

1 Paul Jozef Crutzen  
2 3 December 1933 – 28 January 2021  
3 Elected ForMemRS 2006

4 Rolf Müller<sup>1\*</sup>

<sup>1</sup>*Institute of Energy and Climate Research (IEK-7), Forschungszentrum Jülich, 52425 Jülich, Germany*

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6 **Abstract**

7 Paul J. Crutzen was both a warmhearted person and a brilliant scientist. His  
8 research interests were broad, encompassing topics of relevance in the mesosphere,  
9 the stratosphere and the troposphere. He made fundamental scientific contributions  
10 to a wide range of topics in the science of all these atmospheric regions. For example,  
11 he first described the NO<sub>x</sub> driven ozone loss cycle in the stratosphere, he contributed  
12 key ideas on how to explain the “ozone hole” and he made fundamental discoveries  
13 on the impact of biomass-burning on the atmosphere. Understanding and combating  
14 the origins of air pollution and climate change were driving motivations for his  
15 life’s work. Further, he pioneered the concept that is known as “nuclear winter”,  
16 he initiated the resumption of discussion on geoengineering and coined the term  
17 “Anthropocene”. With all this focus on science, he would not forget the importance  
18 of private life, both as regards his family and his colleagues and students. In 1995,  
19 together with Mario J. Molina, and F. Sherwood Rowland, Paul was awarded the Nobel  
20 Prize in Chemistry for ground-breaking work “in atmospheric chemistry, particularly  
21 concerning the formation and decomposition of ozone”. Finally, Paul’s work not  
22 only had a profound influence on the scientific world, but also on different aspects  
23 of environmental politics throughout many countries. His works on the impact of  
24 human activities on the atmosphere and climate have been influential in the past and  
25 are influential today. They provide a beacon of knowledge for the next generations of  
26 scientists and environmental policymakers.

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\*ro.mueller@fz-juelich.de

<sup>27</sup> **A portrait of Paul Crutzen**



Figure 1: Portrait of Paul Crutzen for the Royal Society. (Picture courtesy Royal Society.)

## 1 Early life and education



Figure 2: Paul Crutzen with his father. (Picture courtesy of Ilona Crutzen.)

29 Paul Jozef Crutzen was born in Amsterdam on 3 December 1933; he was the son of  
30 Anna Gurk and Jozef Crutzen (Fig. 2). His younger sister Elisabeth Helena (Lies) Crutzen  
31 was born on 29 February 1936; she lived all her life in Amsterdam. She passed away in  
32 2003. His mother's parents were of mixed German and Polish origin and lived in the Ruhr  
33 region in Germany. His father had relatives in the Netherlands, Germany, and Belgium.  
34 Through his family and relatives, Paul inherited a cosmopolitan view of the world at an  
35 early age.

36 Paul had a hard youth; his six years of elementary school largely overlapped with  
37 the Second World War. Particularly horrible was the cold winter 1944-1945 known as  
38 "hongerwinter" (winter of famine). In this winter many Dutch died of hunger and disease,  
39 including several of Paul's schoolmates (reference 30) <sup>1</sup>.

40 In 1946 Paul entered the "Hogere Burgerschool" (Higher Citizen School), which he  
41 finished in June 1951. His academic focus lay in the natural sciences, but he also had  
42 to become proficient in French, English, and German. During these years at school, he

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<sup>1</sup>Numbers in this form refer to the bibliography at the end of the text.

43 spent considerable time doing sport: football, cycling, and – his greatest passion – long  
44 distance skating on the Dutch canals and lakes. All through his life, he kept a close eye on  
45 the success of the Dutch skating team at international sports events.

46 Because of a heavy fever, his grades in the final exam of the Higher Citizen School were  
47 not sufficient to qualify him for a university stipend. But as he did not want to continue  
48 to be a financial burden on his parents, he chose to attend the “Middelbare Technische  
49 School” (middle technical school) to train as a civil engineer, which allowed him to earn  
50 a modest salary in the second year of school (30, 45).



Figure 3: Terttu and Paul Crutzen on their wedding day in Amsterdam on 14 February 1958. (Picture courtesy of Ilona Crutzen.)

