

# THE POTENTIAL OF TIME-LAPSE GPR FULL-WAVEFORM INVERSION AS HIGH RESOLUTION IMAGING TECHNIQUE FOR SALT AND ETHANOL TRANSPORT

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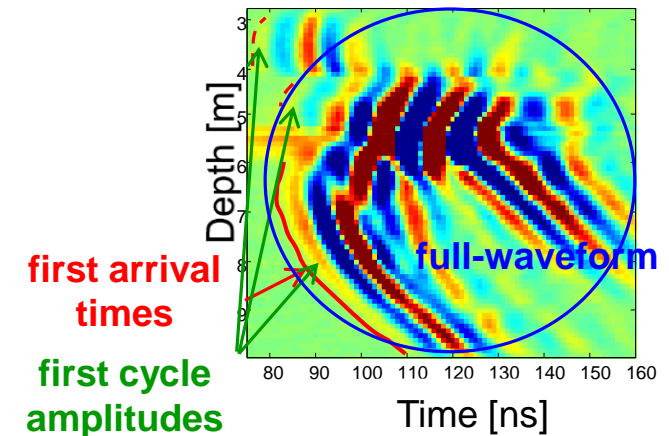
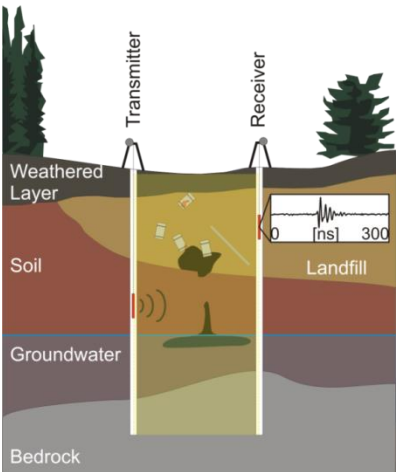
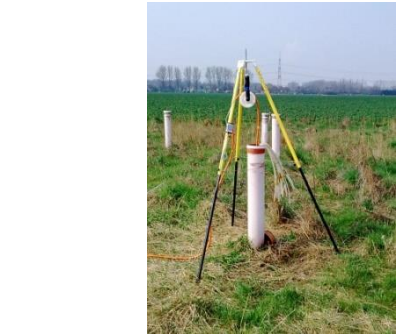
# CROSSHOLE GPR FULL-WAVEFORM INVERSION

## Mapping soil water content variability in the critical zone

GPR is able to minimal-invasively investigate distributions of electromagnetic properties:

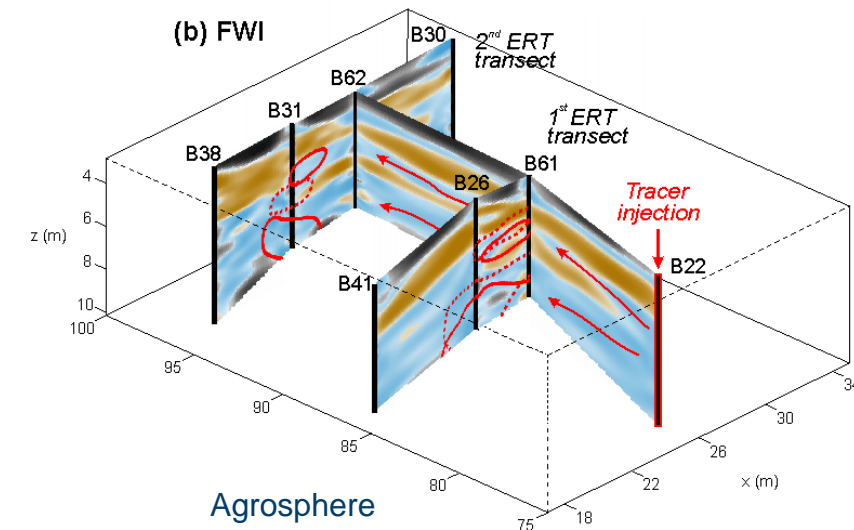
- dielectric permittivity  $\epsilon_r$  / EM velocity:  
➔ soil water content (SWC), porosity
- electrical conductivity  $\sigma$  / EM attenuation:  
➔ clay/silt content, salinity

**Improved characterization compared to other non-invasive methods!**



- facies 1 (sand)
- facies 2 (sandy gravel)
- facies 3 (gravel)

Gueting et al. (2017)



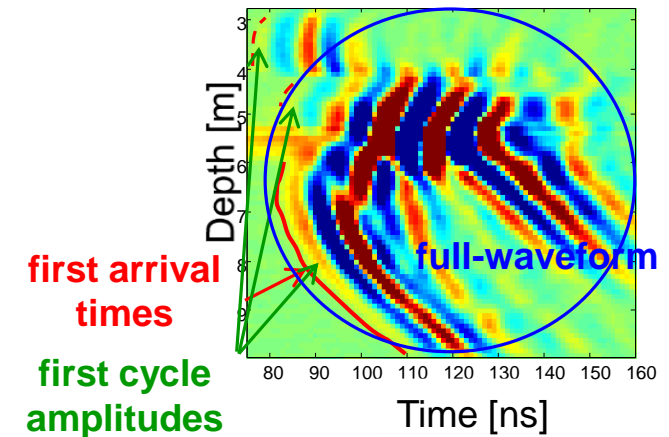
e.g., Ernst et al. (2007a, b); Virieux & Operto (2009); Meles et al. (2010); Klotzsche et al. (2010).

# CROSSHOLE GPR FULL-WAVEFORM INVERSION

## Mapping soil water content variability in the critical zone

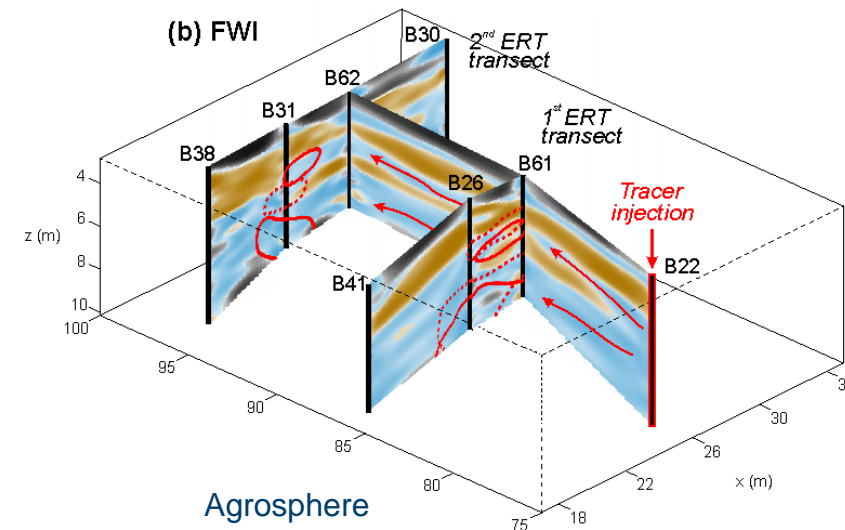
To protect groundwater from pollution and to manage an aquifer, detailed knowledge of the flow and transport processes are necessary:

- No time-lapse data investigated using GPR FWI to obtain in-depth knowledge of tracer distribution within the aquifer.
- Investigate the potential of FWI to map monitor different tracer plumes (numerical study).
  - Positive electrical conductivity: saline, heat tracers
  - Negative electrical conductivity: deionized water.
  - Negative contrast in bulk permittivity and conductivity: ethanol tracer.
- Recently, saline and heat field tracer test was performed at the Krauthausen test site (Germany).



- facies 1 (sand)
- facies 2 (sandy gravel)
- facies 3 (gravel)

Güting et al. (2017)



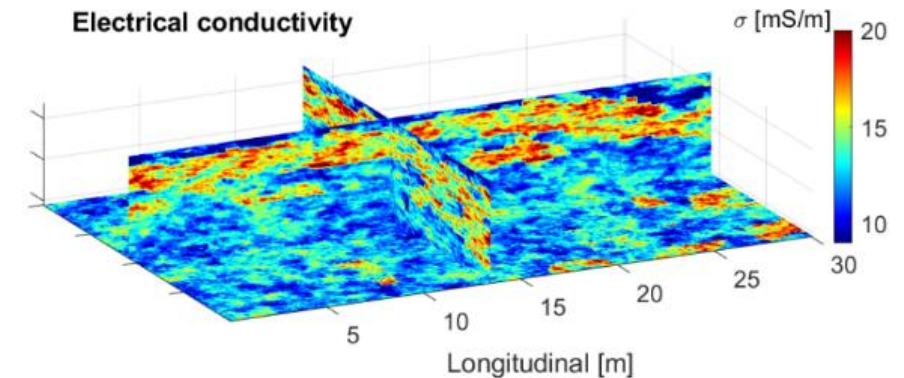
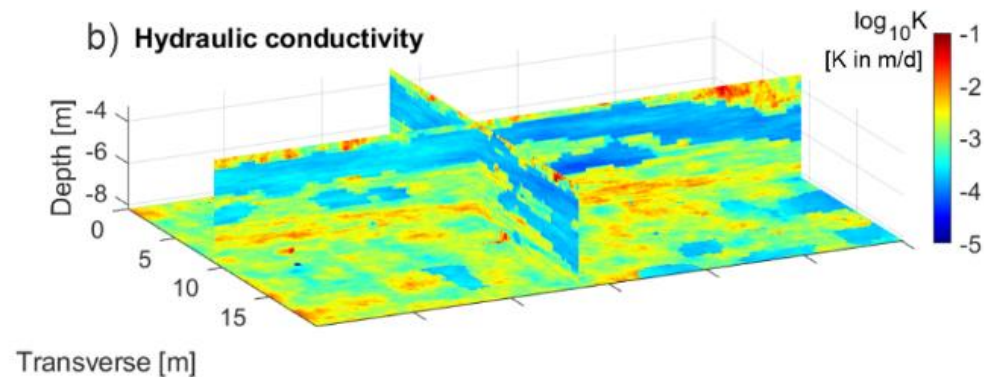
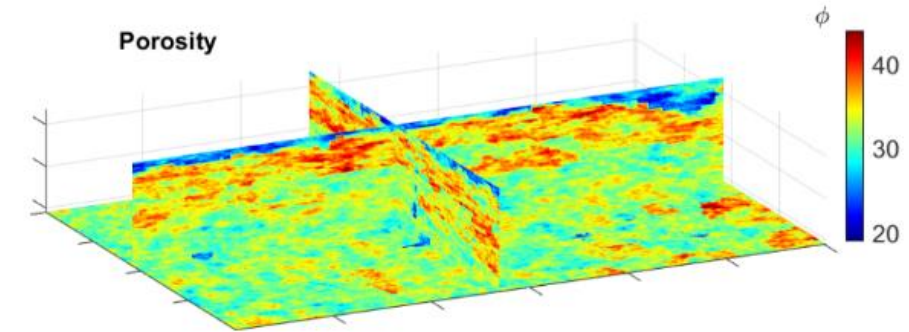
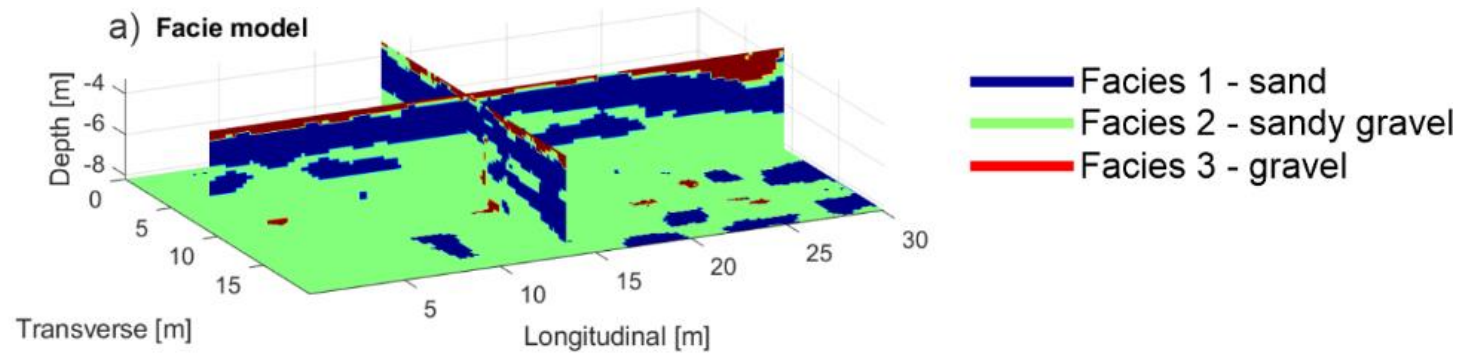
# KRAUTHAUSEN AQUIFER MODEL

High resolution hydrological model:

- 3D facies model was adopted, (Gueting et al., 2017, 2018)
- Hydraulic conductivity  $K$  (Velocity analysis (Englert, 2003), Tracer test (Mueller et al., 2010)).
- Porosity  $\phi$  (Cone penetration tests (Tillman et al., 2008), GPR FWI (Gueting et al., 2017)).
- Dielectric permittivity  $\epsilon_r$  from porosity using CRIM model (Birchak et al., 1974)).
- Electrical conductivity  $\sigma_b$  ( $r=0.5$  correlation between  $\sigma$  and  $\epsilon_r$  of GPR FWI (Gueting et al., 2017)).

