# Reconciling atmospheric water uptake by hydrate forming salts

Bernadette Rosati,\*,†,¶ Andreas Paul,†,§ Emil Mark Iversen,† Andreas Massling,‡ and Merete Bilde\*,†

†Department of Chemistry, Aarhus University, Denmark ‡Department of Environmental Science, Aarhus University, Denmark

¶Faculty of Physics, University of Vienna, Austria

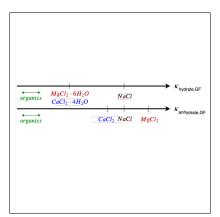
§now at: Institute of Energy and Climate Research, Forschungszentrum Jülich, Germany

E-mail: bernadette.rosati@chem.au.dk; bilde@chem.au.dk

Abstract

Magnesium and calcium chloride salts contribute to the global atmospheric aerosol burden via emission of sea spray and mineral dust. Their influence on aerosol hygroscopicity and cloud forming potential is important but uncertain with ambiguities between results reported in the literature. To address this, we have conducted measurements of the hygroscopic growth and critical supersaturation of dried, size selected nano-particles made from aqueous solution droplets of MgCl<sub>2</sub> and CaCl<sub>2</sub>, respectively, and compare experimentally derived values with results from state-of-the-art thermodynamic modelling. It is characteristic of both MgCl<sub>2</sub> and CaCl<sub>2</sub> salts that they bind water in the form of hydrates under a range of ambient conditions. We discuss how hydrate formation affects the particles' water uptake and provide an expression for hydrate correction factors needed in calculations of hygroscopic growth factors, critical supersaturations, and  $\kappa$  values of particles containing hydrate forming salts. We demonstrate the importance of accounting for hydrate forming salts when predicting hygroscopic properties of sea spray aerosol.

## 16 Graphical TOC Entry



# 18 Keywords

Hydrate forming salts, hygroscopicity, cloud activation, water activity

Atmospheric aerosols containing inorganic salts are ubiquitous in the Earth's atmosphere 20 and play a critical role in climate due to their interactions with sunlight and role in cloud 21 formation. Mineral dust is a major contributor to the atmospheric loading of inorganic salts and is mainly emitted from dry soils as a result of surface winds. 2-5 Mineral dust 23 is dominated by almost insoluble oxides and carbonates with low hygroscopicity. <sup>6</sup> During 24 atmospheric transit, surface reactions of these oxides and carbonates with acidic gases lead 25 to the formation of  $MgCl_2$  and  $CaCl_2$  salts.<sup>7–9</sup> These salts are soluble in water and thus 26 directly impact particle water uptake. <sup>10</sup> The oceans constitute another major emission source 27 of inorganic salt, releasing so-called sea spray aerosols (SSA) through wave breaking at the 28 ocean surface into the atmosphere. 11 SSA are comprised of a mixture of inorganic salts and 29 organic species, <sup>11–13</sup> where the relative ratio varies as a function of size. <sup>14–16</sup> The hygroscopic 30 nature of SSA is mainly driven by the inorganic fraction and although sodium chloride 31 (NaCl) makes up the major inorganic compound in SSA by mass, recent work has shown 32 that it is imperative to include the whole complexity of inorganic sea salt in assessment of 33 its hygroscopicity 17 and cloud forming potential. 18 Rasmussen et al. 19 pointed out that sea salt contains hydrate forming salts which affect volatility and hydrated forms of MgCl<sub>2</sub> and 35 CaCl<sub>2</sub> have been proposed to be responsible for lowering the hygroscopic potential of sea salt compared to pure NaCl. 17

While MgCl<sub>2</sub> is typically stable and solid in its hexahydrate state (MgCl<sub>2</sub>·6H<sub>2</sub>O) at 298 K and relative humidities (RH) above 3%, <sup>20</sup> the hydration state of solid CaCl<sub>2</sub> at 298 K is much less clear. The number of water molecules associated with CaCl<sub>2</sub> may be 0, 2, 4 or 6 for humidities in the range 0-28%. <sup>20</sup> As a further complication the tetrahydrate (CaCl<sub>2</sub>·4H<sub>2</sub>O) can exist in two allotropic forms with different solubilities <sup>21</sup> and several of the hydrates may exist in meta-stable states. <sup>22</sup>

Several recent studies have targeted the hygroscopicity and cloud activation potential of MgCl<sub>2</sub> and CaCl<sub>2</sub> salts<sup>6,9,23–26</sup> revealing ambiguities in how to address and present the water uptake of hydrate forming salts. We here suggest a transparent and clear way of report-