51 From the summer of 1954 until February 1958, with a 21-month interruption for  
52 compulsory military service, Paul worked at the Bridge Construction Bureau of the City  
53 of Amsterdam. During this time, on a vacation trip to Switzerland, he met Terttu Soininen,  
54 a student of Finnish history and literature at the University of Helsinki. They married  
55 in February 1958 (Fig. 3) and moved to Sweden, to the small town Gävle (about 200 km  
56 north of Stockholm) where Paul had found a job in a building construction bureau. In  
57 December of that year their daughter Ilona was born and in March 1964, her sister Sylvia.  
58 Paul and Terttu now have three grand children.

## 59 **2 Starting an academic career**

### 60 **2.1 Stockholm**

61 The opportunity for an academic career arose in 1958, when Paul Crutzen saw an advertise-  
62 ment in a Swedish newspaper by the Department of Meteorology of Stockholm Höögskola  
63 (now Stockholm University) announcing an opening for a computer programmer. He had

no experience in such work, but applied for the job and was chosen (30, 45). So on 1 July 1959 Terttu and Paul moved to Stockholm. This was the start of Paul's scientific career.

At the Meteorology Institute of Stockholm University (MISU) he met many eminent scientists; the director of the institute was Bert Bolin, who later became director of the Intergovernmental Panel on Climate Change (IPCC). Bert Bolin and Georg Witt, an expert on the aeronomy of the upper atmosphere, were Paul's supervisors at MISU. Whilst at MISU, he met his fellow student Henning Rodhe, now (retired) Professor in Chemical Meteorology at MISU, who became a close friend for life.

At that time, Stockholm University housed a computer called BESK, which was among the fastest in the world. However, computer programming was very different from what it is today; coding had to be done in machine language and not even the first variants of Fortran were available. But along with this work Paul followed his scientific interest; in 1963 he could thus fulfil the requirements for a "filosofie kandidat" (Master of Science) degree and, around 1965, got the task of helping J. R. Blankenship, a visiting scientist from the US, to develop a numerical model of the distribution of oxygen in different structural forms in the stratosphere, mesosphere, and lower thermosphere. This was the beginning of Paul's research in atmospheric chemistry and also resulted in his first scientific paper (1). Stratospheric ozone also became the topic of his "filosofie licentiat thesis" (PhD thesis). It is interesting to note that at that time at MISU, the issue of human impact on the atmosphere through "acid rain" was a hot topic, but that Paul decided to stay away initially from research addressing human impact on the atmosphere and decided "to do pure science related to natural processes and therefore [ . . . ] picked stratospheric ozone" (30) – this decision would be revisited later.

## 2.2 Filosofie Doctor

In May 1973, Paul Crutzen submitted his inaugural dissertation entitled "On the Photochemistry of Ozone in the Stratosphere and Troposphere and Pollution of the Stratosphere by High-Flying Aircraft" to the Faculty of Natural Sciences of Stockholm University. His thesis was based on four publications (references 3, 4, 5 and 6, with the last paper being in press at the time of submission).

The exam for the "filosofie doctor" took place in lecture hall 3, on 8 May 1973 at ten in the morning. The title page of his thesis is reprinted here (Fig. 4). At that time it was obligatory to have "opponents" of the candidate; in Paul's exam these were Dr. John Houghton (later Sir John) and Dr. Richard Wayne of the University of Oxford, who wore their university gowns on the occasion; Figure 5 shows a picture of Paul together with Dr. John Houghton (wearing an Oxford gown) and Dr. Richard Wayne (wearing a Cambridge gown) after the exam.

## 3 The chemical balance of the ozone layer

Paul Crutzen was the first to suggest that reactions catalysed by NO and NO<sub>2</sub> control the ozone concentrations in the middle stratosphere (reference 3); in 1970 he put forward the

ON THE PHOTOCHEMISTRY OF OZONE  
IN THE STRATOSPHERE AND TROPOSPHERE  
AND POLLUTION OF THE STRATOSPHERE  
BY HIGH-FLYING AIRCRAFT

INAUGURAL DISSERTATION

BY DUE PERMISSION  
OF THE FACULTY OF MATHEMATICS  
AND NATURAL SCIENCES  
OF THE UNIVERSITY OF STOCKHOLM  
SUBMITTED TO PUBLIC DISCUSSION  
IN LECTURE HALL 3, KUNGSTENSGATAN 45,  
MAY 8, 1973, AT 10 a.m.  
FOR THE DEGREE OF DOCTOR OF PHILOSOPHY

BY  
PAUL CRUTZEN  
FIL. LIC.