Variogram-based Sequential Gaussian Simulation (SGeMS software; Remy, et al., 2009)

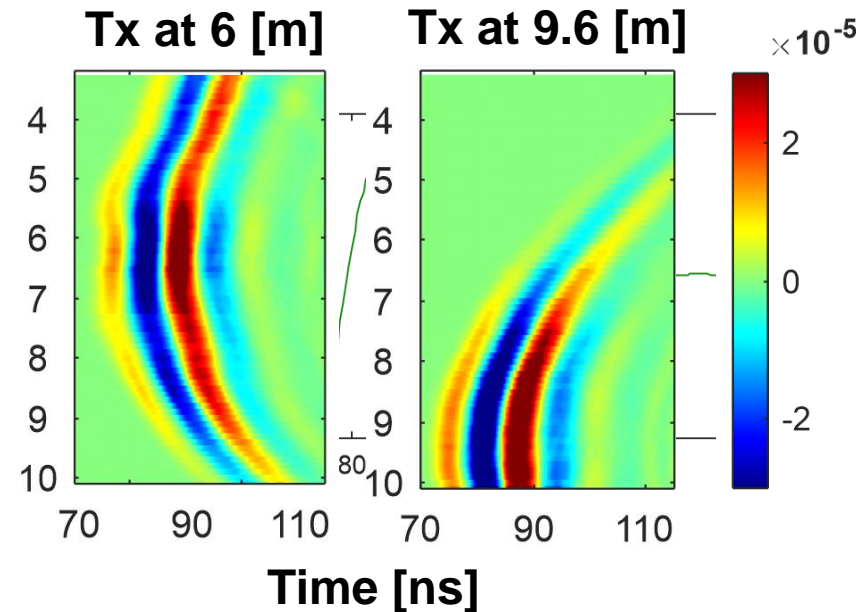
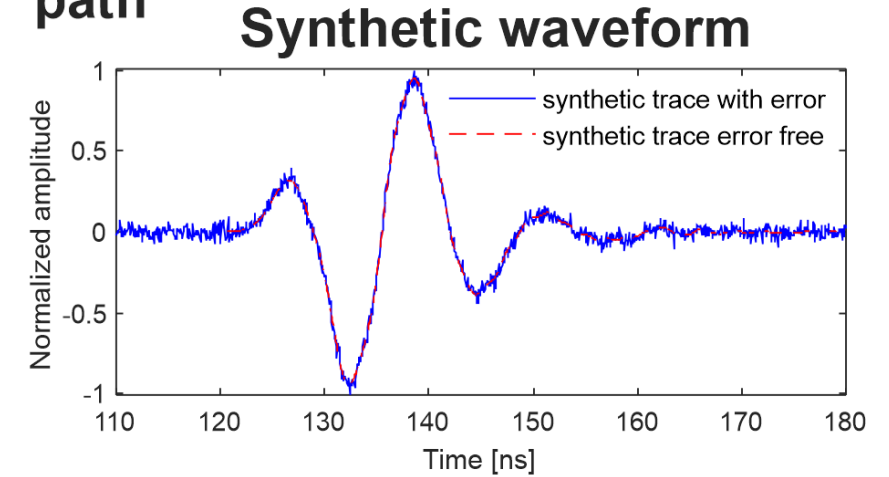
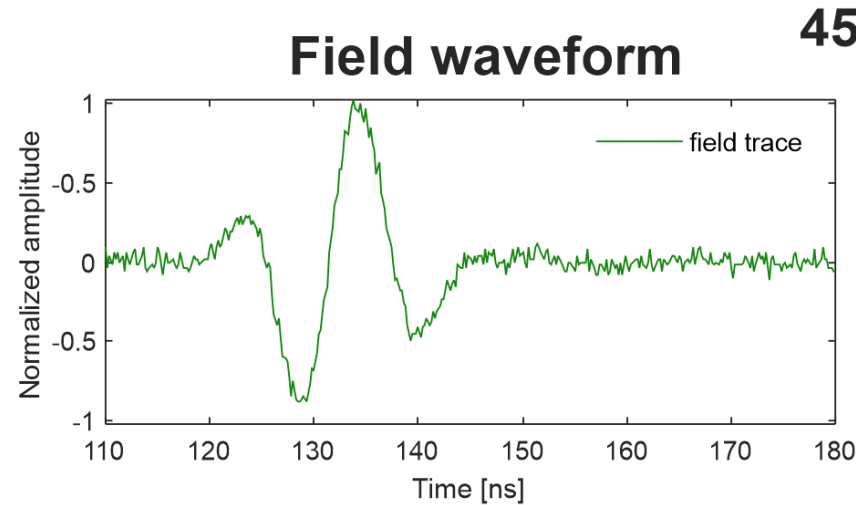
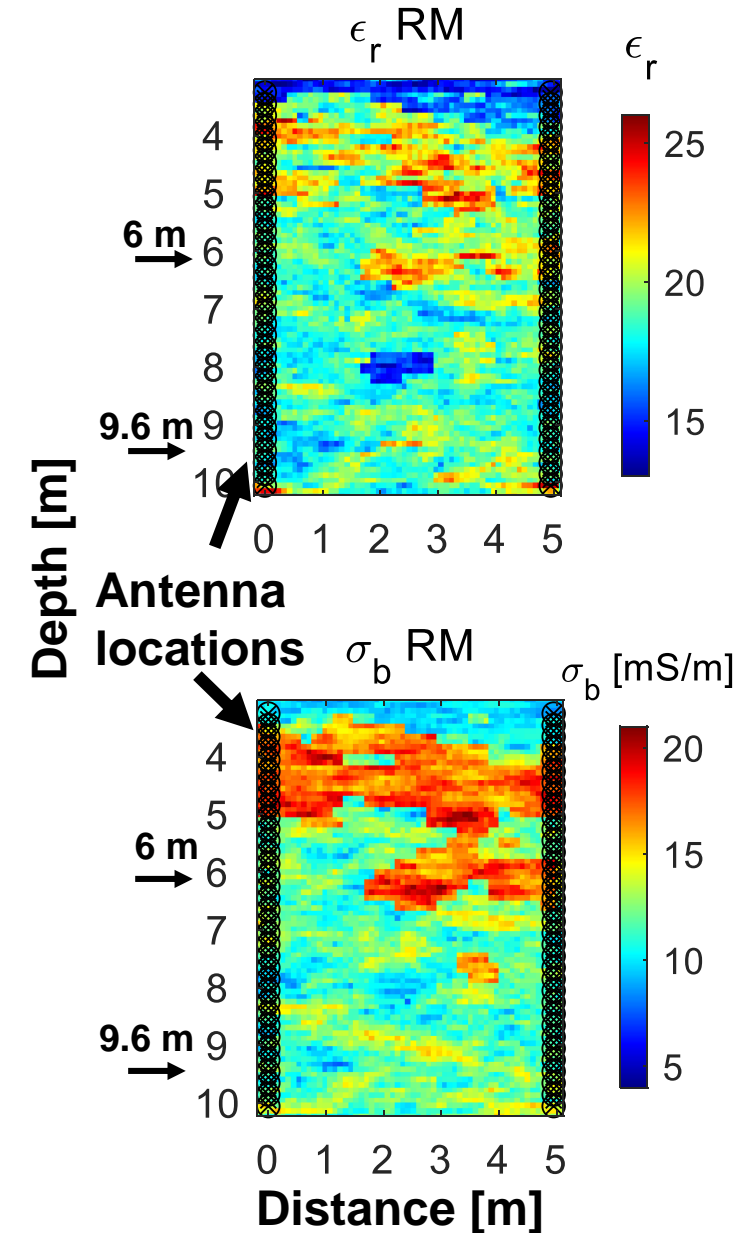
Cell size:  
0.09m



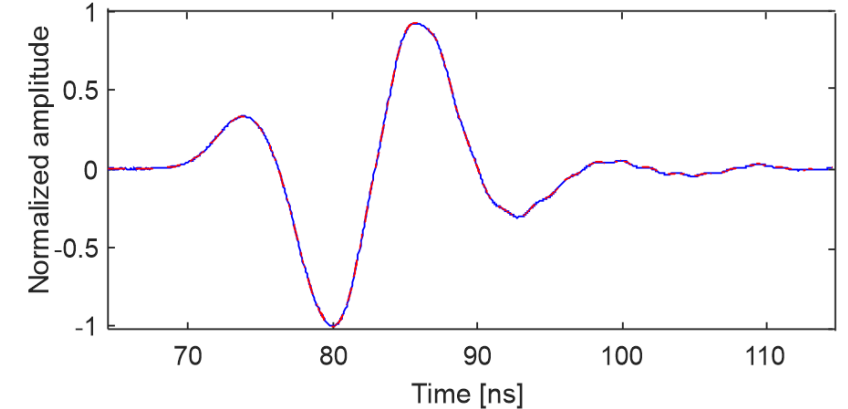


# GPR DATA

- GPR data was generated with 2D FDTD.
- Random noise was added to synthetic waveforms, with a level based on waveforms measured at the field.