ing hygroscopicity and cloud activation potential for hydrate forming salts of atmospheric relevance and provide a comparison to results from a state-of-the-art thermodynamic model. 48 We performed two series of experiments focusing on water uptake by nano-meter sized 49 aerosol particles at sub- and supersaturated conditions of water vapor, respectively. These 50 were complemented by bulk water activity measurements. A full description of the experi-51 ments and tabular values of experimental results are provided in the supplementary informa-52 tion (SI). In short, aerosol particles were generated from aqueous salt solutions (CaCl<sub>2</sub>(aq) or 53 MgCl<sub>2</sub>(aq)) using an atomizer and dried by dilution with dry clean air and passage through 54 diffusion dryers (RH<10%). Water uptake was probed using a Humidified Tandem Differ-55 ential Mobility Analyzer (HTDMA) and a Cloud Condensation Nucleus counter (CCNc). 56 The hygroscopic growth factors (GF(RH)); defined as the ratio between humidified and 57 dry particle diameter) were determined directly by HTDMA measurements and indirectly 58 from measurements on bulk solution using a water activity meter. Figures 1a and 1c, present 59 measured GFs as a function of RH for MgCl<sub>2</sub> and CaCl<sub>2</sub> particles generated from MgCl<sub>2</sub> and CaCl<sub>2</sub> aqueous solutions, respectively, with initial dry diameters of 200 nm. For comparison, experimental values reported in the literature deploying a similar aerosol generation technique and measurement principle for 100 nm<sup>9</sup> and 50 nm<sup>25</sup> particles are shown. The results from this work confirm previously published values; at RH=80% Guo et al. 9 report GF=1.46 and Park et al. <sup>25</sup> GF=1.47, while we find GF=1.50 for MgCl<sub>2</sub>, thus agreeing within the estimated uncertainty in GF of 3%. For CaCl<sub>2</sub>, the discrepancies are larger yielding GF 66 values of 1.51, 9 1.48 25 and 1.69 (this study). All experimental studies show gradual increase 67 in GF(RH) with increasing RH and gradual decrease with decreasing humidity. The lack 68 of clear deliquescence and efflorescence indicates that the particles exist in an amorphous 69 solid phase state after drying. 9 Interestingly, clear deliquescence has been reported for larger 70 (micro-meter sized) particles  $^{24,27,28}$  of MgCl<sub>2</sub> and CaCl<sub>2</sub> and measuring the mass of a bulk 71 sample as function of RH. This confirms particle size to be a critical parameter in the phase 72 state behavior of atmospheric aerosol particles. <sup>29</sup> If particle size is a determining parameter, the actual size needed for crystal formation and thus deliquescence and efflorescence behavior of MgCl<sub>2</sub> and CaCl<sub>2</sub> particles is yet to be uncovered.

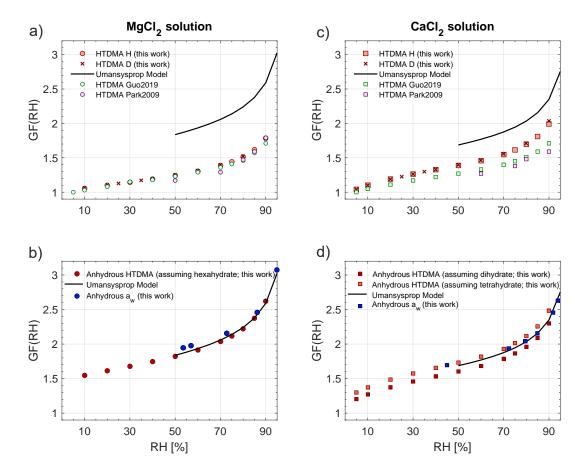


Figure 1: Hygroscopic growth factors vs. RH for dried particles made from aqueous solution of MgCl<sub>2</sub> and CaCl<sub>2</sub> salts. Panels a) and c) show experimental values from this work and literature.  $^{9,25}$  Regarding HTDMA results from this work, crosses indicate values recorded during dehydration (HTDMA D) while all other values stem from hydration branch measurements (HTDMA H). Panels b) and d) show HTDMA results (this work) after correction for hydrates along with GFs obtained from bulk water activity measurements (this work), derived following Zamora et al.  $^{30}$  Solid lines represent model  $^{31}$  predictions for anhydrous salts ( $D_{dry}$ =200 nm).

Predictions from a state-of-the-art thermodynamic model (UManSysProp<sup>31</sup>) are also displayed in Fig. 1. The model runs were executed assuming that the particles consisted of MgCl<sub>2</sub> and CaCl<sub>2</sub>, respectively, in their anhydrous states. Figure 1a and 1c reveal a considerable gap between observed and predicted GFs. We argue that water bound as hydrates in the dry salts is responsible for this difference. It was recently pointed out that airborne aqueous solution droplets containing MgCl<sub>2</sub> and/or CaCl<sub>2</sub> can form stable hydrates upon drying. <sup>19</sup> The number of water molecules associated with the dried salt depends on temperature and RH<sup>20,22</sup> and hence the hydration state of the salt dissolved in the aqueous solution is not necessarily representative for the hydration state of the dried nano-particles generated from the aqueous solution.

