectrometers, respectively. The optical inlets were placed in the back part of the fuselage of the aircraft. The spectrometers operate in the wavelength ranges 350 to 1050 nm and 900 to 2200 nm with spectral resolutions (Full Width at Half Maximum, FWHM) of 2 to 3 nm and 9 to 16 nm, respectively. The temporal resolution of the measurements is in the order of 0.5 to 3 s, corresponding to a spatial resolution of 100 to 600 m at an aircraft velocity of 200 m/s.

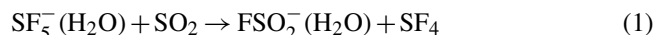
2.5 The Aerosol Mass Spectrometer (AMS)

The chemical composition and size distribution of submicron aerosol particles were determined by an Aerodyne Compact Time-of-Flight Mass Spectrometer (C-ToF-AMS). The instrument samples aerosol through an aerodynamic lens system which focuses the particle beam onto a vaporizer operated at 600°C . Before reaching the vaporizer, the particles pass through a time-of-flight region in a vacuum chamber that allows for particle size determination. The vapor is ionized by 70 eV electrons, and the generated ions are analyzed

in a time-of-flight mass spectrometer. A detailed description of the C-ToF-AMS is given in Drewnick et al. (2005) and Canagaratna et al. (2007). During CONCERT-CHEMISTRY particulate sulfate, ammonium, nitrate, chloride and carbonaceous matter were determined for STP conditions with a time resolution of 10 s. The respective detection limits were 0.02, 0.13, 0.02, 0.03, and $0.11 \mu\text{g m}^{-3}$ (STP) on campaign average. To compensate for the decreasing mass flow into the instrument with altitude, a pressure controlled inlet (PCI) has been installed in front of the standard AMS inlet system, also guaranteeing isokinetic sampling from the aircraft's aerosol inlet. A thorough description of the inlet system is given in Schmale et al. (2010). This Falcon-specific set-up allowed for sampling of particles in the size range between 84 and 735 nm vacuum aerodynamic diameter.

2.6 The Ion Trap Chemical Ionization Mass Spectrometer (ITCIMS)

An ion trap chemical ionization mass spectrometer (ITCIMS) (Fiedler et al., 2005; Speidel et al., 2007; Fiedler et al., 2009) was for the first time operated with SF_5^- reagent ions for the detection of HCl, HNO_3 , SO_2 and HONO in the UTLS. A detailed description of the instrument is given in Jurkat et al. (2010). In the flow reactor SF_5^- reacts with the four trace gases via a fluoride transfer, here exemplarily shown for reactions with SO_2 and with HONO:



The water clusters are removed through collision with helium injected into the ion trap. Mass spectra from 15 to 170 atomic mass units (amu) with a resolution of 0.3 amu were sampled during a trapping time of 200 ms and averaged over five spectra resulting in an overall time resolution of 1.6 s. This allowed for the fast and sensitive detection of SO_2 and HONO in young exhaust plumes of commercial airliners and for measurements of SO_2 , HCl and HNO_3 in an aged stratospheric plume of volcanic origin. Detection limits for HONO and SO_2 for 1.6 s time resolution were 72 and 67 pmol/mol (pptv). Detection limits for HCl and HNO_3 were 22 and 36 pmol/mol respectively with a running mean over 20 spectra (approx. 32 s). A HNO_3 in-flight calibration was performed with a nitric acid permeation source, while HCl, SO_2 and HONO were calibrated in the laboratory. The HONO calibration of the ion trap mass spectrometer was carried out with a LOPAP instrument (Heland et al., 2001) proving the pseudo first order kinetic reaction (2) given above. The calibration factor accounting for the rate coefficient, the ion molecule reaction time, the dilution and possible wall effects for SO_2 and HONO were 20 and 10 nmol/mol (ppbv) with an instrumental error of $\pm 25\%$ and 40% , respectively.

2.7 The NO_y instrument

The NO_y instrument is described in detail by Ziereis et al. (2004) and Voigt et al. (2005). During the first two weeks of the campaign, it was configured to detect mainly gas phase reactive nitrogen $\text{NO}_{y,g} (= \text{NO} + \text{NO}_2 + \text{HONO} + \text{HNO}_3 + 2\text{N}_2\text{O}_5 + \text{PAN} + \dots)$ with a backward facing inlet and total $\text{NO}_{y,t} (= \text{NO}_{y,g} + \text{NO}_{y,\text{part}})$ with a forward facing inlet. Thereafter it monitored mainly gaseous $\text{NO}_{y,g}$ and nitric oxide (NO) with two rear facing inlets.

In the rear facing inlets, particles larger than the cut-off size of $d_{50} = 0.3 \mu\text{m}$ are inertially stripped from the sampled air, so that predominantly gas phase $\text{NO}_{y,g}$ is measured (Voigt et al., 2007). Particles are evaporated in the forward facing inlet releasing gaseous $\text{NO}_{y,g}$. $\text{NO}_{y,g}$ is reduced to NO with CO in a heated gold converter and the chemiluminescence reaction of NO with ozone in the infrared is detected with a photo detector. The NO_y instrument has a detection limit of 1 and 5 pmol/mol for NO and NO_y respectively and an accuracy of $\pm 8\%$.

2.8 The Fast In Situ Hygrometer (FISH)

Water vapor has been measured with the closed path Fast In situ Stratospheric Hygrometer (FISH) (Zöger et al., 1999; Schiller et al., 2008) using the Lyman- α photofragment fluorescence technique. H_2O mixing ratios between 0.5 and 1000 ppmv have been detected with FISH with a time resolution of 1 s and an overall accuracy of 6% or at least 0.3 ppmv. FISH is calibrated regularly before, during and after the field campaign using calibration bench including a frost point hygrometer (MBW DP30) as reference. During CONCERT, air is probed by a backward facing inlet tube, i.e. FISH measured gas phase water (H_2O_g). Including the uncertainty of 0.5 K for the temperature measurements, the relative humidity with respect to ice (RH_I) can be determined with an uncertainty of $\pm 15\%$.

2.9 The frost point hygrometer

In addition, water vapor has been detected with the cryogenic frost point hygrometer CR-2 (Buck Research Instruments, LLC). The instrument measures the temperature of a mirror carrying a thin frost layer that is maintained in equilibrium with the ambient water vapor. The bulk reflectivity of the mirror is measured with an optical detector. It initializes a control circuit to regulate the mirror temperature such that the bulk reflectivity and hence the condensate layer remain constant. The mirror temperature is then equal to the ambient frost point temperature and the water vapor mixing ratio can be calculated using the inverse Clausius-Clapeyron equation.

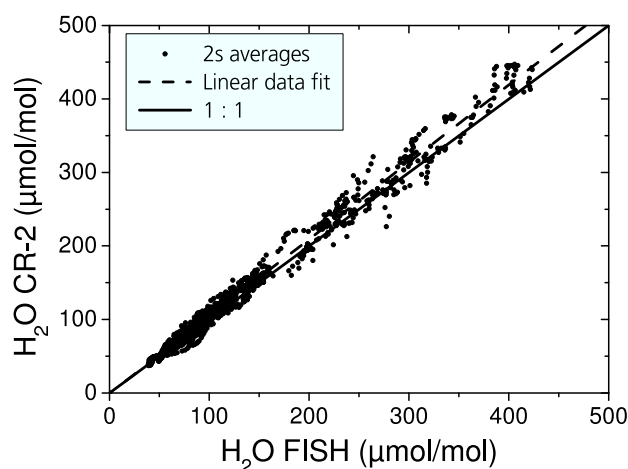


Fig. 1. Correlation of water vapor mixing ratios detected with the frost point hygrometer CR-2 and the Lyman α hygrometer FISH on 20 November 2008 above Germany. The 1:1 line and the linear regression are shown.