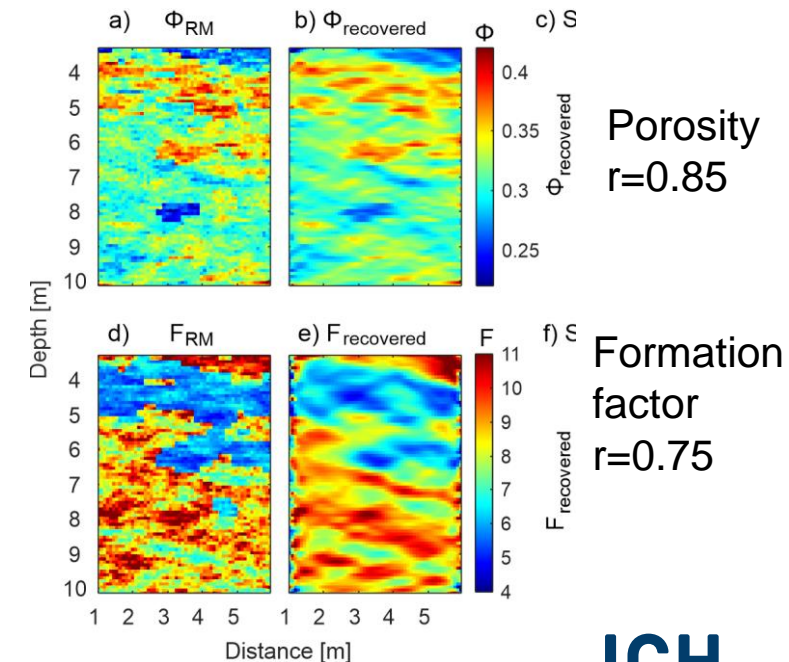
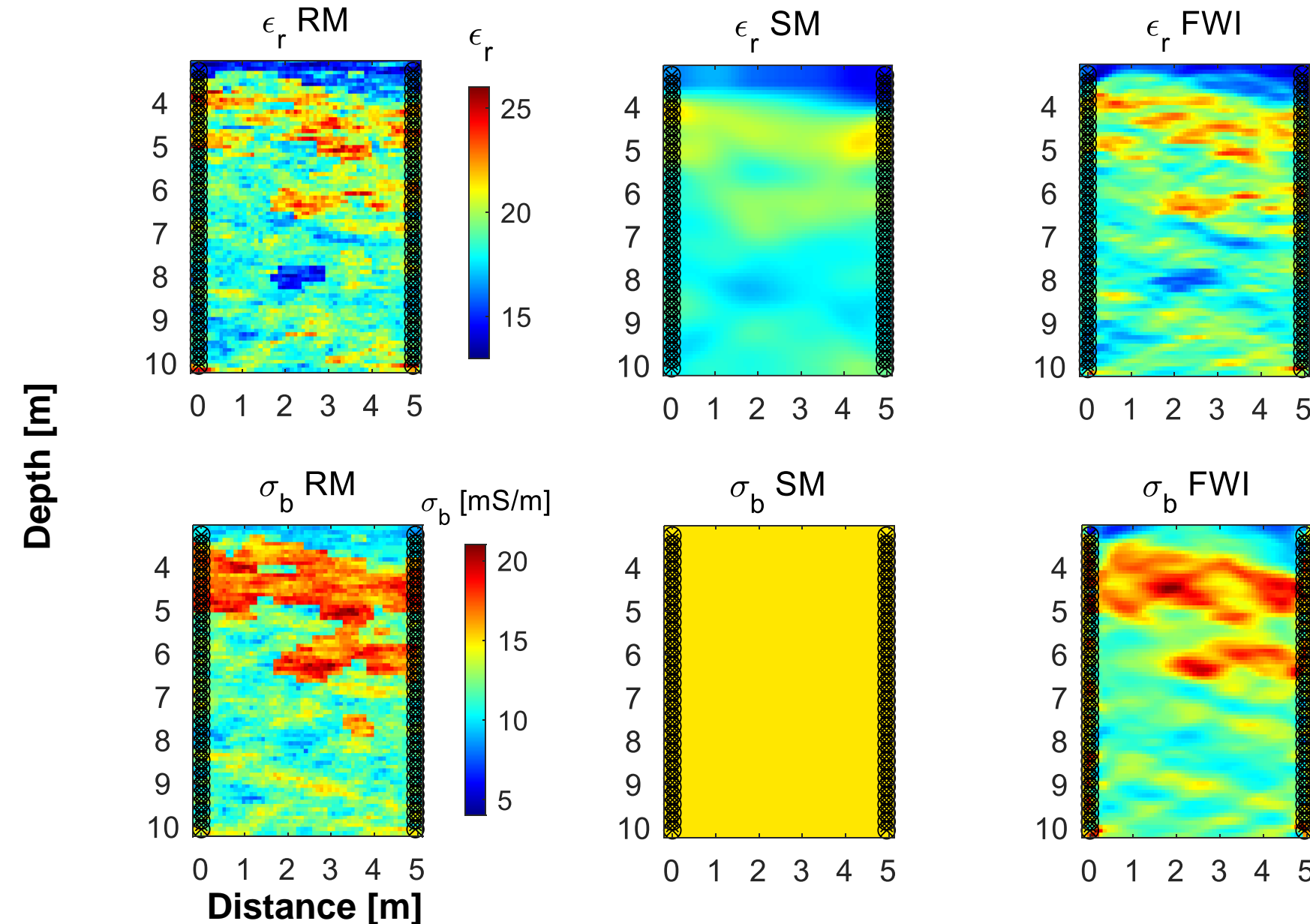


allel path



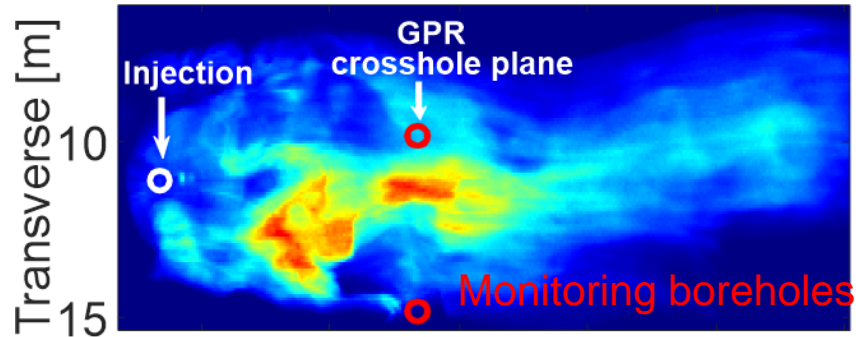
# POROSITY AND FORMATION FACTOR

- The FWI recovery of permittivity is more accurate than electrical conductivity.
- Background FWI models will be used in petrophysical relations:
  - $\epsilon_r$  to calculate porosity.
  - $\sigma_b$  to calculate the formation factor.



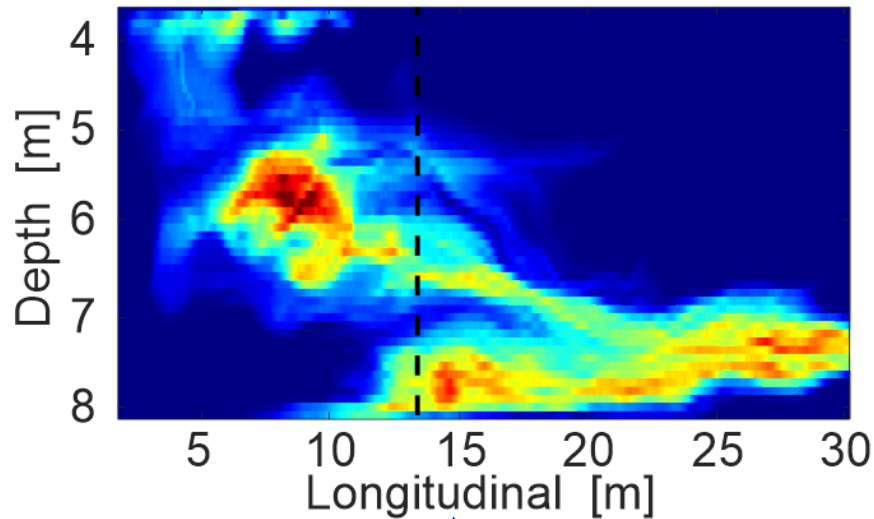
# FLOW AND SOLUTE TRANSPORT MODEL

Plan view

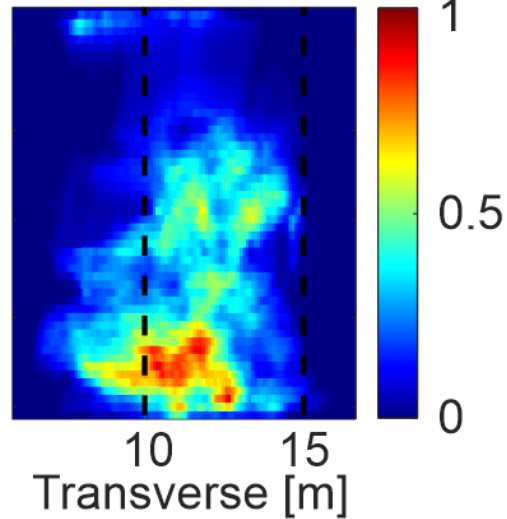


Flow model: TRACE code  
(Vereecken et al., 1994)  
Solute transport model: PARTRACE  
code (Bechtold et al., 2012)

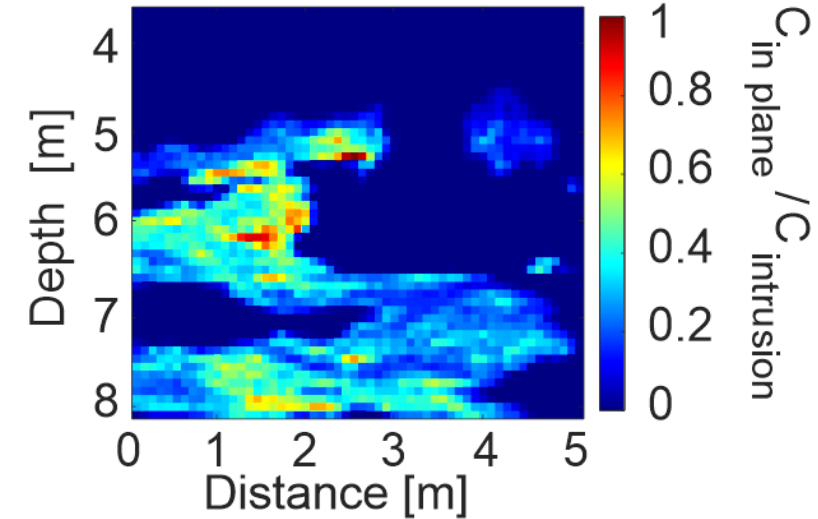
Side elevation



Front elevation  $C/C_{\max}$



Intrusion in plane (norm.)



Flow direction

At Day 15 after injection, the snapshot of tracer concentration between the boreholes is used to generate GPR data through petrophysical relations.

# PETROPHYSICAL RELATIONS FOR SALT TRACER

- No changes in  $\varepsilon_r$   
(Sreenivas & Venkatarantan, 1975).
- $\sigma_f$  increases with salinity.  
 $\sigma_b$  increases linearly with  $\sigma_{f..}$
- Concentrations from transport simulation were used to calculate  $\sigma_{f..}$

## Generation of petrophysical aquifer model from tracer changes

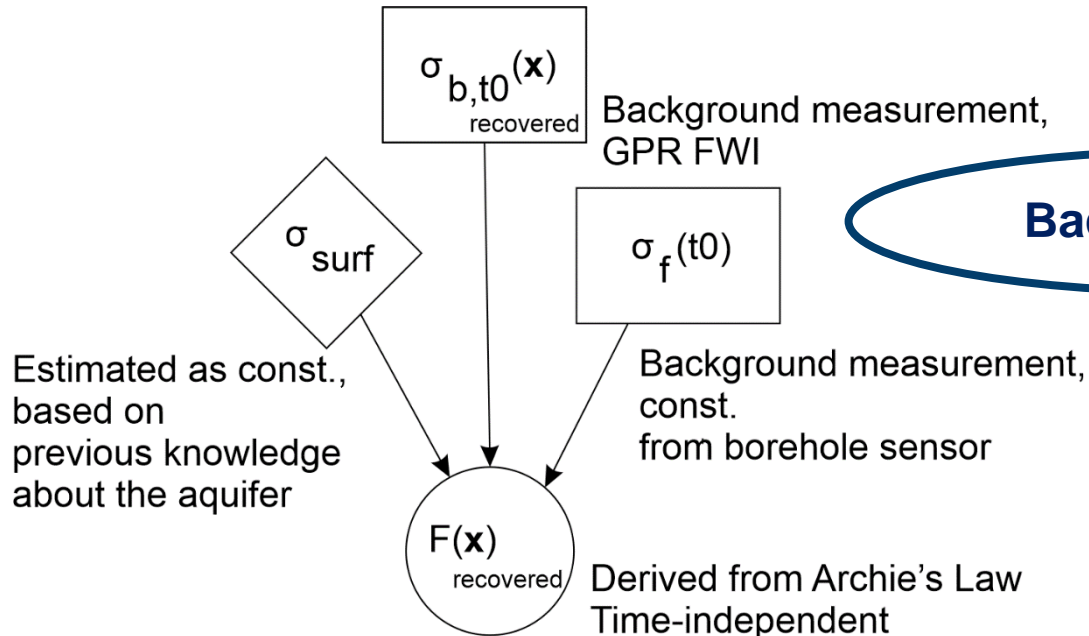
Archie's Law:  $\sigma_b(\mathbf{x},t) = \sigma_f(\mathbf{x},t)/F(\mathbf{x}) + \sigma_{surf}(\mathbf{x})$  Archie, 1942

$\sigma_{surf}(\mathbf{x}) \sim N(1.2, 0.3^2)$  mS/m with the same correlation lengths as  $\sigma_b$ .