STOCKHOLM 1973  
TRYCKERIBOLAGET IVAR HÆGGSTRÖM AB

Figure 4: The title page of the inaugural dissertation of Paul Crutzen submitted in May 1973. This was one of the last occasions in which the classical “Filosofie Doktor” (similar to the Habilitation in other countries) was awarded (30).

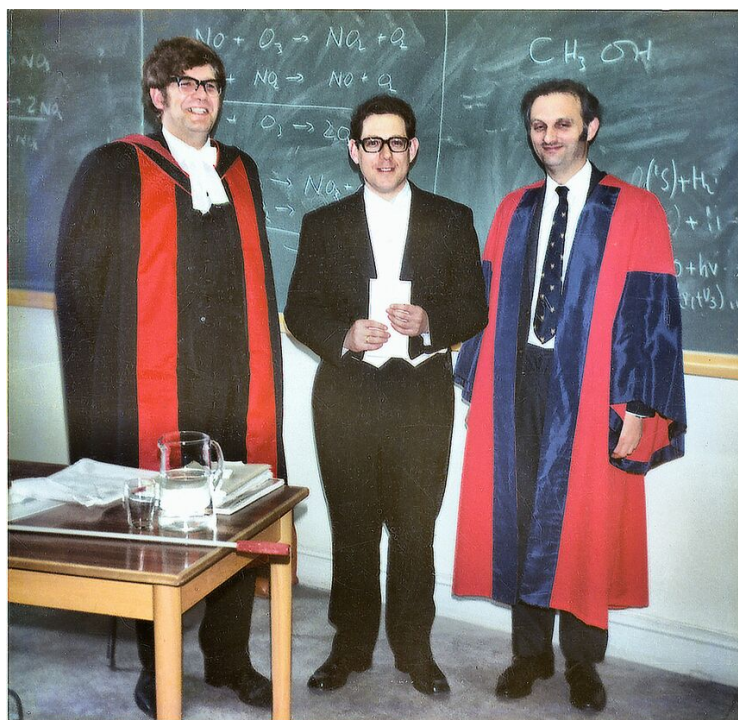
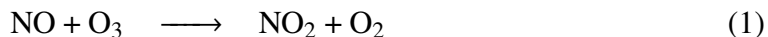


Figure 5: After Paul's exam in Stockholm in 1973. From left to right: Richard P. Wayne, Paul Crutzen, John Houghton (later Sir John). Note reactions (1) and (2) on the blackboard. Picture was taken with the camera of R. P. Wayne, probably by G. Witt. (Picture courtesy of Richard P. Wayne, see also reference 45.)

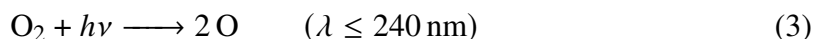
103 following catalytic cycle:



104 where the sum of NO and NO<sub>2</sub> is defined as NO<sub>x</sub> = NO + NO<sub>2</sub> and O indicates an oxygen  
105 atom in ground-state (O(<sup>3</sup>P)).

106 The discovery of the mechanism was a major achievement and an important step  
107 towards facilitating a quantitative description of the chemistry of the stratospheric ozone  
108 layer. It formed part of the basis of Paul being awarded the Nobel Prize in Chemistry in  
109 1995. In 1970, many uncertainties regarding stratospheric nitrogen remained; in particular  
110 the first measurements of nitrogen compounds in the stratosphere had just become available  
111 (Murcray et al., 1968; Rhine et al., 1969). Today it is established that reactions (1) and  
112 (2) are the most important ozone loss reactions in the stratosphere in the altitude region  
113 between about 25 and 45 km (e.g. Portmann et al., 2012, see also references 29, 30).

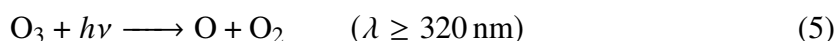
114 In the early 1970s, the tools for numerical simulations were one-dimensional column  
115 models for mid-latitude conditions (30°N), with about a 1 km vertical resolution, where  
116 the entire atmospheric transport was described by an empirical vertical eddy diffusion  
117 coefficient ( $K_z$ ) (Brasseur, 2019). Research on the chemical balance of the ozone layer had  
118 started with the British scientist Sydney Chapman (1930) who proposed that the formation  
119 of ozone is initiated by the photolysis of O<sub>2</sub>, by solar radiation at short wavelengths



120 followed by the association process



121 (where M is a collision partner such as atmospheric N<sub>2</sub>, O<sub>2</sub> or CO<sub>2</sub> that can remove energy  
122 released in the association reaction). The photodissociation of O<sub>3</sub> produces O