The frost point hygrometer has been calibrated to a high accuracy laboratory frost point hygrometer (MBW373 Calibration) during and after the campaign.

The uncertainty of the water vapor mixing ratio is between $\pm 3\%$ and $\pm 5\%$ for water vapor mixing ratios between $340 \mu\text{mol/mol}$ and $40 \mu\text{mol/mol}$. This uncertainty does not include time-dependent oscillations of the mirror temperature induced by the control circuit after steep gradients in water vapor. Hence the response time of the frost point hygrometer is on the order of several seconds to one minute depending on the water vapor gradient. The CR-2 instrument has been operated for the first time on the Falcon during the CONCERT mission, therefore we show a comparison of the water vapor mixing ratios detected with the CR-2 instrument and Lyman- α hygrometer.

Figure 1 shows the correlation of the water vapor mixing ratios detected with the CR-2 and the FISH instrument during a flight on 20 November 2008 above Germany. The mean water vapor mixing ratio measured by the frost point mirror was 4% higher than the mixing ratio observed by the Lyman α instrument and lies within the errors of both instruments. The standard deviation of the ratio of the water vapor mixing ratios (CR-2/FISH) is $\pm 7\%$.

3 Flight planning and meteorological forecasts

Research flight planning was based on deterministic meteorological forecasts from the European Centre for Medium-Range Weather Forecasts (ECMWF) and on chemical forecasts computed by the MATCH-MPIC model (Lawrence et al., 2003). In addition to the forecast quantities provided by ECMWF, we derived several campaign-specific forecast quantities, for instance, the Schmidt-Appleman cri-

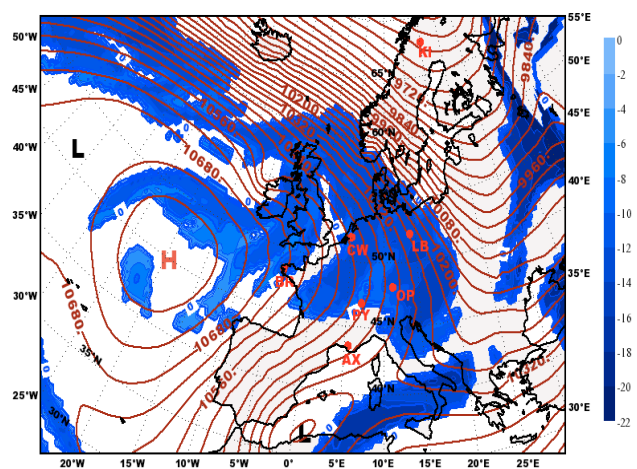


Fig. 2. 36 hour forecast of the Schmidt-Appleman criterion at 250 hPa, valid on 19 November 12 UTC. Shown is the difference between the ambient temperature (K) as forecasted by the ECMWF and the threshold temperature for contrail formation as computed from the Schmidt-Appleman theory. Values are only plotted in regions where the relative humidity is larger than 80%, that is contrails that form in such a region have a high probability to be persistent. Red contours show the geopotential height (m).

terion (Schumann, 1996) to predict regions with a high likelihood of contrail occurrence. Furthermore, the trajectory code LAGRANTO (Wernli et al., 1997) was run on-demand on the ECMWF predicted wind fields in order to estimate how contrails in the flight target regions would be advected.

ECMWF forecasts were visualized with the Metview system (Bonifacio, 1999) and made available to the campaign scientists through a web-based user interface. For fast on-demand LAGRANTO computations, a Python-based user interface was developed. With the information given by the forecast system, we were able to identify target regions with atmospheric conditions well suited to fulfil the mission objectives. Figure 2 shows a plot of the 36 h forecast of the Schmidt-Appleman criterion combined with regions of high relative humidity, valid on 19 November 2008, 12:00 UTC, at 250 hPa. The predictions indicated good conditions for contrail occurrence over Germany and the North Sea. The day was subsequently chosen for targeted contrail observations.

4 Flight survey

During the CONCERT campaign, the Falcon performed 12 mission flights on 27, 28, 29 and 31 October and on 2, 7, 17, 19 and 20 November 2008 including 3 return flights on 31 October and on 17 and 19 November. The flight paths from all flights are shown in Fig. 3. Extensive contrail measurements were performed on 2, 17, 19 and 20 November and contrails were also observed on 28 and 29 October, but without FSSP data. The flights on 27 October and on 7 and

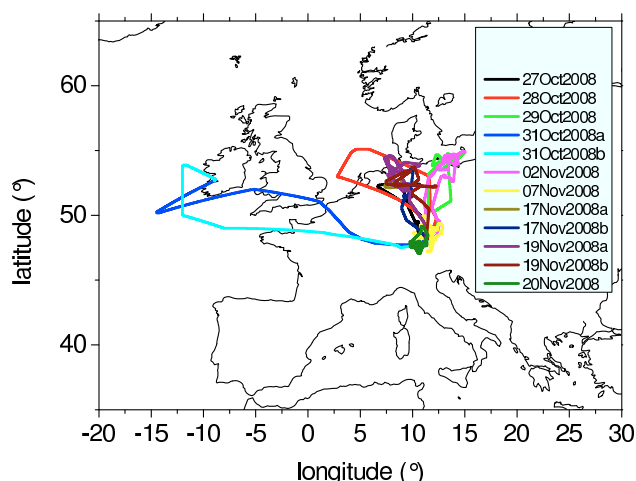


Fig. 3. Trajectories of 12 mission flights performed during the CONCERT (CONtrail and Cirrus ExpeRiment) campaign in October/November 2008 over Western Europe. The flight dates are given in the legend.

20 November also concentrated on the detection of cirrus clouds. The 31 October return flight from Oberpfaffenhofen to Shannon had the aim to detect regional aircraft emissions from the Northern Atlantic flight corridor. Sampling a stratospheric intrusion was the main goal for the flights on 28 and 31 October.

5 Detection of young contrails

Microphysical and optical properties of 22 young contrails were detected by the instruments aboard the Falcon on 8 mission flights. Contrails from 11 different aircraft types were detected, including an A380, several A340s and B767s and B737s, an A319, a CRJ-2 and a Fokker20. The contrails were probed at altitudes between 8.5 to 11.6 km and at temperatures between 213 and 229 K (Fig. 4). Contrails were identified from a simultaneous increase in the extinction $>0 \text{ km}^{-1}$ and the NO mixing ratio $>0.2 \text{ nmol/mol}$, latter representing the upper limit for the upper tropospheric NO levels. Further the threshold temperature $T < 240 \text{ K}$ was used to exclude lower tropospheric cloud observations. In total 1.7 h of contrail measurements were achieved. The RH_i in the contrails ranged between 122 and 55%, with 80% of the observations concentrating between 105 and 75% RH_i . Figure 4 (bottom panel) shows that the clear sky RH_i is also mostly below 100% with the most frequent observations between 70% and 100%. Since air from the environment is continuously entrained into the contrails, the slightly sub-saturated conditions in the contrails might portray the clear sky conditions. Including the error in the temperature measurements, the RH_i has an accuracy of $\pm 15\%$.