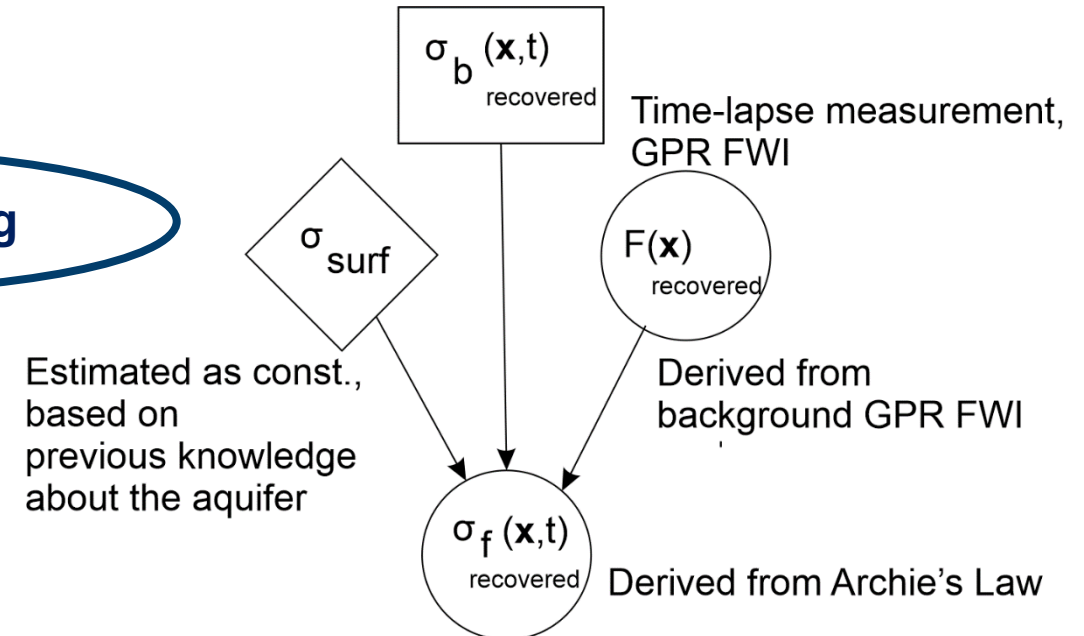
$F(\mathbf{x})_{real}$  is calculated from  $\sigma_{b^{real}}$  and  $\sigma_{f(t0)_{real}} \approx const.$

Forward processing

### 1. Derive $F(\mathbf{x})_{recovered}$ from background GPR FWI $\sigma_{b,t0}(\mathbf{x})_{recovered}$



### 2. Derive $\sigma_f(\mathbf{x},t)_{recovered}$ from time-lapse GPR FWI $\sigma_b(\mathbf{x},t)_{recovered}$



Back processing



# PETROPHYSICAL RELATIONS FOR ETHANOL TRACER

➤  $\epsilon_r$  and  $\sigma_b$  decrease with ethanol volumetric concentration  $S_{EtOH}$  in groundwater (Glaser et al., 2012).

→ Concentrations from transport simulation were used to calculate  $S_{EtOH}$ .

## Generation of petrophysical aquifer model from tracer changes

### $\epsilon_r$ model:

1. Empirical relation:

$$\epsilon_{f\_mixture}(\mathbf{x}, t) = 84.05 - 42.6 \cdot S_{EtOH}(\mathbf{x}, t) - 15.7 \cdot S_{EtOH}(\mathbf{x}, t)^2$$

2. CRIM mixing model (Birchak et al., 1974)

$$\epsilon_{r\_eff}(\mathbf{x}, t) = (\epsilon_{f\_mixture}^{0.5} \cdot \phi(\mathbf{x}) + \epsilon_{matrix}^{0.5} \cdot (1 - \phi(\mathbf{x})) )^2$$

$\phi(\mathbf{x})_{real}$  is calculated from

background GPR FWI, and  $\sigma_f(t_0) = const.$

### $\sigma_b$ model:

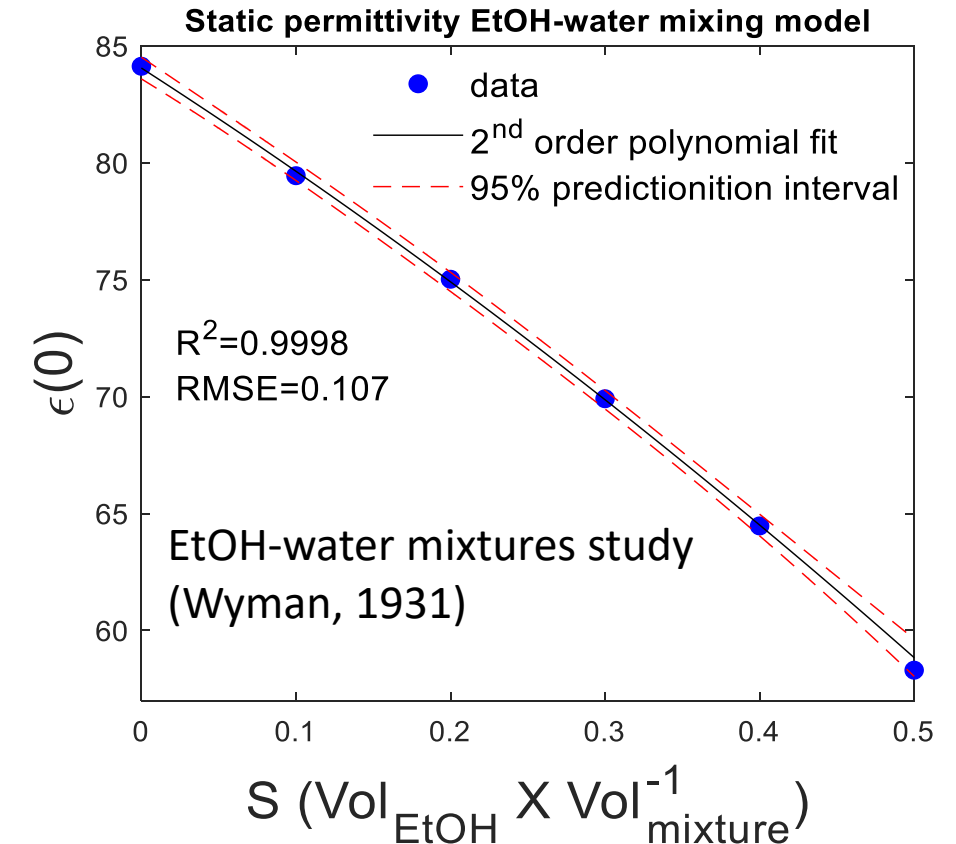
1. Mixing model:

$$\sigma_{f\_mixture} = (S_{EtOH} \cdot \sigma_{EtOH}^\alpha + (1 - S_{EtOH}) \cdot \sigma_{f,background}^\alpha)^{1/\alpha}$$

With  $\alpha = 0.3$  for  $S_{EtOH} \leq 0.4$ .

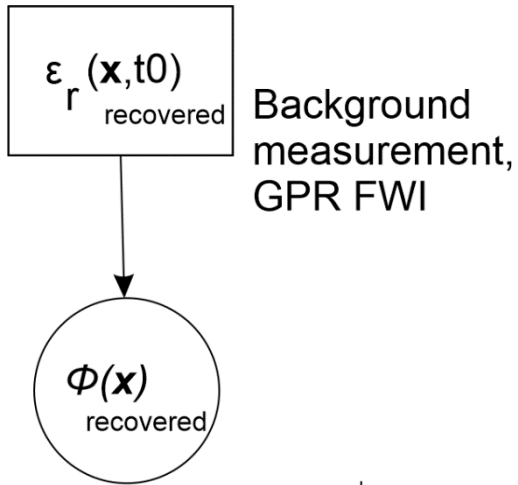
(adopted from Personna et al., 2013)

2.  $\sigma_{b,eff}$  is calculated using Archie's model:

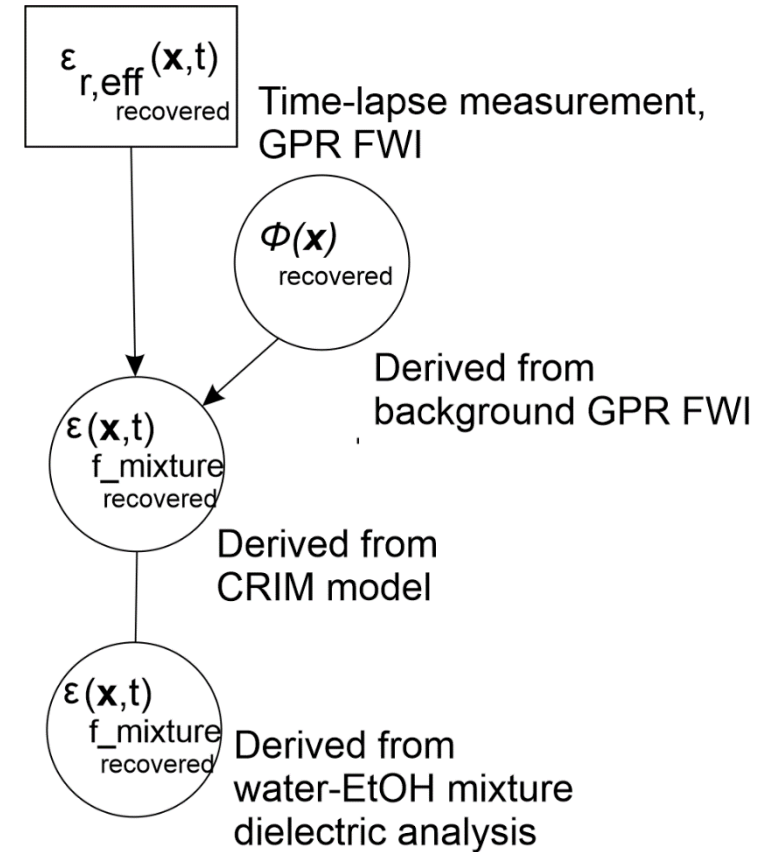


# PETROPHYSICAL RELATIONS FOR ETHANOL TRACER

1. Derive  $\phi(x)_{\text{recovered}}$  from background  
GPR FWI  $\epsilon_r(x, t_0)_{\text{recovered}}$



2. Derive  $S_{\text{EtOH}}(x, t)_{\text{recovered}}$  from time-lapse  
GPR FWI  $\epsilon_{r, \text{eff}}(x, t)_{\text{recovered}}$