123 but does not constitute a loss of O<sub>3</sub> because there is a steady state between O and O<sub>3</sub> (30).  
124 Therefore, “odd oxygen” O<sub>x</sub> is defined as O<sub>x</sub> = O + O<sub>3</sub>. Destruction of odd oxygen occurs  
125 through the reaction



126 Both the production and destruction of an odd oxygen species are equivalent to the pro-  
127 duction and destruction of an ozone molecule. This scheme, referred to as the “Chapman  
128 reactions” was considered for many years to be the theoretical basis for a quantitative de-  
129 scription of stratospheric ozone, including its vertical distribution in the stratosphere (see  
130 e.g. reference 30). However, this scheme (Chapman, 1930) greatly overestimates ozone  
131 concentrations compared to measurements and can not explain the observed latitudinal  
132 distribution of ozone (Brasseur, 2019).

133 Twenty years later, Bates and Nicolet (1950) suggested that reactions involving OH  
134 and HO<sub>2</sub> radicals should lead to significant catalytic ozone loss in the mesosphere and  
135 thermosphere. Further studies focused on the impact of HO<sub>x</sub>-catalysed reactions (HO<sub>x</sub> =  
136 OH + HO<sub>2</sub> + H) on stratospheric ozone and new measurements of the rate constants for



the Chapman reactions were conducted (e.g. Müller, 2009; Brasseur, 2019, and references therein; see also reference 30).

Paul, in his “filosofie licentiat” thesis of 1968 concluded that these earlier theories could not explain the vertical distribution of ozone above 25 km (and in the troposphere) and that other photochemical processes beyond OH and HO<sub>2</sub> radicals should be introduced into the theory of stratospheric ozone (2, 30). Two years later, in 1970, Paul demonstrated that NO<sub>x</sub> species (i.e. reactions (1) and (2) are essential for the chemical balance of the stratospheric ozone layer (3).

In summer 1969 Paul had joined the Department of Atmospheric Physics at the Clarendon Laboratory of Oxford University as a postdoctoral fellow. He stayed there for a two year period and returned to Stockholm in July 1971. In Oxford, the head of the research group, Sir John Houghton, supported Paul’s idea about the potential role of NO<sub>x</sub> in stratospheric ozone chemistry. At that time the first measurements of nitrogen compounds in the atmosphere and their implications for the source of nitrogen in the stratosphere were a focus of research (see e.g. references 4 and 30); we know much more today (sections 4.2 and 4.3).

## 4 The impact of nitrogen on stratospheric ozone

### 4.1 Supersonic transport

In a paper on the vertical distributions of trace gases in the O<sub>x</sub>/NO<sub>x</sub>/HO<sub>x</sub>/HNO<sub>x</sub> system in 1971 (4) Paul Crutzen discussed the difficulties of doing model calculations in the absence of key information on chemical rate coefficients. But he also stated that the increase of NO<sub>x</sub> caused by the planned supersonic transport in the stratosphere should lead to observable changes in stratospheric ozone.

In the early 1970s a controversy about supersonic transport had started in the United States. At that time, large fleets of stratospheric supersonic aircraft were considered (US: Boeing, Britain/France: Concorde, Soviet Union: Tupolev) and a fleet of 500 supersonic planes seemed a reasonable estimate. The concern was an enhanced catalytic ozone destruction. Originally ozone destruction by OH and HO<sub>2</sub> radicals (resulting from the release of water vapour in the engine exhausts) was considered, but soon it was realised that the catalytic destruction of ozone by NO<sub>x</sub> posed a much greater threat to the ozone layer (Johnston, 1971).

In 1972, Paul made an important statement on the issue of supersonic transport (5). He made this statement despite (or perhaps because of) the limited knowledge on stratospheric processes at this time. He said that the “minimum requirement is therefore that extensive supersonic air traffic should not take place in the stratosphere before reliable predictions can be made of the possible environmental consequences of such operations”.

### 4.2 Solar Proton Events

In August 1972, a major solar proton event occurred, during which, within a few hours, large quantities of NO were produced at high latitudes (17). Paul showed, together with I. S. A. Isaksen and G. C. Reid, that solar proton events should lead to a depletion of

177 stratospheric ozone (9). A substantial depletion of ozone was indeed observed after the  
 178 solar proton event in 1972 (about 20% between 40 and 45 km) and could be reproduced by  
 179 model calculations by Susan Solomon and Paul (17); these calculations also considered  
 180 chlorine chemistry. Thus, the analysis of a natural phenomenon provided further support  
 181 for the impact of catalytic reactions involving  $\text{NO}_x$  on stratospheric ozone (reactions 1 and  
 182 2).