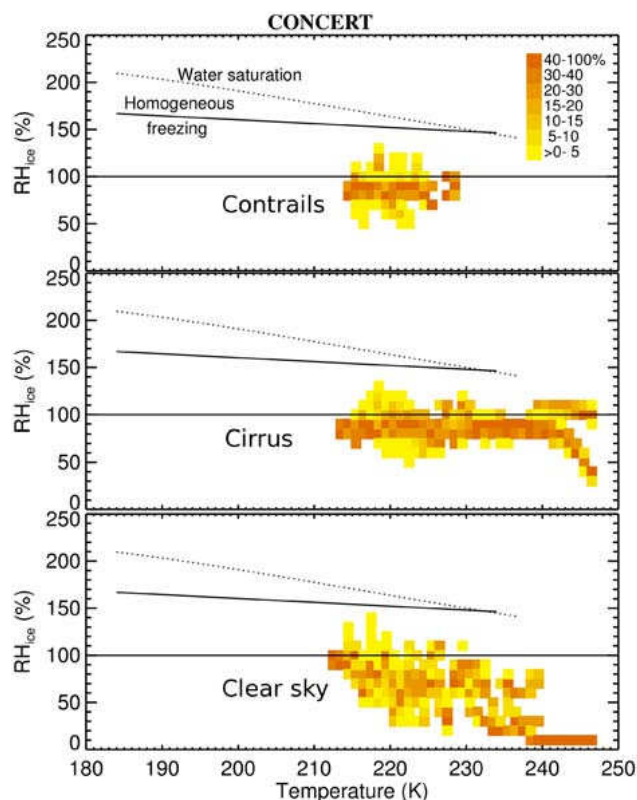


Fig. 4. Frequencies of occurrence of relative humidities over ice versus temperature in contrails (upper panel), cirrus clouds (middle panel) and clear sky (lower panel) observed during the CONCERT campaign. The thin solid line represents the homogeneous freezing threshold, the dotted line refers to water saturation. The data are sorted in 1 K temperature bins. In each 1 K bin the sum of all data is 100%.

Contrails were mainly probed in the top region or above visible cirrus clouds, as this sampling strategy was found to be very effective. The contrail sampling strategy was the following: Predictions of the IWC from the ECMWF were used to send the aircraft into the top of a cirrus cloud. The Falcon further ascended above the cloud till contrails were observed by the pilots. The contrail formation altitude was communicated to German Air Traffic Control (DFS), and commercial airliners crossing that region were asked to change their flight altitude to contrail formation altitudes. Thereafter the Falcon was directed behind the airliners and contrails were probed at 10 to 86 nautical miles distance corresponding to contrail ages of 55 to 572 s. During the contrail samplings, the pilots of the commercial aircraft communicated their distance to the Falcon as well as aircraft parameters such as velocity, weight and fuel flow. These parameters were then used to derive emission indices for NO_x at flight altitude (Schulte et al., 1997; Döpelheuer and Lecht, 1999).

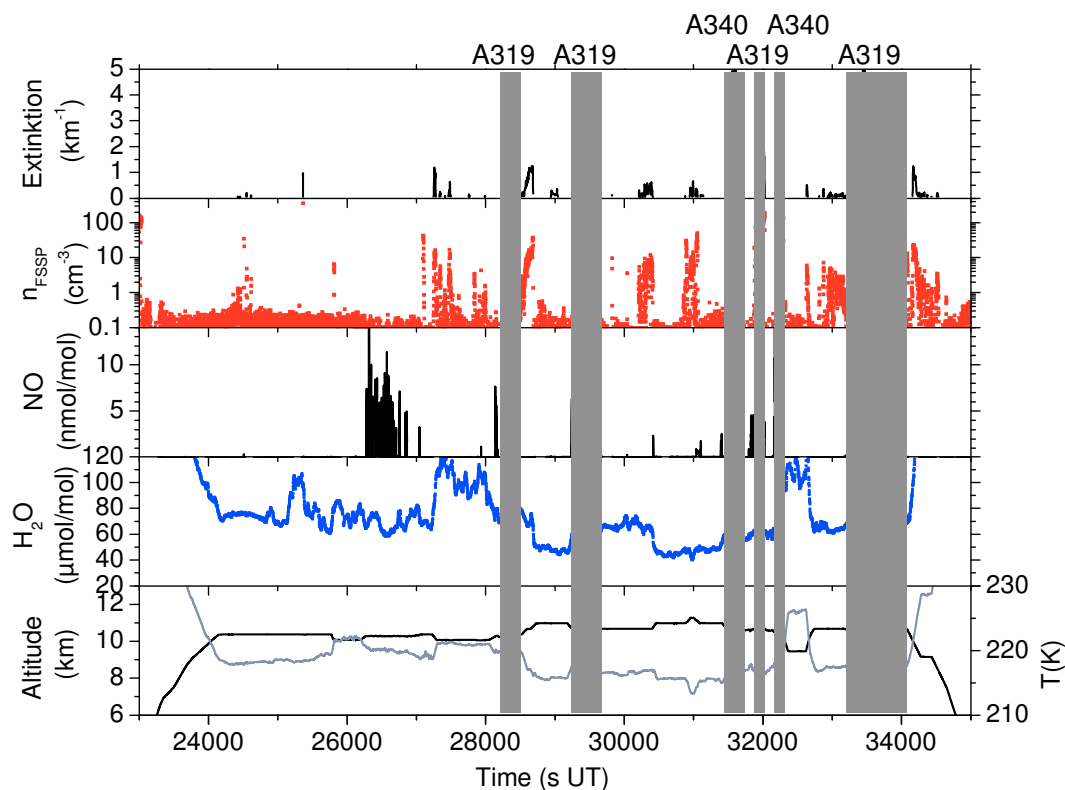


Fig. 5. Contrail flight A on 19 November 2008 above Germany from Oberpfaffenhofen to Hamburg. The flight altitude, temperature, and the H_2O and NO mixing ratios as well as particle number density detected with the FSSP and the extinction for particles larger than $3\text{ }\mu\text{m}$ are shown. Periodic gaps in the NO data are due to calibration events. Sequences of contrail encounters of an A319 and an A340 are marked in gray.

5.1 Survey of two contrail flights on 19 November 2008

On 19 November 2008, extensive contrail measurements were performed with the Falcon. We detected 10 contrails above Germany during a return flight from Oberpfaffenhofen to Hamburg and back. During the first flight shown in Fig. 5, we frequently chased an A319, which was a charter aircraft flying specifically for the contrail measurements. The exhaust of the A319 without contrail was detected in the first encounter, thereafter contrails from the A319 were probed four times with the longest contrail encounter of 13 minutes. In addition, the contrail from an A340 was sampled twice. During the return flight, we chased four commercial aircraft and detected the contrails of a B767, a CRJ-2, an A380 and an A320 over Germany as shown in Fig. 6.

Contrail encounters are marked by fast simultaneous increases in the NO mixing ratio, the particle number density detected with the FSSP and the extinction detected with the PN. We attribute contrails to measurements where the extinction is $>0\text{ km}^{-1}$ and the NO mixing ratio is $>0.2\text{ nmol/mol}$. In addition, the contrails were filmed by the camera in the cockpit of the Falcon.

Highest NO mixing ratios and extinctions were detected in the contrail of the A380. The contrail was observed for 756 s. NO values up to 58 nmol/mol were observed and the extinction of particles with $d > 3\text{ }\mu\text{m}$ increased to 5.4 km^{-1} in the A380 contrail. Enhanced particle concentrations ($0.4 < d < 17.7\text{ }\mu\text{m}$) up to 537 cm^{-3} were detected. Average values in the A380 contrail were lower with a particle concentration of 179 cm^{-3} , NO mixing ratio of 6 nmol/mol and extinction of 1.6 km^{-1} .