For MgCl<sub>2</sub> the stable hydration state at RH in the range 3-33% and a temperature of 298 K is MgCl<sub>2</sub>·6H<sub>2</sub>O.<sup>20</sup> The dry particle diameter as selected by the first DMA in the HTDMA system at these conditions is thus not representative of an anhydrous particle but a particle containing MgCl<sub>2</sub>·6H<sub>2</sub>O. To obtain the diameter  $(D_p(RH,a))$  and growth factor (GF(RH,a)) of the anhydrous salt particle, we correct the selected dry diameter in the first DMA  $(D_p(RH,h))$  to account for hydrate water by applying a hydrate correction factor  $(c_H)$ :

$$D_p(RH, a) = c_H \cdot D_p(RH, h),$$

$$GF(RH, a) = \frac{GF(RH, h)}{c_H},$$
(1)

where GF(RH,h) is the observed (hydrated) growth factor. As derived in the SI the hydrate correction factor  $c_H$  is calculated by assuming spherical particles with a specific molar mass (M) and density ( $\rho$ ):

$$c_H = \left(\frac{M_a \cdot \rho_h}{M_h \cdot \rho_a}\right)^{1/3}.$$
 (2)

Here the subscripts a and h refer to the anhydrous and hydrated salt, respectively. As seen in Fig. 1b, HTDMA and modeled data agree almost perfectly (within 1%) for MgCl<sub>2</sub> when the GF(RH) values are corrected for contributions of hydrate water. This result is further supported by the good agreement between GFs for the anhydrous salt derived from water activity (a<sub>w</sub>) measurements (this work) following the procedure described in Zamora et al. The hydration state of CaCl<sub>2</sub> in dried particles is more difficult to infer as several possibilities are plausible. Small changes in RH, e.g. from 9-21% to 21-28%, can change CaCl<sub>2</sub>·4H<sub>2</sub>O

to CaCl<sub>2</sub>·6H<sub>2</sub>O<sup>20</sup> at 298 K, and a change in temperature from 298 to 303 K can induce a hydration change from CaCl<sub>2</sub>·6H<sub>2</sub>O to CaCl<sub>2</sub>·4H<sub>2</sub>O.<sup>20,32</sup> It is thus challenging to access the 103 hydration state of CaCl<sub>2</sub> particles generated by drying of aqueous solution droplets as in 104 this work, in literature <sup>6,9,23,25,26</sup> and in the atmosphere. Figure 1d shows GF(RH) for CaCl<sub>2</sub> 105 corrected assuming dihydrate  $(2 \cdot H_2 O)$  and tetrahydrate  $(4 \cdot H_2 O)$  states. It is clear that a 106 correction is needed to reach agreement with the model output for anhydrous salt particles 107 but it is not possible to infer the actual hydration state as the best correlation depends on 108 the actual RH. This result likely reflects that we cannot exclude that individual particles 109 may contain several hydration states of the salt. Additionally, Fig. 1d shows GF(RH) val-110 ues retrieved from a<sub>w</sub> providing excellent agreement with model data. For completeness, 111 GF(80%) was measured as a function of initial dried particle diameter (40-200 nm) for both 112 salts and in all cases agreement with the model was only reached after application of the 113 hydrate correction factors (see SI). 114

Table 1 provides hydrate correction factors  $(c_H)$  for the different hydration states of MgCl<sub>2</sub> and CaCl<sub>2</sub>. Since in the atmosphere hydrate forming salts are likely present in their hydrated state, modelled hygroscopic growth factors based on the assumption of anhydrous salts should be corrected to account for the effect of hydrates using  $c_H$ .

Table 1: Hydrate correction factors for the atmospherically relevant hydration states of MgCl<sub>2</sub> and CaCl<sub>2</sub>, assuming spherical particles at T=298.15 K. Densities and molar masses used for this calculation can be found in the SI.

Compound	$c_H$
MgCl <sub>2</sub> ·6H <sub>2</sub> O	0.68
$CaCl_2 \cdot 2H_2O$	0.87
$CaCl_2 \cdot 4H_2O$	0.80
CaCl <sub>2</sub> ·6H <sub>2</sub> O	0.74

The particle ability to act as cloud condensation nuclei (CCN) is also affected by the salt hydration state. Figure 2 shows critical supersaturation (SS<sub>crit</sub>) vs. dry particle diameter for MgCl<sub>2</sub> and CaCl<sub>2</sub> particles, respectively. As in the case of the HTDMA experiments, the dried particle diameter likely represents the diameter of a particle containing hydrated

salts. The hydrate correction factor, introduced in Eq. 2 and presented in Table 1, can be used to calculate the diameter of the anhydrous particle from the diameter of the hydrated 124 particle (and vice versa). SS<sub>crit</sub> measured for dried particles made from MgCl<sub>2</sub> and CaCl<sub>2</sub> 125 aqueous solutions, respectively, are in relatively good agreement with previous findings 6,23,26 126 (Fig. 2a and c). For hydrated MgCl<sub>2</sub> particles with a diameter of 50 nm we find SS<sub>crit</sub>=0.57 127  $(\kappa=0.34^{33})$ , while Gaston et al.<sup>23</sup> found SS<sub>crit</sub>=0.48 for the same size  $(\kappa=0.47)$ . For hydrated 128 CaCl<sub>2</sub> Gaston et al.<sup>23</sup> report  $SS_{crit}=0.24$  for particles with a diameter of 75 nm ( $\kappa=0.56$ ), 129 while Sullivan et al.<sup>6</sup> find  $SS_{crit}=0.265$  ( $\kappa=0.46$ ). We obtained  $SS_{crit}=0.275$  ( $\kappa=0.42$ ) which 130 is in best agreement with Sullivan et al. 6 (see Fig. 2b). The most striking feature is again 131 the significant divergence in observed CCN activity of dried salt particles and model results 132 assuming anhydrous salts. Both literature studies assumed that CaCl<sub>2</sub> particles were present 133 in their dihydrate state, while MgCl<sub>2</sub> was in its hexahydrate state.<sup>6,23</sup> 134