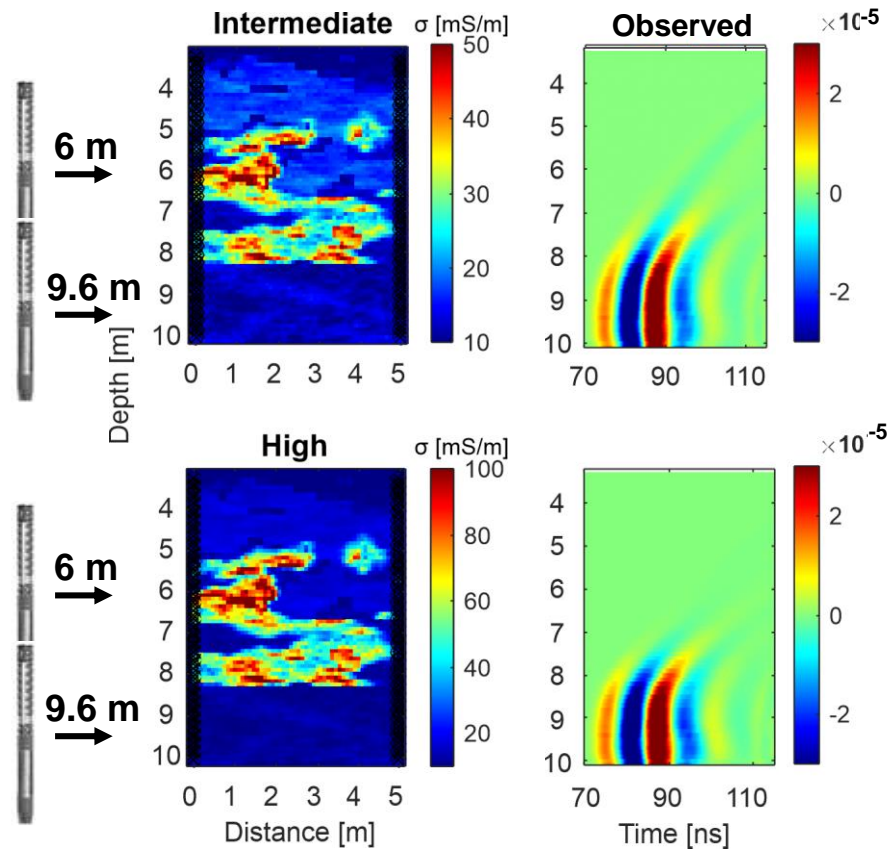
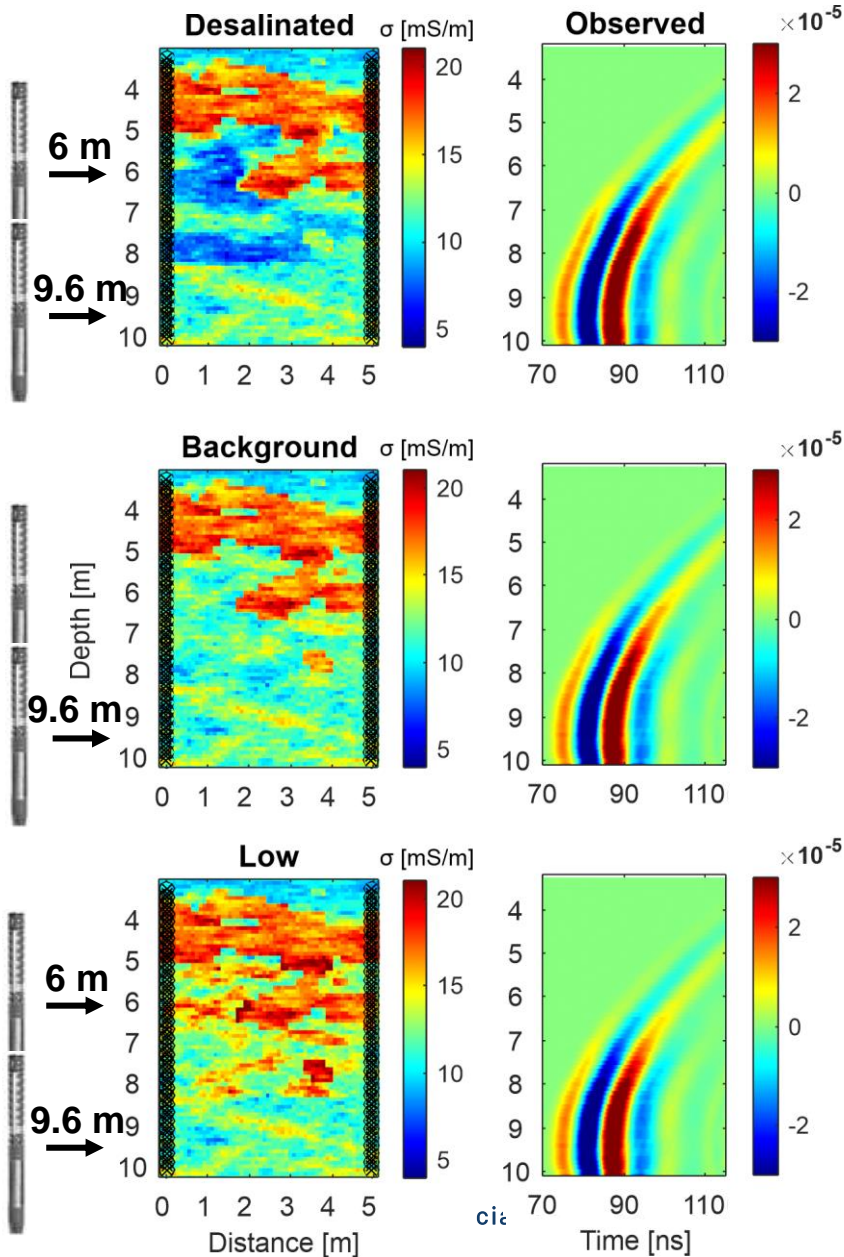
Back processing

$\sigma_b$  model:

Time-lapse GPR FWI  $\sigma_b$  enables deriving  $\sigma_{f\_mixture}$ , like for the salt tracer.

# EFFECT OF SALT TRACER ON DATA

Amplitude scale is ten-fold lower.



- Five cases (including background) of changes in conductivity levels are tested.
- Amplitude decrease with increasing fluid conductivity.

Each case  $\sigma_b$  is calculated from the transport simulation by setting different  $\sigma_{f\_injection}$

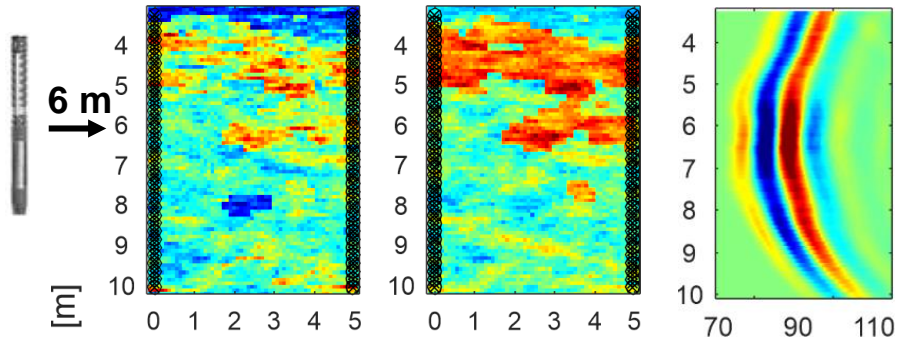
$$\sigma_f(x,t) = (n_p(x,t)) \cdot (\sigma_{f\_injection} - \sigma_{f\_background}) + \sigma_{f\_background}$$

$$\sigma_b(x,t) = \sigma_f(x,t) / F(x) + \sigma_{surf}(x)$$

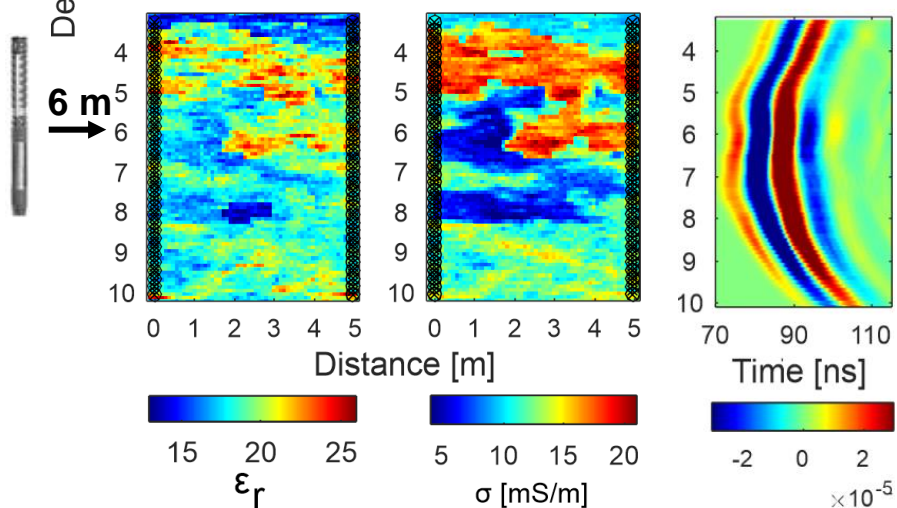
# EFFECT OF ETHANOL TRACER ON DATA

- $\epsilon_r$  and  $\sigma_b$  decrease with **ethanol volumetric concentration**  $S_{EtOH}$

Background

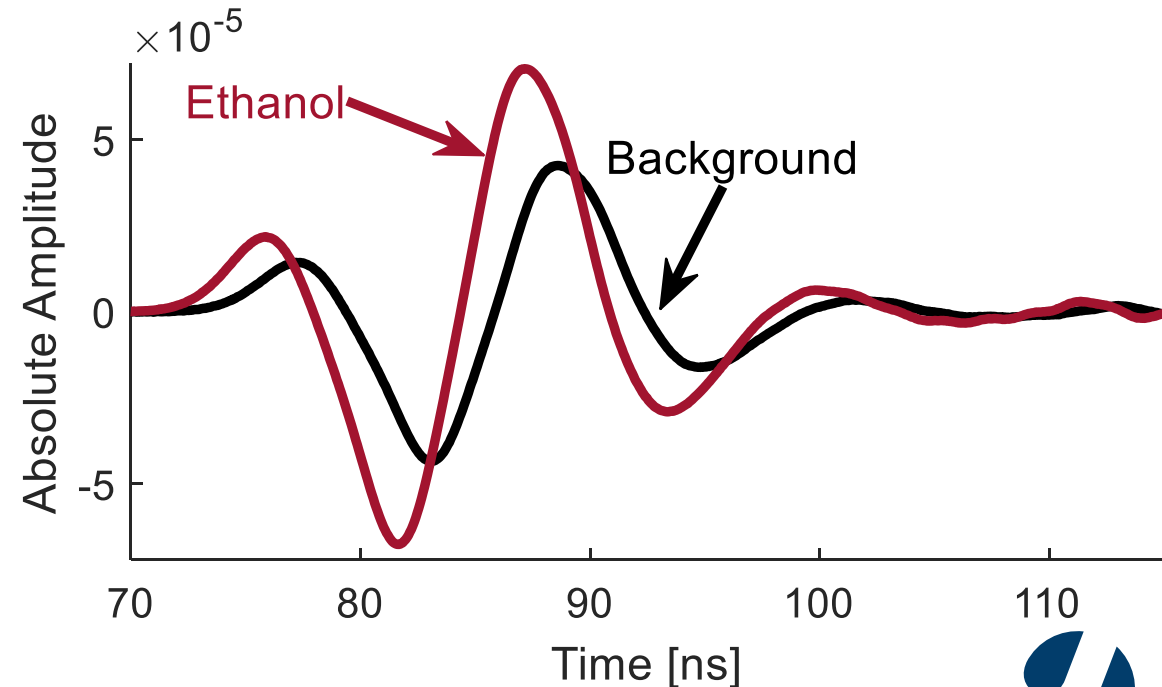


Ethanol



Maximal permittivity change of about  $\Delta\epsilon_r = -4$  between ethanol and background models lead to phase shift of  $\sim 1.5$  ns.