The age of the contrails has been calculated based on the positions of the Falcon and the contrail producing aircraft taking into account the drift of the contrail at the measured wind speed. The ages of the detected contrails were between one and three minutes for the A319, the A340, the B767 and the CRJ-2. The A380 contrail with ages of 1 to 5 min was detected for 16 minutes. Having left the contrail of the A380, the contrail of an A320 with ages of 8 to 10 min was sampled, representing the oldest contrail observed during the CONCERT campaign. Given contrail ages of less than 10 minutes, all contrails are expected to be line shaped. Hence the contrail measurements were performed in the vortex and the early dispersion regime.

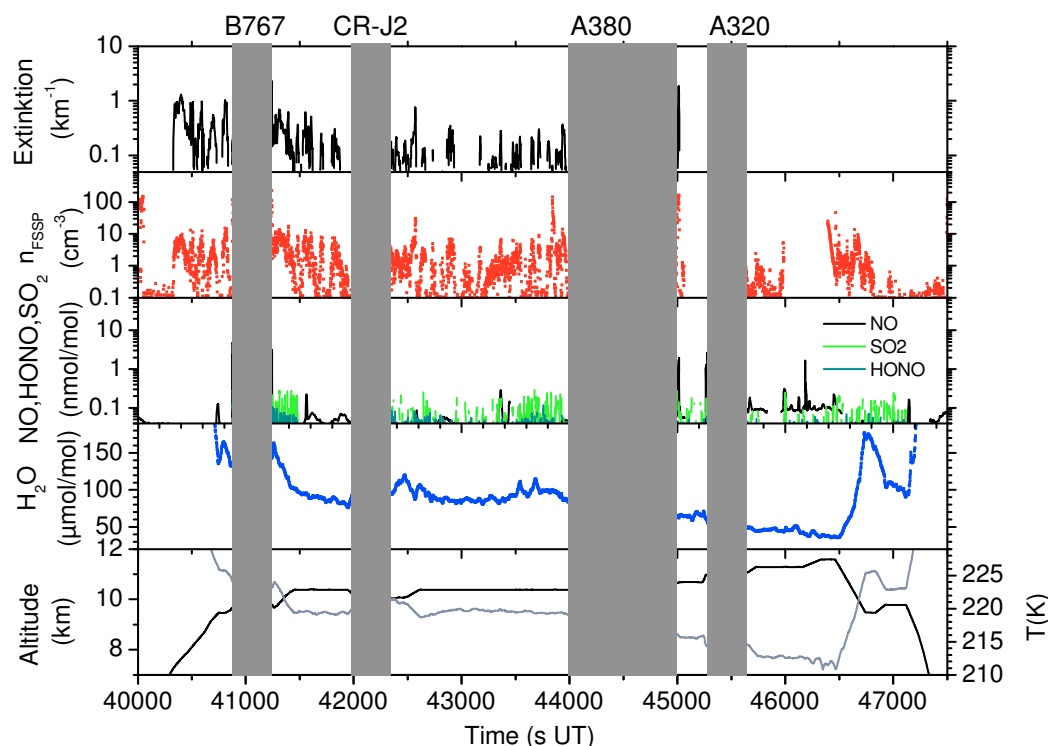


Fig. 6. Contrail flight B on 19 November 2008 above Germany from Hamburg to Oberpfaffenhofen. The flight altitude, temperature, and the H_2O , NO, SO_2 and HONO mixing ratios as well as the particle number density detected with the FSSP and the extinction for particles larger than $3\text{ }\mu\text{m}$ are shown. Sequences of contrail encounters of a B767, a CRJ-2, an A380 and an A320 are marked in gray. Gaps in the NO, SO_2 and HONO mixing ratios are related to instrument calibrations. The PN stopped working at 45050 s UT and there is a gap in the FSSP data between 45050 s and 45410 s UT.

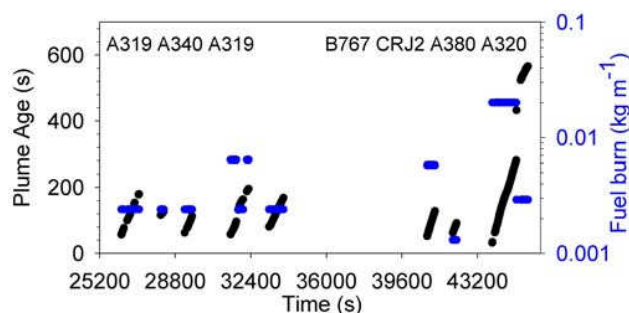


Fig. 7. Plume ages versus Falcon flight time and fuel consumption of the aircraft observed during two flights on 19 November 2008.

6 Contrail modeling

A new “Contrail Cirrus Prediction Tool” (CoCiP) has been developed to simulate contrail cirrus resulting from a single flight as well as from a fleet of cruising aircraft, flight by flight, regionally or globally (Schumann, 2010). The method predicts contrail cirrus for given aircraft, traffic and weather prediction data. The method describes the life cycle of individual contrails using a Lagrangian Gaussian plume

model, following early concepts (Schumann and Konopka, 1994; Schumann et al., 1995; Konopka, 1995). The turbulent mixing of the plume is mainly controlled by shear but also controlled by turbulent diffusivities; these parameters were selected such that the computed dilution factor (i.e. the mass of the plume per mass of fuel burnt) fits available measurements (Schumann et al., 1998) and LES results (Dürbeck and Gerz, 1996). Contrails are initiated when the Schmidt-Appleman criterion is satisfied (Schumann, 1996). The evolution of the individual contrails is computed using wind, temperature, humidity, and ice water content from numerical weather prediction (NWP) output. In our case we use the operational forecast data of the Integrated Forecast System of the ECMWF. This model computes the humidity, allowing for ice supersaturation in air masses without and with cirrus clouds, in fair agreement with observations (Tompkins et al., 2007). The plume trajectory follows horizontal and vertical winds using established methods (Stohl et al., 2001). The ice phase is modelled as a function of bulk contrail ice properties (ice mass and ice particle number). It is assumed that ice particles form on emitted soot particles. The ice water content is computed as a function of mixing and ambient humidity. The model assumes ice saturation inside the contrail. Contrails in

subsatuated air masses evaporate at time scales depending on the emitted mass of water vapor, the degree of subsaturation and plume dilution. In supersaturated air masses, the contrail ice water content grows by uptake of humidity mixed into the plume from ambient air. The final life time of the contrails in the model is controlled by ice particle loss processes (turbulent mixing, aggregation, sedimentation, radiative heating). These aspects will be described elsewhere and are not important for the present study with young contrails. The climate impact of the contrails (radiative forcing and radiative energy gain per flight distance with contrail) is computed as a function of the contrail properties (width, length, life time, optical depth, temperature, effective radius, optical depth of ambient cirrus above the contrail) using radiative fluxes without contrails from NWP output (Schumann, 2010).

The contrail properties are strongly dependent on the ambient meteorological field (in particular humidity, shear, stratification, and upward motion). In addition the computed contrail properties are aircraft depending (Schumann, 1996). The fuel properties and the propulsion efficiency of the aircraft influence the threshold temperature of contrail formation (Busen and Schumann, 1995). The amount of water vapor emitted in the early contrail contributes to the initial ice water content and hence the life time of the contrail in unsaturated ambient air. The number of soot particles emitted controls the number of ice particles forming in the young contrail (Kärcher and Yu, 2009). Besides fuel consumption, the emission indices (EI) control the computed plume concentrations. The emission index for NO_x was determined as in Döpelheuer and Lecht (1999) based on aircraft and engine type and the fuel consumption rates reported by the aircraft pilots. The mass, speed and span width of the individual aircraft determine the downward sinking of the wake vortex (Holzapfel, 2006). The shear driven plume dispersion is strongly enhanced for large wake vortex sinking depths (Schumann et al., 1995). The actual air traffic data (radar observed aircraft type and aircraft positions versus time) were provided for the CONCERT days by the Deutsche Flugsicherung (DFS). The position of the Falcon, measured onboard was found to be in close agreement (50 m laterally, 10 m vertically) with the DFS observations.