So called intrinsic  $\kappa$  values can be calculated as: <sup>23</sup>

135

$$\kappa_{int} = \frac{\nu \cdot \rho_s \cdot M_w}{\rho_w \cdot M_s},\tag{3}$$

where  $\nu$  is the number of ions,  $\rho_s$  and  $\rho_w$  are the densities of solute and water and  $M_w$  and  $M_s$  are the molecular weights of solute and water.  $\kappa_{int}$  for  $CaCl_2 \cdot 2H_2O$  is 0.68 while experimental results assuming  $CaCl_2 \cdot 2H_2O$  are notably lower. 6,23 Gaston et al. 23 accounted for solution non-idealities using the van't Hoff factor and report on this basis a  $\kappa$  value of 0.589 for  $CaCl_2 \cdot 2H_2O$ . This is closer to their experimental result (0.56) but still higher than results from Sullivan et al. 6 and from this study .

Applying the hydrate correction factors to the dry diameters from our study yield the results presented in Fig. 2b and 2d. In case of MgCl<sub>2</sub>·6H<sub>2</sub>O we obtain excellent agreement with model data. In the case of CaCl<sub>2</sub> the tetrahydrate or hexahydrate hypothesis leads to the best agreement. In comparison, the HTDMA results where best agreement was obtained assuming dihydrate or tetrahydrates and could be caused by slightly differing RH when the aerosols were size selected. Nevertheless, the results emphasize that including some hydration

even if the precise state is not known still yields a smaller error than assuming anhydrous particles.

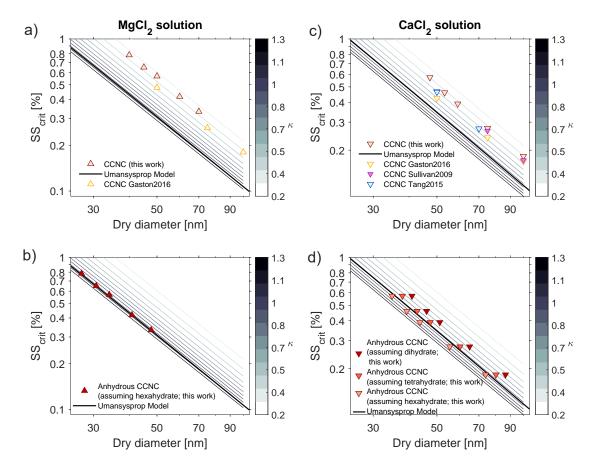


Figure 2: Critical super-saturations (SS<sub>crit</sub>) as a function of size for MgCl<sub>2</sub> and CaCl<sub>2</sub> together with  $\kappa$ -isolines. Additionally, available literature data<sup>6,23,26</sup> and model data are displayed. Panels a) and c) show data as directly obtained from the measurement, while b) and d) were corrected for hydrate contributions.

To reconcile hygroscopicity data gathered at sub- and supersaturated water vapour conditions, the hygroscopicity parameter  $\kappa^{33}$  is typically used. Figure 3 displays  $\kappa$  values obtained from HTDMA and CCNc measurements from this study and literature plotted vs. values calculated using the UManSysProp model. Generally, a large disagreement between experimental and modeled (assuming anhydrous particles) values is observed for both MgCl<sub>2</sub> and CaCl<sub>2</sub> (open symbols). Agreement between experimental and modelled  $\kappa$  values is obtained if particle diameters are modified to account for hydrate water following Eq. 2. We conclude

that the hydration state of MgCl<sub>2</sub> in dried salt nano-particles is 6 while for CaCl<sub>2</sub> hydration states in the range 2-6 are likely depending on drying procedure. Since results from subsatu-158 rated conditions suggest a hydration state of 2-4 and results from supersaturated conditions 159 suggest a hydration state of 4-6, we suggest use of the tetrahydrate state if no firm knowledge 160 on hydration state is available. Consistent with other studies for both organic and inorganic 161 compounds,  $^{33,34}$  there is a significant difference in  $\kappa$  calculated from subsaturated (GF(RH)) 162 and supersaturated (SS<sub>crit</sub>) conditions: in the case of MgCl<sub>2</sub> the model results yield  $\kappa_{\rm GF}$  of 163 1.89, while  $\kappa_{\rm SS_{crit}}$  is merely of 1.23, similarly a  $\kappa_{\rm GF}$  of 1.38 and  $\kappa_{\rm SS_{crit}}$  of 0.97 were found for 164  $CaCl_2$ . 165