Transmitter and receiver at 6 m



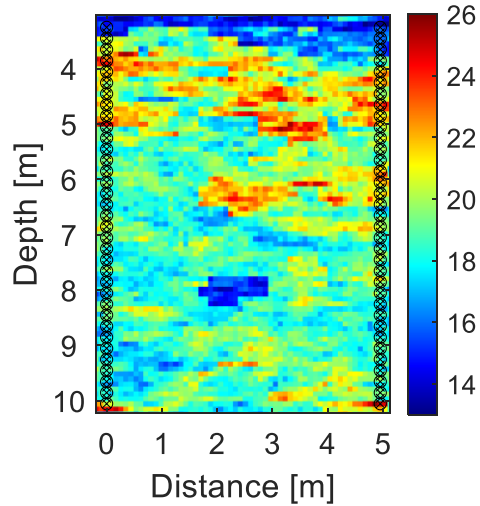


# TIME-LAPSE STARTING MODEL INVERSION STRATEGIES

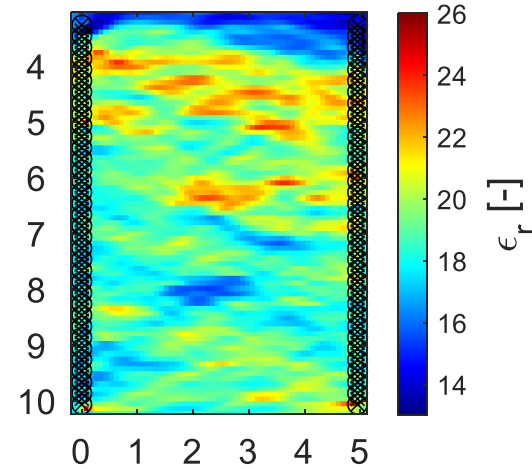
Strategy 1 – SM from background FWI recovered

Strategy 2 – SM of a previous day FWI recovered

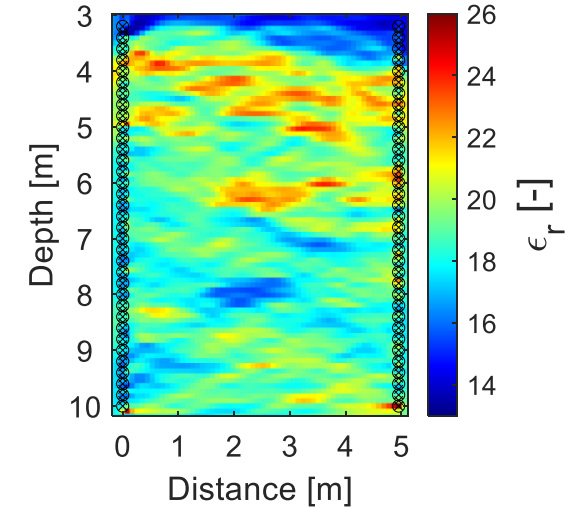
$\epsilon_r$  RM Day 15



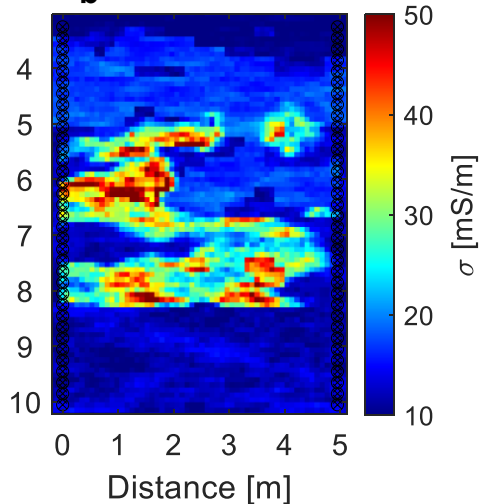
$\epsilon_r$  FWI background



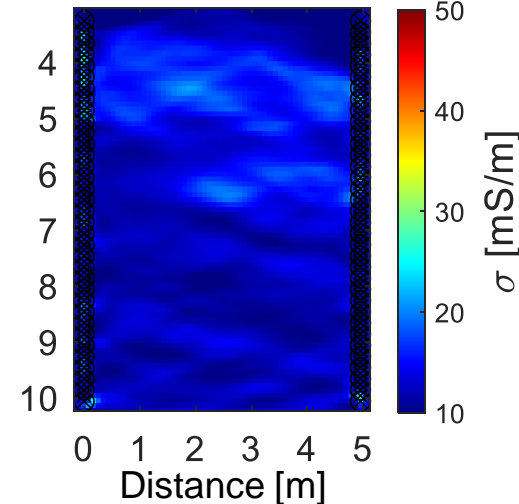
$\epsilon_r$  FWI of a previous day (13)



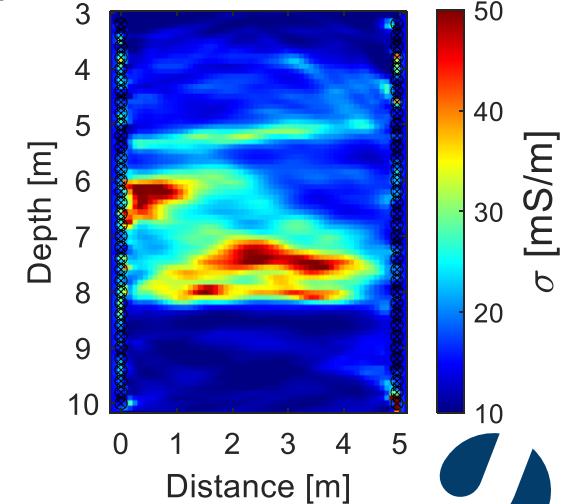
$\sigma_b$  RM Day 15



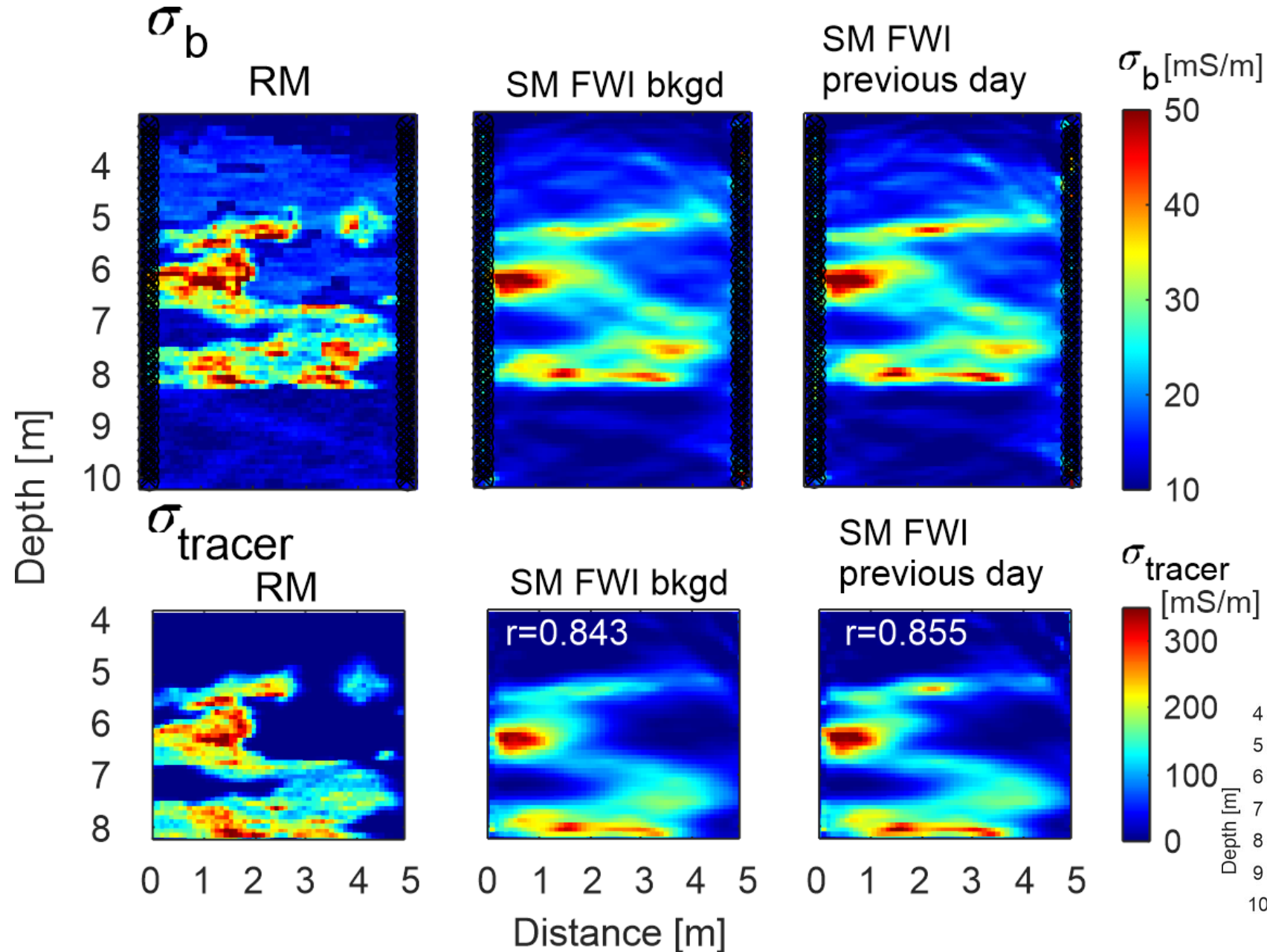
$\sigma_b$  FWI background



$\sigma_b$  FWI of a previous day (13)