6.1 Validation of CoCiP model results with CONCERT observations of young contrails

From the position of the Falcon as a function of time, and the ECMWF forecast wind field (wind speed of order 40 m/s), the model computes the age of the exhaust plume at the time of measurement. The data were accurate enough to decide when the Falcon was in or near the exhaust of aircraft for young plumes (a few minutes) to an age dependent accuracy of order 300 m horizontally, and 20 m vertically. Figure 7 shows the plume ages (between 60 and 600 s) for the two Falcon flights of 19 November 2008. Also plotted is the fuel

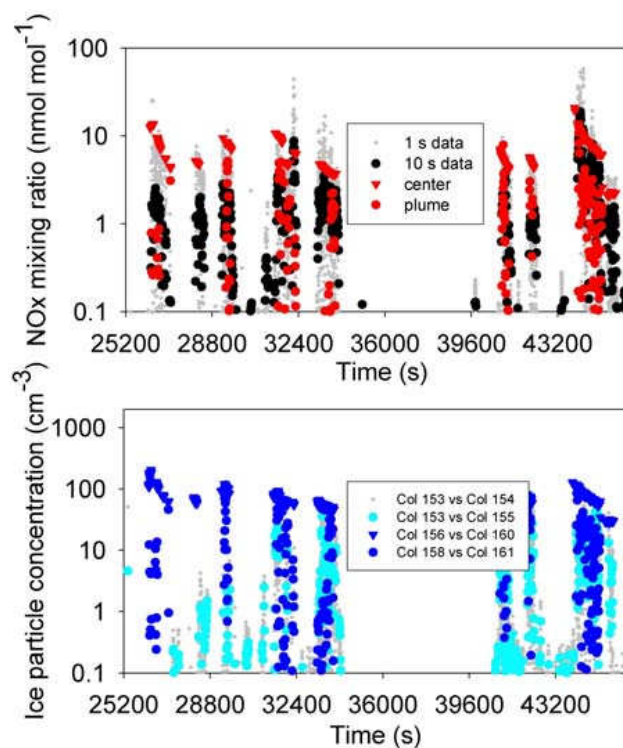


Fig. 8. Comparison of CoCiP model results to contrail observations during two flights on 19 November 2008. Upper panel: NO_x mixing ratio versus Falcon flight time. Measured data (grey and black dots) and computed results (red circles: along computed flight path; red triangles: on the plume axis). Lower Panel: Concentration of ice particles larger than $1\ \mu\text{m}$ versus Falcon flight time. Measured data (grey and cyan dots) and computed results (blue circles: along computed flight path; blue triangles: on the plume axis).

consumption of the aircraft, which varies by a factor of 20 between the various aircraft. From a comparison of the computed and measured times of entering the exhaust plume we could see that the model correctly predicts the sinking of the initial wake vortices, up to 270 m for the heaviest aircraft, to within about 20 m vertical accuracy. Figure 8 shows the NO_x mixing ratio versus Falcon flight time. Both 1-s data and 10-s running averages of the measured data are plotted. The latter are commensurable with the model results which do not resolve the turbulent details of the aircraft wake. In the model, the simulated Falcon track does not always lead through the rather narrow (order of 100 m) exhaust plume. Hence the model tends to underestimate the maximum measured exhaust concentrations. However, the model can also be applied to compute the plume properties under the assumption that the Falcon at the time of measurement would be flying exactly in the center of the Gaussian plume. In this way, the model provides the maximum exhaust values that are consistent with the Gaussian plume. The NO_x data reach up to 50 nmol/mol, i.e., far above the background concentrations. Maximum concentrations were measured in the plume

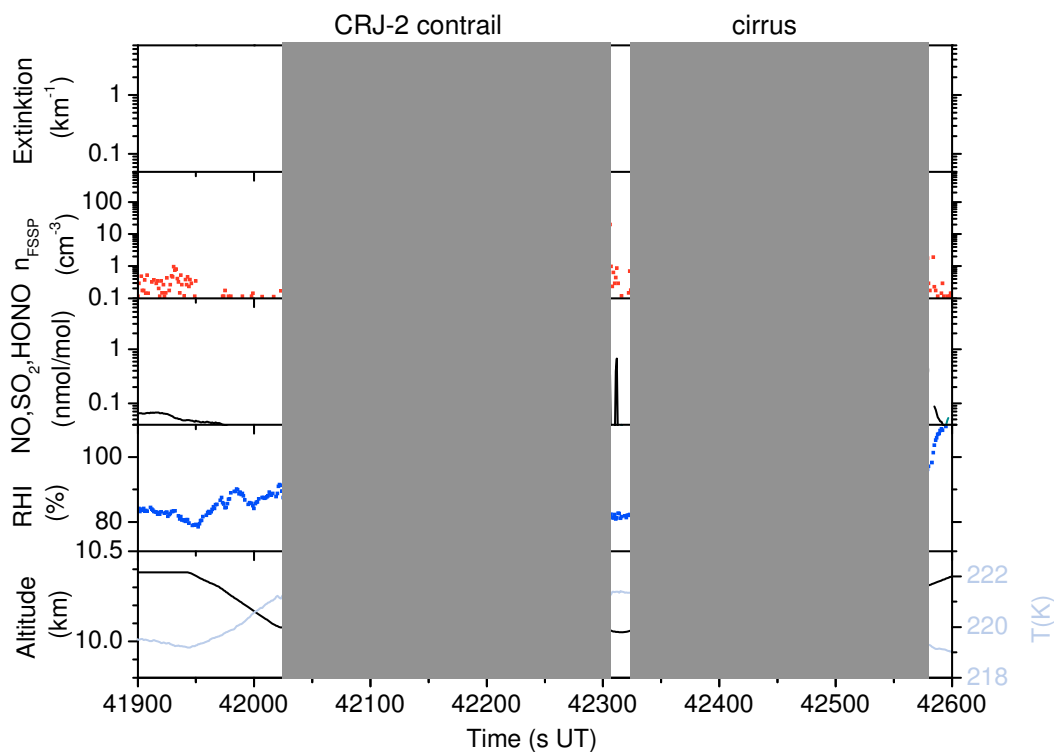


Fig. 9. Observations in the contrail from a CRJ-2 on 19 November 2008 above Germany. The flight altitude, temperature, RH_I, and trace gas mixing ratios of NO, SO₂ and HONO as well as particle number density detected with the FSSP and extinction for particles larger than 3 μm are shown. The contrail (cirrus) sequence is marked in dark (light) gray.

of the biggest aircraft (A380). The maximum concentrations decrease with plume age because of dilution. Figure 8 shows that the model provides NO_x mixing ratios that are within the scatter of the measured data. The model gives different NO_x levels in the various plumes which differ because of different plume ages, dilution, fuel consumption and emission indices. For comparable plume ages, the NO_x concentration is larger for heavier aircraft, mostly with higher emission indices. The results indicate that the model parameters were suitably selected to correctly simulate the dilution in the sheared and turbulent ambient atmosphere. It should be noted that this agreement was obtained after the emission indices had been adjusted for the various aircraft/engine types. The results indicate quite large NO_x emission indices in some cases up to 27 g/kg. Figure 8 shows similar results for the ice crystal number concentration in the contrail. The model results and the measurements (for ice particles larger than 1 μm) agree within the scatter of the data. The concentrations reach as high as 100 cm⁻³ in the contrails. During the first flight the model tends to overestimate the ice particle concentrations and predicts very small ice water content because of the low ambient humidity. The agreement with observations is better during the second flight at higher ambient humidity. The computed ice particle concentration is linearly dependent on the number of soot particles emitted with the exhaust. Hence, this is also a check for the soot number emission indices used

(2.4–3.4 $\times 10^{15}$ kg⁻¹). The absolute level of the model results is sensitive to the plume dilution and the loss of ice particles in the adiabatically sinking and thus heating wake core contrail (Sussmann and Gierens, 1999; Unterstrasser and Gierens, 2010; Lewellen and Lewellen, 2001). Again the agreement is such that it supports the basic validity of the model.