For comparison, Fig. 3 additionally presents results for the notoriously hygroscopic NaCl salt, calculated from GF(RH) and SS<sub>crit</sub> values (black symbols). It is evident that the an-hydrous forms of MgCl<sub>2</sub> and CaCl<sub>2</sub> are actually extremely hygroscopic, with CaCl<sub>2</sub> reaching values close to those of NaCl ( $\kappa_{\text{GF,CaCl}_2}$ =1.38,  $\kappa_{\text{GF,NaCl}}$ =1.49) and MgCl<sub>2</sub> being much more hygroscopic than NaCl ( $\kappa_{\text{GF,MgCl}_2}$ =1.89,  $\kappa_{\text{GF,NaCl}}$ =1.49). When, however, considering the hydrated states of MgCl<sub>2</sub> and CaCl<sub>2</sub>, their  $\kappa$  reach comparable values around 0.5. It should be noted that NaCl is typically present in its anhydrous state except at temperatures below 273 K.<sup>22</sup>

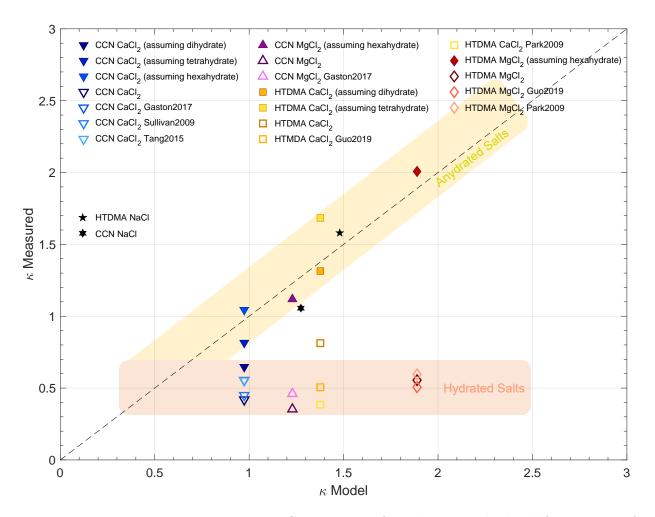


Figure 3: Hygroscopicity parameter  $\kappa$ : Comparison of  $\kappa$  values as calculated from HTDMA (taken at RH=90%) and CCNc (calculated as mean over all sizes, see SI) as well as literature data  $^{6,9,23,25,26}$  for MgCl<sub>2</sub> and CaCl<sub>2</sub>. The  $\kappa$ -Model data were obtained using the UManSysProp model. The dashed line represents the 1:1 line. All filled symbols were corrected for hydrates, while open symbols represent uncorrected data. Additionally, results for NaCl as measured and modeled from HTDMA and CCNc data are illustrated.

Figure 3 demonstrates the importance of hydrates for correct interpretation and pre-174 diction of hygroscopic growth and cloud droplet activation and shows that closure between 175 experimental data and state-of-the-art model predictions can be achieved when hydrates are 176 considered. These findings also underline that hygroscopicity of complex aerosols containing hydrate forming salts (such as sea salt) cannot be directly predicted by using models that 178 combine properties of anhydrous salts as this approach leads to an overestimation of the 179 hygroscopic growth as was shown in previous studies. 17

177

The results from this study can be used to predict the hygroscopicity of oceanic sea 181 salt. The major components of dried sea salt are NaCl, MgCl<sub>2</sub>, Na<sub>2</sub>SO<sub>4</sub> and CaCl<sub>2</sub>. <sup>35</sup> As 182 discussed above, NaCl is expected to be anhydrous at most ambient conditions. Na<sub>2</sub>SO<sub>4</sub> has 183 been shown to be present both in its anhydrous and decahydrate state.  $^{22}$  A previous study, 184 however, found that suspended  $\mathrm{Na_2SO_4}$  microparticles are anhydrous.<sup>36</sup> This is supported 185 by volatility experiments on dried mixed NaCl/Na<sub>2</sub>SO<sub>4</sub> particles that did not evaporate 186 at temperatures ranging from 50 to  $300^{\circ}\mathrm{C^{19}}$  and GF measurements for  $\mathrm{Na_{2}SO_{4}}$  that were 187 well comparable to theoretical calculations assuming anhydrates and Köhler theory. <sup>37</sup> Thus, 188  $Na_2SO_4$  is treated as anhydrous in the following. Table 2 presents estimated GF,  $\kappa_{GF}$  and 180  $\kappa_{SS_{crit}}$  values for two cases: (i) assuming sea salt consists of anhydrous salts only and (ii) 190 considering hydrated MgCl<sub>2</sub> and CaCl<sub>2</sub> salts in sea salt. By applying a volume weighted 191 mixing rule<sup>33</sup> sea salt GF<sub>a</sub> of 2.38,  $\kappa_{a,GF}$ =1.46 and  $\kappa_{a,SS_{crit}}$ =1.21 are found. When includ-192 ing hydrated MgCl<sub>2</sub> and CaCl<sub>2</sub>, GF<sub>h</sub> of 2.14,  $\kappa_{h,GF}=1.08$  and  $\kappa_{h,SS_{crit}}=0.89$  are found. GF 193 measurements of artificial sea salt particles yielded GF(90%)=2.09 and 2.19 ( $\kappa=0.96$  and 194 1.12) using a sea spray simulation tank and a nebulizer to generate sea salt particles, respec-195 tively,  $^{17}$  while  $\kappa$ =0.92 was found from CCNc measurements.  $^{38}$  The herein calculated sea salt 196 results including hydrated MgCl<sub>2</sub> and CaCl<sub>2</sub> are evidently in much better agreement with 197 the sea salt experimental data than calculations based on the anhydrous salts alone.