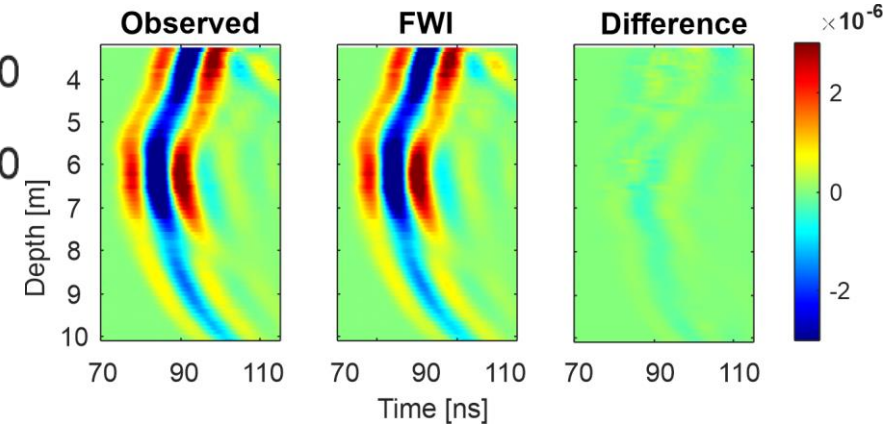
# POSITIVE “INTERMEDIATE” ELECTRICAL CONDUCTIVITY TRACER



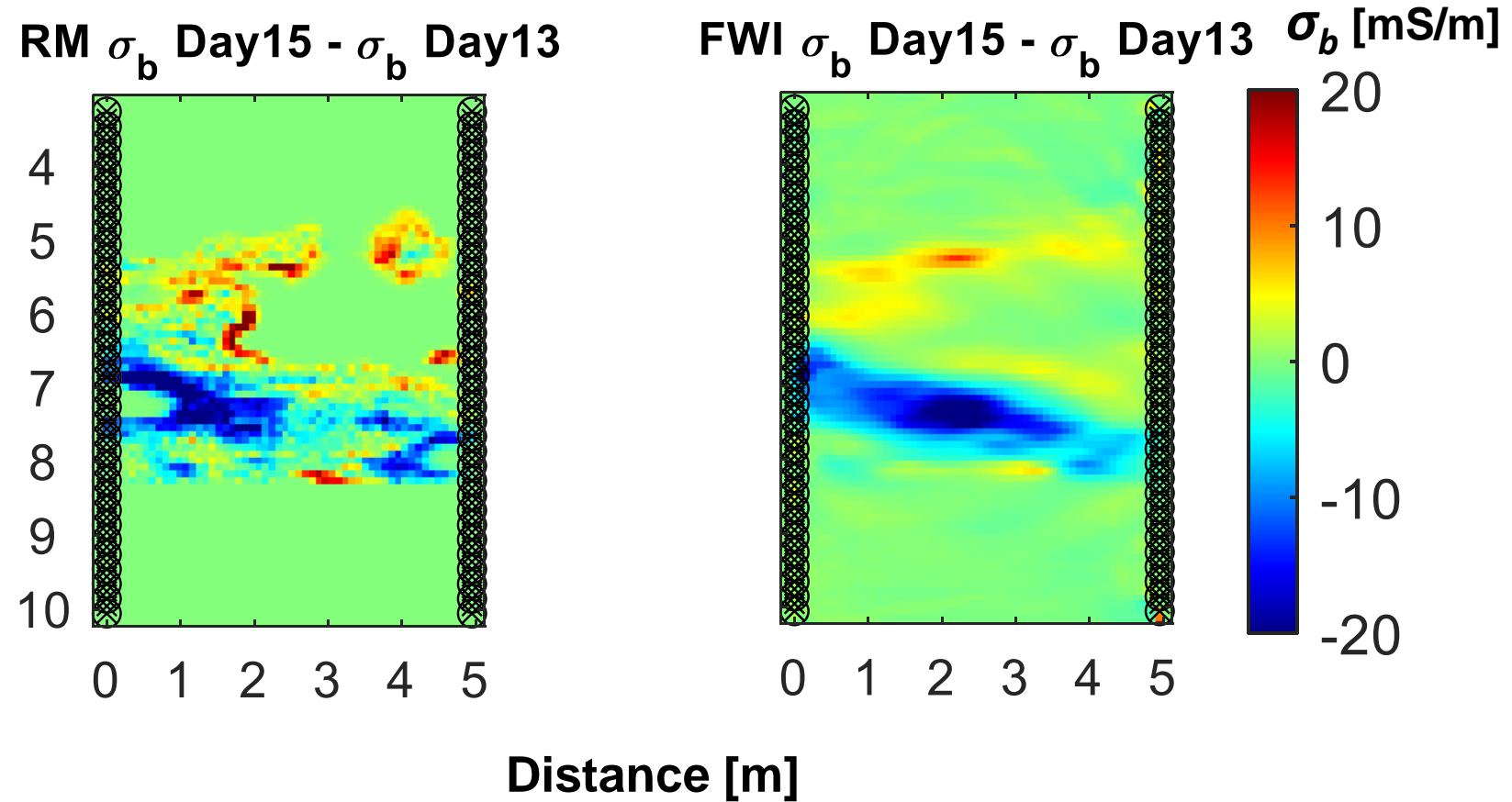
Starting model with FWI from the previous day results in better FWI model recovery (higher correlation with the RM  $\sigma_{\text{tracer}}$ ). However also SM from FWI background results in good recovery.

## Intermediate case

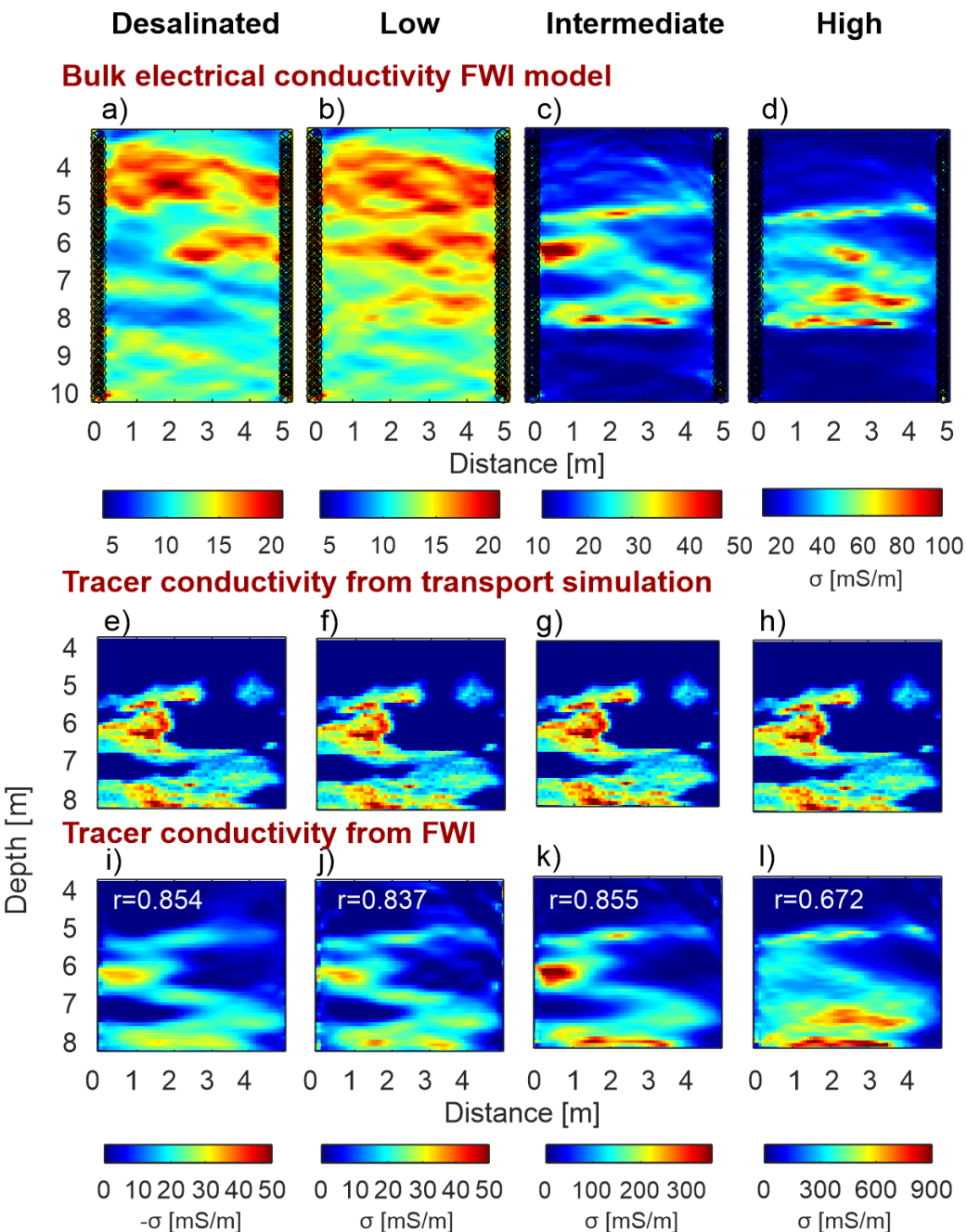
Good FWI data fit



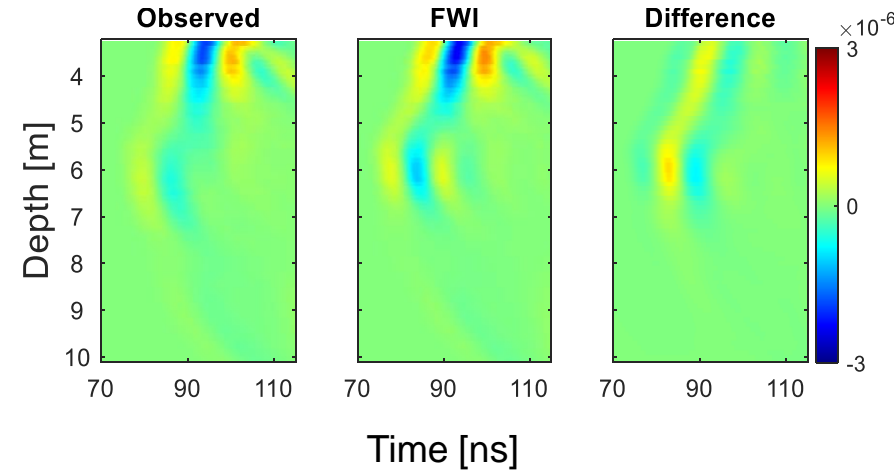
# RECOVERY OF TIME-LAPSE FWI SALT CHANGES BETWEEN DAY 15 AND DAY 13



# FLUID CONDUCTIVITY RECOVERY: COMPARISON OF SALT TRACER CASES



## High case Image scan for Tx at 6 m depth

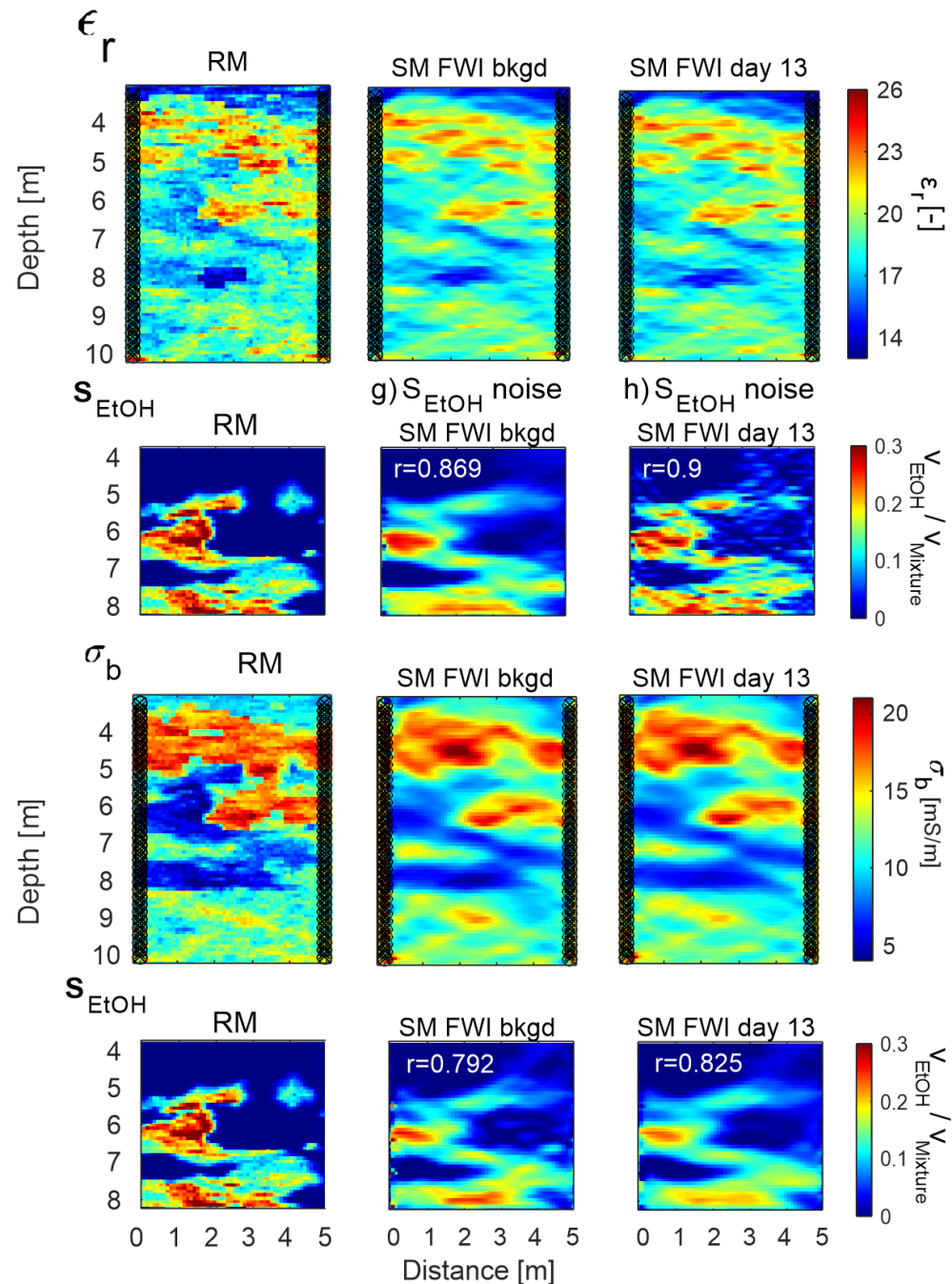


Difference in data is at the level of the observed/FWI data.