CoCiP seems to compute realistic ice particle concentrations. There were no indications of additional ice formation beyond soot induced ice particles. In particular, there are no indications for large contributions from volatile material contributing to ice nucleation for these cases with temperatures of 210 to 222 K. Hence, it seems that the rather complex issue of ice formation in fresh contrails can be explained at accuracy comparable with measurements with a Gaussian plume model and with bulk ice physics and nucleation controlled by the soot particles in the contrail.

The model is computationally efficient enough to be used in forecast mode for operational route optimization with respect to minimum climate impact. Further model tests are underway to explain observed aviation induced cirrus cover diurnal changes over the North Atlantic. Further, the model can be used to evaluate the radiative forcing induced by contrail cirrus formation globally.

7 Observation of the contrail from a CRJ-2 aircraft

Below we discuss in detail the contrail sampling from the CRJ-2 on 19 November 2008. The aircraft was chased between 11:40:23 and 11:45:06 UT above Germany (Fig. 9). The contrail was detected for almost 4.5 minutes at an altitude of 10.02 to 10.11 km and at temperatures near 221.4 ± 0.1 K. Elevated extinction for particles larger than $3 \mu\text{m}$ up to 3.3 km^{-1} , particle number concentrations up to 406 cm^{-3} and NO mixing ratios up to 6 nmol/mol were measured. Average values in this contrail are 0.8 km^{-1} in the extinction, 125 cm^{-3} in the particle concentration and 1.7 nmol/mol in the NO mixing ratio. Peak SO_2 and HONO mixing ratios of up to 0.69 and 0.26 nmol/mol were detected. In the contrail the RH_i increased to values between 82 and 95% and was below these values direct before and after the encounter. The contrail was detected at an age of 61 to 122 s and hence was probed in the vortex regime.

The mean particle surface area and volume distributions detected with the FSSP and the 2D-C in the contrail are shown in Fig. 10. The gap in the particle data between 17.7 and $100 \mu\text{m}$ refers to the particle size range detected by the FSSP ($0.45 < d < 17.7 \mu\text{m}$) and the 2D-C ($d > 100 \mu\text{m}$).

Also shown is the particle size distribution of the surrounding cirrus field. The cirrus cloud has been detected after leaving the contrail from 11:45:06 UT for 5 min as indicated by background NO and SO_2 mixing ratios and elevated extinction and particle number density. The cirrus cloud has been observed at altitudes between 10.1 and 10.4 km and temperatures of 221 to 219 K. The cirrus particle size distribution (see Fig. 10) suggests that the contrail was embedded in a cirrus cloud and that large cirrus crystals ($d > 100 \mu\text{m}$) might have been entrained in the contrail.

It has been questioned whether high ice crystal concentrations often observed with the FSSP are real or were caused by shattering of large ice crystals on protruding probe inlets (McFarquhar et al., 2007; Jensen et al., 2009). Techniques have been proposed (Field et al., 2003, 2006) to correct the particle size distributions based on information of the particles interarrival times (Fast-FSSP or 2D-C) and new particle image probes with high pixel resolution (including CIP and 2D-S) may be used to quantify shattering artefacts. In our cirrus case, the particle concentration may be affected by particle shattering as large ice crystals were detected by the 2D-C and CPI instruments. In contrast, the contrail FSSP-300 measurements are not strongly affected by ice-crystal shattering since the cirrus contribution to the contrail ice crystal surface or volume distribution for particles smaller than $17.7 \mu\text{m}$ is less than 1% (see Fig. 10), ruling out that the contrail particle size distribution is significantly influenced by shattering. During contrail encounters, the small particle mode with $d < 17.7 \mu\text{m}$ is clearly dominated by contrail ice crystals.

The contrail particle size distribution for particles with $0.45 < d < 17.7 \mu\text{m}$ has a mean total ice crystal surface A of $0.51 \mu\text{m}^2 \text{ mm}^{-3}$ and a mean contrail ice volume V of

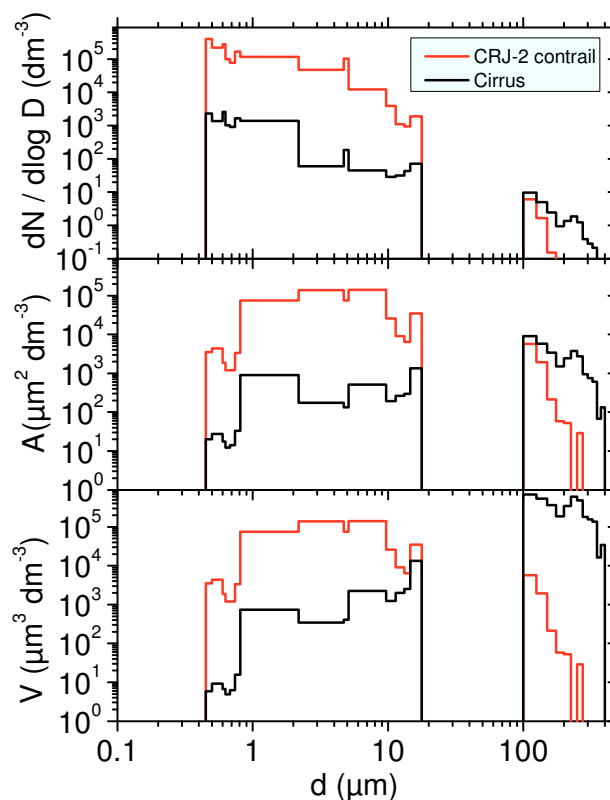


Fig. 10. Particle size, surface and volume distributions of the CRJ-2 contrail (red) embedded in a cirrus cloud (black). The ice surface area (ice volume) is given in $\mu\text{m}^2 \text{ dm}^{-3}$ air ($\mu\text{m}^3 \text{ dm}^{-3}$ air) for each size bin. The total ice surface area and total ice volume can be calculated by summing up the individual size bins.

$1.80 \mu\text{m}^3 \text{ mm}^{-3}$. Using the equation $r_{\text{eff}} = (3/4) V/A$ yields an effective radius r_{eff} of $2.6 \mu\text{m}$. Few observation of young contrails with ages of less than 3 minutes exist. The contrail particle size distribution detected during CONCERT broadly agrees with observations in a less than 3 minutes old contrail by Schröder et al. (2000). Similar concentrations of ice crystals ($50\text{--}300 \text{ cm}^{-3}$) with sizes below $10 \mu\text{m}$ were detected by Baumgardner and Gandrud (1998) in a 30 seconds old contrail at 210 K. Poellot et al. (1999) observed slightly larger particles with mean diameters of $8 \mu\text{m}$ at mean concentrations of 10 cm^{-3} in contrails older than 3 minutes.