Table 2: Calculations of the properties of sea salt considering (i) anhydrous salts only and (ii) hydrated forms of MgCl<sub>2</sub> and CaCl<sub>2</sub>. The volume fractions of the anhydrous salts  $\epsilon_{va}$  are recalculated from the specifications by ASTM.<sup>35</sup> For case (i) anhydrous GFs,  $\kappa_{a,GF}$  and  $\kappa_{a,SS_{crit}}$  were calculated using the UManSysProp model. Hydrated GFs,  $\kappa_{h,GF}$  and  $\kappa_{h,SS_{crit}}$  were calculated assuming that NaCl and Na<sub>2</sub>SO<sub>4</sub> were anhydrous, while MgCl<sub>2</sub>·6H<sub>2</sub>O and CaCl<sub>2</sub>·4H<sub>2</sub>O were present. The volume fractions of hydrated sea salt  $\epsilon_{vh}$  were recalculated from the anhydrous case assuming volume additivity (see SI). GFs stem from measurements at D<sub>dry</sub>=200 nm and RH=90%, while  $\kappa_{SS_{crit}}$  were calculated as mean over all measured sizes.

Anhydrous sea salt				Partly hydrated sea salt					
Compound	$\epsilon_{va}$ [%]	$GF_a$	$\kappa_{a,GF}$	$\kappa_{a,SS_{crit}}$	Compound	$\epsilon_{vh}$ [%]	$GF_h$	$\kappa_{h,GF}$	$\kappa_{h,SS_{crit}}$
NaCl	72.5	2.40	1.48	1.27	NaCl	53.8	2.40	1.48	1.27
$\mathrm{MgCl}_2$	14.3	2.59	1.89	1.16	$MgCl_2 \cdot 6H_2O$	34.1	1.79	0.56	0.35
$Na_2SO_4$	9.7	1.93	0.72	0.88	$Na_2SO_4$	7.2	1.93	0.72	0.88
$CaCl_2$	3.5	2.35	1.38	0.92	CaCl <sub>2</sub> ·4H <sub>2</sub> O	5.0	1.99	0.81	0.42
mixture	100	2.38	1.46	1.21	mixture	100	2.14	1.08	0.89

In summary, the hygroscopic and cloud activation properties of MgCl<sub>2</sub> and CaCl<sub>2</sub> are 199 dictated by their hydration state, which substantially lowers their ability to take up water. 200 The behavior at sub- and supersaturated water vapour conditions differs remarkably, yielding 201 higher  $\kappa$  values when calculated from the hygroscopic growth compared to when calculated 202 from the SS<sub>crit</sub>. When accounting for the hydration state of the salts, very good agreement 203 can be achieved between measured and modeled data. This work demonstrates, that it 204 is important to account for hydrate formation when predicting hygropscopic properties of 205 complex aerosols containing hydrate forming salts, such as sea salt aerosols. 206

### 207 Acknowledgement

This research was supported by the Austrian Science Fund (FWF: J 3970-N36) and Aarhus University. The authors thank the Villum Foundation for support for instrumentation (HT-DMA and pre-humidifier) involved in this study.

# <sup>211</sup> Supporting Information Available

The Supplementary Information is available free of charge. Details of the calculation procedure for anhydrous particle size, experimental materials and methods as well as figures and tabular values for the hygroscopicity results from HTDMA, CCNc and water activity measurements are presented.

#### 216 References

- 217 (1) IPCC, Climate Change 2013: The Physical Science Basis. Contribution of Working
  218 Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate
  219 Change.; Cambridge University Press: Cambridge, United Kingdom and New York,
  220 NY, USA, 2013; Chapter SPM, pp 1–30, Stocker, T.F. and Qin, D. and Plattner, G.-K.
  221 and Tignor, M. and Allen, S.K. and Boschung, J. and Nauels, A. and Xia, Y. and Bex,
  222 V. and Midgley, P.M.
- <sup>223</sup> (2) Tegen, I.; Fung, I. Contribution to the atmospheric mineral aerosol load from land <sup>224</sup> surface modification. *Journal of Geophysical Research: Atmospheres* **1995**, *100*, 18707– <sup>225</sup> 18726.
- 226 (3) Chou, C.; Formenti, P.; Maille, M.; Ausset, P.; Helas, G.; Harrison, M.; Osborne, S.