- Desalinated, Low and Intermediate recover  $\sigma_f$  well.
- Lower correlation in Desalinated, Low are not due to FWI bad fit but due to petrophysical relations, as the uncertainty in  $\sigma_{f, \text{recovered}}$  estimated from background FWI is propagated into  $\sigma_f$ , with a larger impact on low  $\sigma_b$  (see Archie's Law).
- In contrast, High case has lower correlation ( $r=0.672$ ) because of bad fit of the data, as the high conductivity depressed the signal.



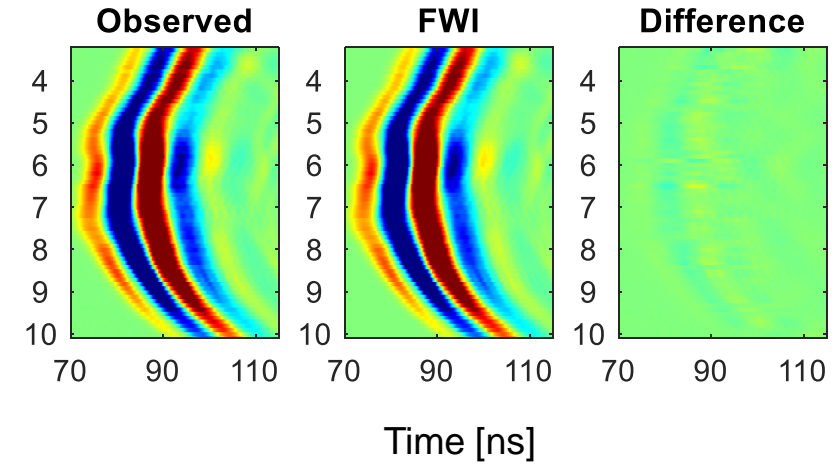
# ETHANOL TRACER RECOVERY



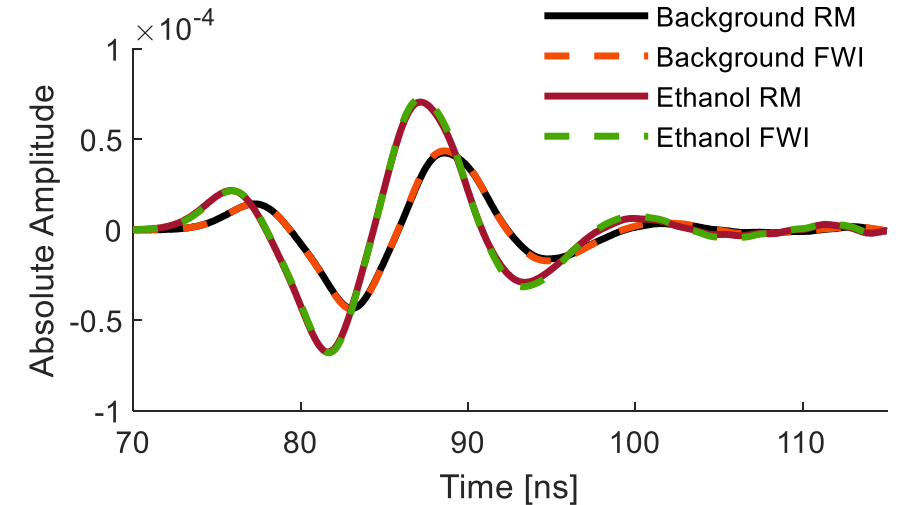
Starting model with FWI from the previous day results in better FWI model recovery.

Preferential paths from **permittivity** ( $r=0.9$ ) are imaged in high-resolution.

In ethanol intrusion both the amplitude and the phase are fitted successfully.



## Transmitter and receiver at 6 m

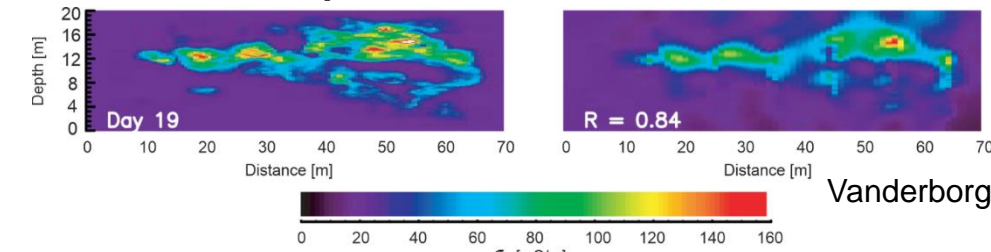


# SPECTRAL ANALYSIS OF TRACER PLUME DISTRIBUTION

## ERT study transport simulation

Real transport

ERT recovered



Vanderborght et al., 2005

## This study transport simulation – salt tracer (from $\sigma_b$ changes)

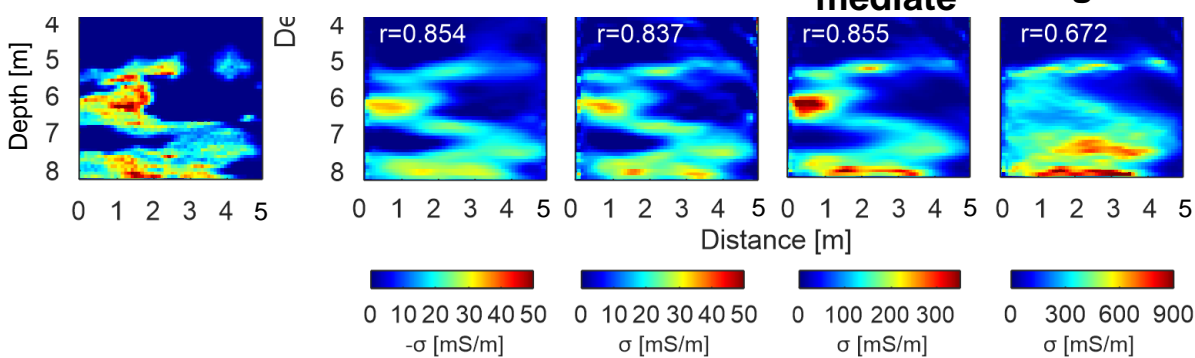
Real transport

Desalinated

Low

Inter-  
mediate

High

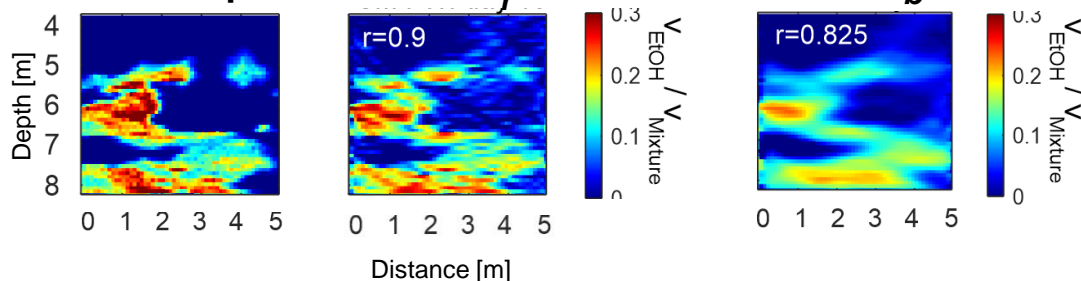


## This study transport simulation – ethanol tracer (from $\epsilon_r$ and $\sigma_b$ changes)

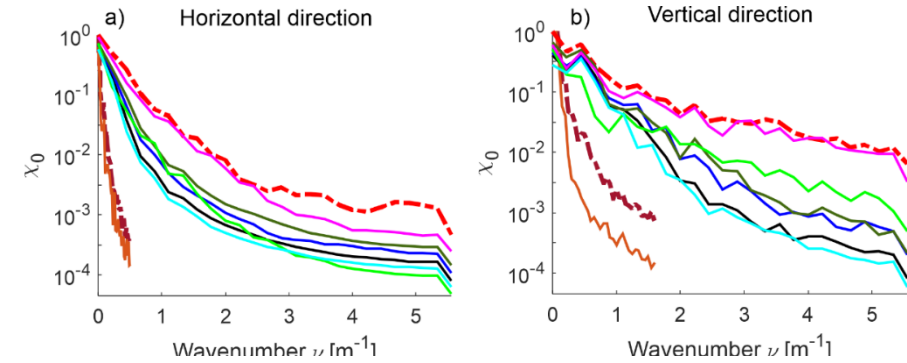
Real transport

From  $\epsilon_r$

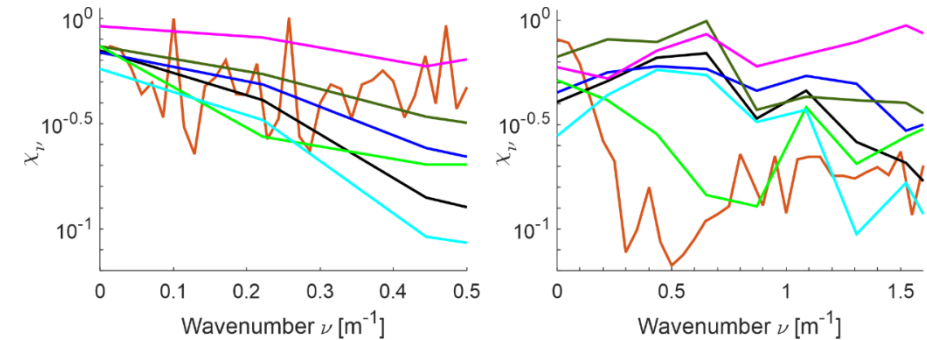
From  $\sigma_b$



$$\chi_0(u) = \text{PSD}_{\text{recovered}}(u) / \text{PSD}_{\text{transport}}(u=0)$$



$$\chi_\nu(u) = \text{PSD}_{\text{recovered}}(u) / \text{PSD}_{\text{transport}}(u)$$



--- Tracer transport simulation for ERT synthetic experiment  
 --- Inverted ERT  
 --- Tracer transport simulation for GPR FWI synthetic experiment  
 --- Desalinated    --- Low    --- Intermediate    --- High  
 --- Ethanol from  $\sigma$     --- Ethanol from  $\epsilon_r$

- Ethanol tracer distribution that is derived from the GPR FWI  $\epsilon_r$  recovery is almost overlapping that of true tracer distribution transport.
- The Intermediate tracer conductivity case shows the best reconstruction amongst the saline tracer scenarios.
- More spectral information recovered in the vertical direction with GPR FWI. That may be related to higher sampling (tx 0.2 and rx 0.1 m receiver spacing; ERT 0.6 m electrode vertical separation).

# CONCLUSION

- A high resolution transport model of the Krauthausen aquifer has been built to investigate the potential of time-lapse GPR FWI.
- Synthetic time-lapse GPR FWI allowed to investigate different inversion strategies which showed that higher accuracy of tracer imaging can be achieved by using a previous fay FWI models as SM for the current day.
- Time-lapse GPR FWI has the potential for high resolution imaging of tracer tests that cause changes in electrical conductivity and/or permittivity, with better potential for pemittivity, to delinaite flow paths of about 0.1 m thickness.
- Resolution of permittivity changes is better than conductivity and have the potential to delineate flow paths of about 0.1 m thickness.
- First experimental data for saline and heat tracer experiments show an effect on the GPR data.



# Thank you for your attention!

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## Acknowledgments:

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