7.1 Processing of NO and SO_2 in the engine and the plume/contrail from the CRJ-2 aircraft

Amongst gaseous aircraft emissions, SO_2 and NO_x are the most important precursors for condensable acids that influence particle formation and composition in the upper troposphere. The rate of conversion of fuel sulfur into H_2SO_4 , ϵ_S , has been subject to many investigations (Fahey et al., 1995; Curtius et al., 1998; Miake-Lye et al., 1998) with recent estimates of ϵ_S ranging from 2 to 10%. Simultaneous

measurements of HONO, NO and SO₂ in aircraft exhaust plumes under cruise conditions do not exist. HONO/NO ratios have been measured in flight (Arnold et al., 1992) and HONO/NO_x ratios on the ground (Wormhoudt et al., 2007; Wood et al., 2008) with an average of 2% for thrust settings between 30 and 100%.

As an example we discuss here the encounter of the CRJ-2 contrail with a CF34-3B engine detected on 19 November 2008 shown in Fig. 6. We observed a good correlation of the NO mixing ratios detected with a chemiluminescence detector and the HONO and SO₂ mixing ratios measured with the ITCIMS instrument. Within the plume HONO mixing ratios range from 0.26 nmol/mol to values below the detection limit and SO₂ shows peak values of 0.69 nmol/mol with a background of 0.08 to 0.1 nmol/mol.

The HONO/NO ratio was calculated by integrating over all measurements in the contrail to account for the different time resolution of the two instruments. For this aircraft and plume ages of 62 to 122 s, HONO/NO (HONO/NO_y) ratios were 0.037 (0.024) with an instrumental error of $\pm 42\%$. The high HONO/NO value suggests that NO conversion to NO₂ has occurred within in the first minute after emission. Higher HONO/NO values are therefore expected in aging plumes compared to measurements at the engine exit (Wood et al., 2008).

In order to derive the conversion efficiency of fuel sulfur ϵ_S in the engine we compare the two formation pathways of HONO and HSO₃ (a precursor of H₂SO₄) in the turbine segment of the engine. A major part of the OH-induced conversion of the precursor gases SO₂ and NO takes place in the combustion chamber and the turbine segment of the engine (Kärcher et al., 1996; Somnitz et al., 2005). Here SO₂ reacts with OH via the Stockwell-Calvert-Mechanism to form H₂SO₄ and NO reacts directly with hydroxyl radicals to form HONO. Under the assumption that the ratio of the effective rate coefficients for the two reactions is near unity the sulfur conversion efficiency can be inferred.

In the case of the CRJ-2 the fuel sulfur conversion efficiency yields a value of 2.9 %. This fraction can be regarded as a lower limit because our method accounts specifically for the OH-induced conversion of SO₂. This lower limit is in good agreement with recent model simulations of the sulfur conversion efficiency (Starik et al., 2002), and compares well with former direct in-situ measurements of total sulfuric acid (Curtius et al., 2002) in aircraft exhaust plumes. The uncertainty of the method arises from the uncertainty of the effective rate coefficients for the reaction SO₂ + OH and NO + OH. The advantage of this method to derive the sulfur conversion efficiency is that it is independent of dilution and the initial OH concentration and it benefits from reliable trace gas measurements.

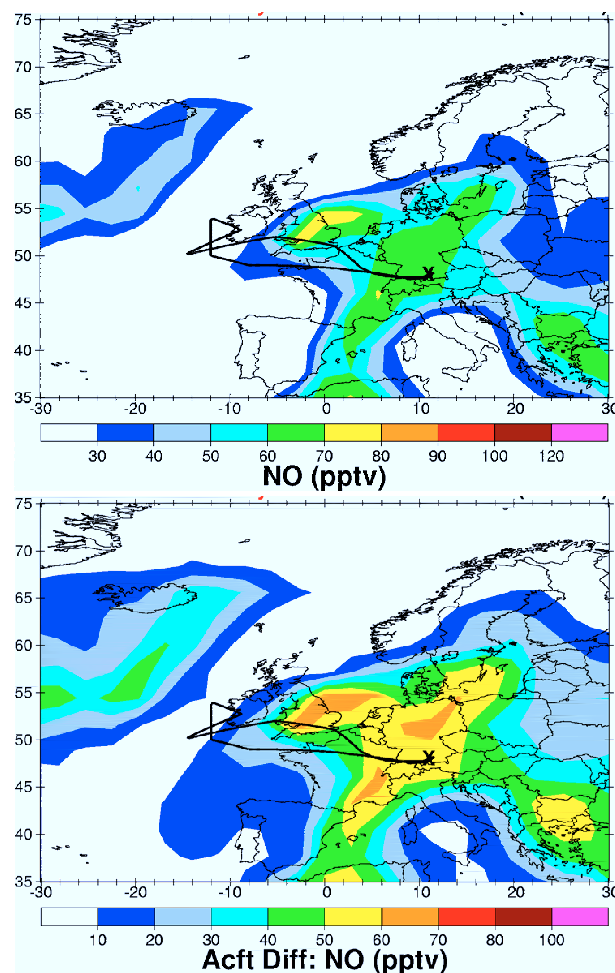


Fig. 11. Simulations with the MATCH-MPIC model of background NO (upper panel) and aircraft NO emissions (lower panel) on 31 October 2008, 12 UT, at 250 hPa above Europe. Also shown is the flight trajectory of the Falcon (a return flight from Oberpfaffenhofen to Shannon, black line).

8 Aircraft NO corridor effects in the lowermost stratosphere

Besides small scale plume processing also the detection of regional scale aircraft emissions were within the scope of the CONCERT campaign.

Therefore, the MATCH-MPIC model (Lawrence et al., 2003) was used for the prediction of regional aircraft NO emissions. The MATCH-MPIC model is a global, three dimensional chemical transport model driven by meteorological analyses and forecasts from NCEP (National Centers for Environmental Prediction). During the CONCERT campaign, MATCH-MPIC was run semi-operationally in two configurations, one with, and the other without aircraft NO emissions. Aircraft emissions used for these simulations are described by Schmitt and Brunner (1997). Examination of