  Size distribution, shape, and composition of mineral dust aerosols collected during

  the African Monsoon Multidisciplinary Analysis Special Observation Period 0: Dust

  and Biomass-Burning Experiment field campaign in Niger, January 2006. Journal of

  Geophysical Research: Atmospheres 2008, 113.
- (4) Formenti, P.; Elbert, W.; Maenhaut, W.; Haywood, J.; Andreae, M. O. Chemical composition of mineral dust aerosol during the Saharan Dust Experiment (SHADE) airborne campaign in the Cape Verde region, September 2000. *Journal of Geophysical Research: Atmospheres* 2003, 108.

- 235 (5) Nowak, S.; Lafon, S.; Caquineau, S.; Journet, E.; Laurent, B. Quantitative study of the
  236 mineralogical composition of mineral dust aerosols by X-ray diffraction. *Talanta* **2018**,
  237 186, 133 139.
- 238 (6) Sullivan, R. C.; Moore, M. J. K.; Petters, M. D.; Kreidenweis, S. M.; Roberts, G. C.;
  239 Prather, K. A. Effect of chemical mixing state on the hygroscopicity and cloud nucle240 ation properties of calcium mineral dust particles. *Atmospheric Chemistry and Physics*241 **2009**, 9, 3303–3316.
- <sup>242</sup> (7) Usher, C. R.; Michel, A. E.; Grassian, V. H. Reactions on Mineral Dust. *Chemical Reviews* **2003**, *103*, 4883–4940.
- 244 (8) Santschi, C.; Rossi, M. J. Uptake of CO2, SO2, HNO3 and HCl on Calcite (CaCO3)
  245 at 300 K:âĂL Mechanism and the Role of Adsorbed Water. The Journal of Physical
  246 Chemistry A 2006, 110, 6789–6802.
- <sup>247</sup> (9) Guo, L.; Gu, W.; Peng, C.; Wang, W.; Li, Y. J.; Zong, T.; Tang, Y.; Wu, Z.;

  Lin, Q.; Ge, M. et al. A comprehensive study of hygroscopic properties of calcium
  and magnesium-containing salts: implication for hygroscopicity of mineral dust and

  sea salt aerosols. Atmospheric Chemistry and Physics 2019, 19, 2115–2133.
- (10) Tang, M.; Cziczo, D. J.; Grassian, V. H. Interactions of Water with Mineral Dust
   Aerosol: Water Adsorption, Hygroscopicity, Cloud Condensation, and Ice Nucleation.
   Chemical Reviews 2016, 116, 4205–4259, PMID: 27015126.
- <sup>254</sup> (11) Lewis, E.; Schwartz, S. Sea Salt Aerosol Production: Mechanisms, Methods, Measure-<sup>255</sup> ments and Models; American Geophysical Union, Washington DC, 2004; p 413.
- 256 (12) Blanchard, D. C. The ejection of drops from the sea and their enrichment with bacteria 257 and other materials: A review. *Estuaries* **1989**, *12*, 127–137.

- <sup>258</sup> (13) Leck, C.; Bigg, E. K. Biogenic particles in the surface microlayer and overlaying atmo-<sup>259</sup> sphere in the central Arctic Ocean during summer. *Tellus B: Chemical and Physical* <sup>260</sup> *Meteorology* **2005**, *57*, 305–316.
- (14) Cochran, R. E.; Jayarathne, T.; Stone, E. A.; Grassian, V. H. Selectivity Across the
   Interface: A Test of Surface Activity in the Composition of Organic-Enriched Aerosols
   from Bubble Bursting. The Journal of Physical Chemistry Letters 2016, 7, 1692–1696.
- (15) Facchini, M.; Rinaldi, M.; Decesari, S.; Carbone, C.; Finessi, E.; Mircea, M.; Fuzzi, S.;
   Ceburnis, D.; Flanagan, R.; Nilsson, E. et al. Primary submicron marine aerosol dominated by insoluble organic colloids and aggregates. Geophysical Research Letters 2008,
   35.
- 268 (16) Salter, M.; Hamacher-Barth, E.; Leck, C.; Werner, J.; Johnson, C.; Riipinen, I.; Nils269 son, E.; Zieger, P. Calcium enrichment in sea spray aerosol particles. *Geophysical Re-*270 search Letters **2016**, 43, 8277–8285.
- (17) Zieger, P.; Väisänen, O.; Corbin, J.; Partridge, D.; Bastelberger, S.; Mousavi-Fard, M.;
   Rosati, B.; Gysel, M.; Krieger, U.; Leck, C. et al. Revising the hygroscopicity of inorganic sea salt particles. *Nature Communications* 2017, 8.
- 274 (18) King, S.; Butcher, A.; Rosenoern, T.; Coz, E.; Lieke, K.; De Leeuw, G.; Nilsson, E.;
  275 Bilde, M. Investigating primary marine aerosol properties: CCN activity of sea salt
  276 and mixed inorganic-organic particles. Environmental Science and Technology 2012,
  277 46, 10405–10412.
- 278 (19) Rasmussen, B. B.; Nguyen, Q. T.; Kristensen, K.; Nielsen, L. S.; Bilde, M. What
  279 controls volatility of sea spray aerosol? Results from laboratory studies using artificial
  280 and real seawater samples. *Journal of Aerosol Science* **2017**, 107, 134 141.
- 281 (20) Kelly, J. T.; Wexler, A. S. Thermodynamics of carbonates and hydrates related to