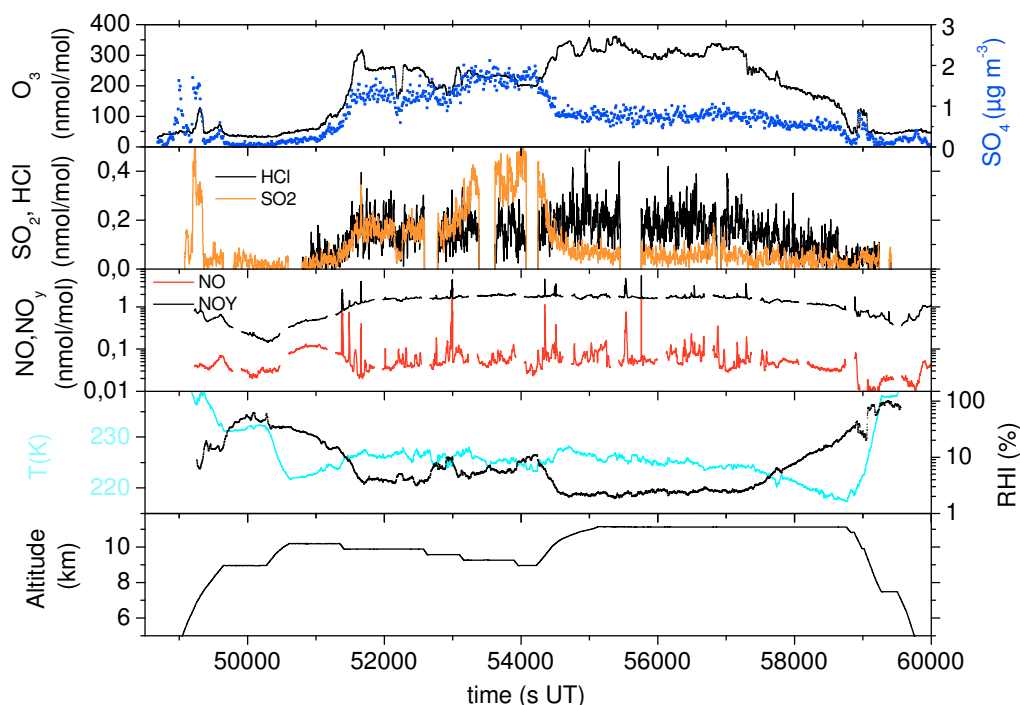


Fig. 12. Altitude, temperature, RH_I, NO, NO_y, O₃, SO₂, HCl, and sulfate concentrations detected on 31 October 2008 on flight B from Shannon to Oberpfaffenhofen. Large scale aircraft emissions in the lowest stratosphere and the aged eruption plume from volcano Mt. Kasatochi have been observed.

the differences between the output from these two model configurations allowed the prediction of large-scale regions influenced by aircraft emissions. Fig. 11 shows examples of the background and the aircraft NO predictions for 31 October 2008, 12 UTC, at 250 hPa including the flight path of the Falcon. The simulations show an extended feature above Great Britain with elevated NO concentrations of up to 80 pmol/mol. In this region aircraft emissions additionally contribute up to 70 pmol/mol to atmospheric background concentrations, suggesting that more than 40% of the total NO might originate from aircraft emissions. This day was subsequently chosen for a flight to detect aircraft NO corridor effects.

The region with enhanced NO was crossed twice by the Falcon on 31 October during a return flight from Oberpfaffenhofen to Shannon and back in the lowest stratosphere at altitudes between 10.5 and 11.6 km (Fig. 12). The detection of spikes in the NO and NO_y data within that region shows the presence of aircraft exhaust plumes with ages of minutes up to few hours (Schumann et al., 1998). The detection of aircraft plumes and their superposition is a first indication for the observation of aircraft corridor NO_x effects. In addition to the individual aircraft plumes, large scale elevated NO mixing ratios of 50 to 100 pmol/mol were detected in the lower stratosphere within the region influenced by the aircraft emissions compared to unperturbed stratospheric NO levels <50 pmol/mol. The measurements qualitatively agree

with simulations with the MATCH-MPIC model of background and aircraft NO mixing ratios above Europe (Fig. 11). Hence we suggest that elevated NO levels in that region might be attributed to the superposition and dispersion of aged aircraft plumes and to regional aircraft NO effects. Our first rough analysis invites for a detailed study of the impact of aircraft NO in atmospheric composition. While several studies of the aircraft NO_x impact in the upper troposphere exist (Schlager et al., 1997, 1999; Ziereis et al., 1999; Schumann et al., 2000), our measurements present a rare observation (Kondo et al., 1999; Koike et al., 2000) of a substantial large scale aircraft NO effect in the lowermost mid latitude stratosphere.

9 Detection of aged stratospheric aerosol layers from volcano Mt. Kasatochi

A natural source of particles and trace gases in the UTLS region are major volcanic eruptions. Volcanic eruptions inject sulfur into the stratosphere, preferably as gaseous SO₂, which undergoes stratospheric conversion to sulfuric acid and ultimately to sulfate aerosols. Increased aerosol surface areas after major volcanic eruptions increase the planetary albedo, resulting in transient global surface cooling lasting for about 1 to 2 years. In addition, volcanic eruptions may also inject chlorine species into the stratosphere, as gaseous

HCl or as NaCl particles. Depending on the RH_I within the eruption plume, a significant fraction of the emitted gaseous HCl can reach the stratosphere (Textor et al., 2003). Halogen activation on volcanic aerosol can contribute to ozone loss in the lowest stratosphere (Borrmann et al., 1997; Solomon et al., 1998).

An aged volcanic eruption plume was observed during 3 CONCERT flights in the lowest stratosphere. The plume originated from the 7/8 August 2008 eruption of the Aleutian volcano Kasatochi (52° N), which injected about 1.5 Mt of SO₂ into the lower stratosphere, representing a rare case of a major volcanic eruption in northern high latitudes. We measured the eruption plume three month later on 28 October and on two flights on 31 October (Fig. 12) in a stratospheric tongue extending over Western Europe. As the Falcon cruised at 8.5 to 11 km altitude, two stratified plumes were sighted as dark bands against the horizon. The SO₂ and the sulfate aerosol measurements revealed the presence of two major sulfate aerosol layers extending up to 5 km above the local tropopause. The observed sulfate concentration of 2.0 µgm⁻³ was significantly enhanced compared to background concentrations of 0.5 µgm⁻³ (Schmale et al., 2010). The increase of sulfate is most likely due to stratospheric OH-induced conversion of injected volcanic SO₂ leading to gaseous sulfuric acid, followed by condensation. The maximum SO₂ mixing ratio of 510 pmol/mol in the plume very substantially exceeded the atmospheric background SO₂ of about 40 pmol/mol (Jurkat et al., 2010). The ratio of particulate sulfate to total sulfur inside the plume of 0.8±0.1 reflects slow OH-induced SO₂ conversion to sulfuric acid with an upper limit of the e-folding SO₂ lifetime of 60 days (Jurkat et al., 2010).

Additional trace gas measurements of HCl, HNO₃ and O₃ allow to investigate heterogeneous reactions taking place on the sulfate particles. The ratio HCl/O₃ tends to be increased by 19% in the SO₂-rich plumes compared to outside plume conditions suggesting direct injection of volcanic HCl into the stratosphere and very slow processing of HCl on the aerosol in the polar summer stratosphere. HNO₃/O₃ was enhanced by 50% in the plumes suggesting increased N₂O₅ hydrolysis on enhanced sulfate aerosols (Jurkat et al., 2010).

21% of the volcanic aerosol consisted of carbonaceous material, which is factor of 1.4 more than observed in UTLS background aerosol (Schmale et al., 2010). The increase of organic species may be due to direct injection of organic aerosol into the stratosphere. Alternatively, organic aerosol may form in the lower stratosphere from injected organic precursor gases, which became entrained into the tropospheric segment of the eruption column. The increased organic aerosol fraction might influence uptake coefficients and heterogeneous processing on volcanic sulfate aerosol.

10 Conclusions

Chemical, microphysical and radiative properties of contrails, cirrus clouds and an aged volcanic aerosol layer were detected during the CONCERT campaign in October/November 2008 over Western Europe and were compared to model simulations. The observations of microphysical and chemical contrail properties provide a solid base for future studies.

Of particular interest is the first detection of the contrail from an A380-800 aircraft and the temporal evolution of particle size distribution and particle shape in the contrail. Further the mixing of contrail ice crystals with ambient cirrus cloud particles and the evaporation of contrail ice crystals under ice sub-saturated conditions can be addressed with model simulations. Particle properties such as the ratio between the volume mean and the effective radius of contrail ice crystals can be discussed. The chemical processing of trace gases in contrails and the large scale perturbation of the UTLS composition by aircraft emissions can be studied in detail. The contrail observations during the CONCERT campaign further allow to calculate the radiative forcing from contrails from a typical present-day aircraft fleet.

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