- heterogeneous reactions involving mineral aerosol. Journal of Geophysical Research:

  Atmospheres 2005, 110.
- 284 (21) Raj, G. Phase Rule, 5th Edition; GOEL Publishing House, Meerut, 2014; p 214.
- <sup>285</sup> (22) Lide, D. CRC Handbook of Chemistry and Physics, 85th Edition; CRC Press, Boca Raton, FL, 2005; p 2661.
- (23) Gaston, C. J.; Pratt, K. A.; Suski, K. J.; May, N. W.; Gill, T. E.; Prather, K. A. Laboratory Studies of the Cloud Droplet Activation Properties and Corresponding Chemistry
   of Saline Playa Dust. Environmental Science & Technology 2017, 51, 1348–1356.
- <sup>290</sup> (24) Gupta, D.; Eom, H.-J.; Cho, H.-R.; Ro, C.-U. Hygroscopic behavior of NaCl-MgCl<sub>2</sub>

  mixture particles as nascent sea-spray aerosol surrogates and observation of efflorescence

  during humidification. *Atmospheric Chemistry and Physics* **2015**, *15*, 11273–11290.
- <sup>293</sup> (25) Park, K.; Kim, J.-S.; Miller, A. L. A study on effects of size and structure on hygroscopicity of nanoparticles using a tandem differential mobility analyzer and TEM. *Journal* of Nanoparticle Research 2009, 11, 175–183.
- <sup>296</sup> (26) Tang, M. J.; Whitehead, J.; Davidson, N. M.; Pope, F. D.; Alfarra, M. R.; McFig<sup>297</sup> gans, G.; Kalberer, M. Cloud condensation nucleation activities of calcium carbonate
  <sup>298</sup> and its atmospheric ageing products. *Phys. Chem. Chem. Phys.* **2015**, *17*, 32194–32203.
- <sup>299</sup> (27) Ha, Z.; Chan, C. K. The Water Activities of MgCl2, Mg(NO3)2, MgSO4, and Their <sup>300</sup> Mixtures. Aerosol Science and Technology **1999**, 31, 154–169.
- (28) Cohen, M. D.; Flagan, R. C.; Seinfeld, J. H. Studies of concentrated electrolyte solutions
   using the electrodynamic balance. 1. Water activities for single-electrolyte solutions.
   The Journal of Physical Chemistry 1987, 91, 4563–4574.
- <sup>304</sup> (29) Cheng, Y.; Su, H.; Koop, T.; Mikhailov, E.; Pöschl, U. Size dependence of phase transitions in aerosol nanoparticles. *Nature Communications* **2015**, 6.

- of common organic aerosol solutes, including humic substances, as derived from water activity measurements. *Journal of Geophysical Research: Atmospheres* **2011**, *116*.
- 309 (31) Topping, D.; Barley, M.; Bane, M. K.; Higham, N.; Aumont, B.; Dingle, N.; McFig-310 gans, G. UManSysProp v1.0: an online and open-source facility for molecular prop-311 erty prediction and atmospheric aerosol calculations. *Geoscientific Model Development* 312 **2016**, 9, 899–914.
- 313 (32) Ropp, R. Encyclopedia of the Alkaline Earth Compounds, 1st Edition; Elsevier, 2012; p 1216.
- 315 (33) Petters, M. D.; Kreidenweis, S. M. A single parameter representation of hygroscopic growth and cloud condensation nucleus activity. *Atmospheric Chemistry and Physics* 2007, 7, 1961–1971.
- Whitehead, J. D.; Irwin, M.; Allan, J. D.; Good, N.; McFiggans, G. A meta-analysis of particle water uptake reconciliation studies. *Atmospheric Chemistry and Physics* **2014**, 14, 11833–11841.
- 321 (35) D1141-98, A. Standard Practice for the Preparation of Substitute Ocean Water; ASTM

  322 International: West Conshohocken, PA, USA, 2013.
- 323 (36) Tang, I. N.; Fung, K. H.; Imre, D. G.; Munkelwitz, H. R. Phase Transformation and
  324 Metastability of Hygroscopic Microparticles. Aerosol Science and Technology 1995, 23,
  325 443–453.
- 326 (37) Hu, D.; Qiao, L.; Chen, J.; Ye, X.; Yang, X.; Cheng, T.; Fang, W. Hygroscopicity of

  Inorganic Aerosols: Size and Relative Humidity Effects on the Growth Factor. Aerosol

  and Air Quality Research 2010, 10, 255–264.

329 (38) Nguyen, Q. T.; Kjær, K. H.; Kling, K. I.; Boesen, T.; Bilde, M. Impact of fatty acid 330 coating on the CCN activity of sea salt particles. *Tellus B: Chemical and Physical* 331 *Meteorology* **2017**, *69*, 1304064.