### 183 4.3 Nitrogen fertilisers

184 Paul also established that the biosphere has a direct impact on stratospheric ozone; the  
 185 biosphere (i.e. the soil) is a source of  $\text{N}_2\text{O}$  and therefore a source of stratospheric  $\text{NO}_x$   
 186 (4, 11). In the early 1970s many aspects of nitrogen chemistry in the stratosphere were  
 187 unknown or at least uncertain. However, it soon became clear that the reaction



188 constitutes an important source of  $\text{NO}_x$  in the stratosphere (e.g., McElroy and McConnell,  
 189 1971, see also reference 4). The dominant mechanism for loss of  $\text{N}_2\text{O}$  in the stratosphere  
 190 is photolysis



191 which, however, does not produce  $\text{NO}_x$ .

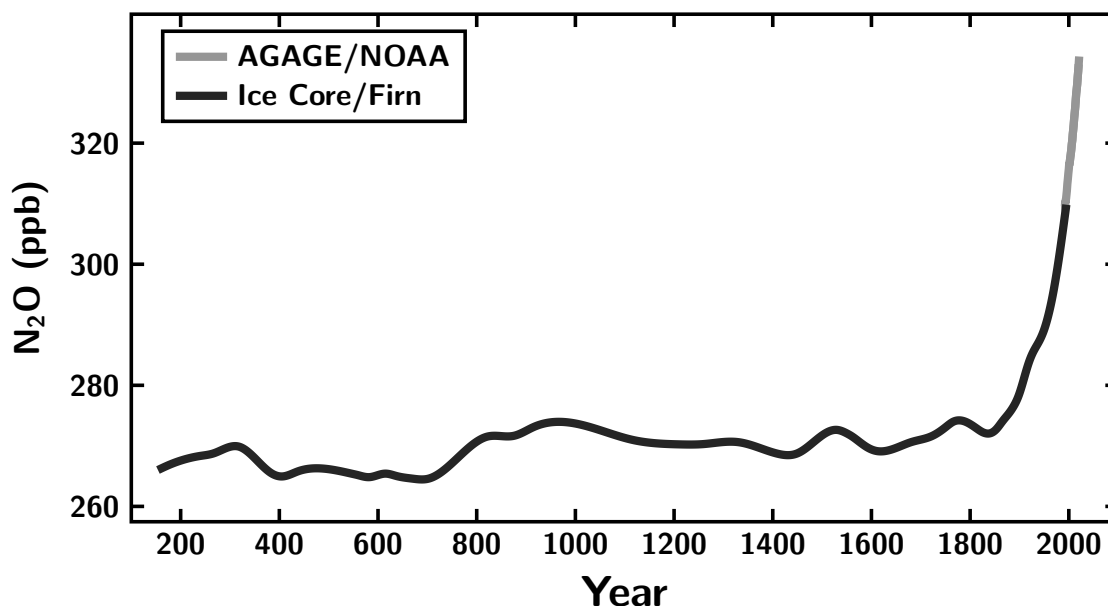


Figure 6: The temporal development of atmospheric concentrations of  $\text{N}_2\text{O}$  over the past 2000 years. (Updated to the year 2021 from Müller, 2021). Historical data (dark line, MacFarling Meure et al., 2006; Park et al., 2012; Rubino et al., 2019) are deduced from ice core and firn measurements; data after 1995 (light grey line) are global mean data from the atmospheric monitoring networks AGAGE (Prinn et al., 2018) and NOAA (Dlugokencky et al., 1994). (See acknowledgements for the data sources).

192  $\text{N}_2\text{O}$  is a natural product of microbiological processes in soils and waters; but human  
 193 activities, such as the application of nitrogen fertilisers in agriculture, cause substantial

194 additional emissions of N<sub>2</sub>O to the atmosphere (Park et al., 2012). The point of the impact  
 195 of human activities on the atmospheric nitrogen budget was emphasised in 1977 in another  
 196 influential paper by Paul and Dieter Ehhalt (reference 11, see also Solomon, 2021b).

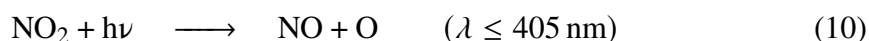
197 We know today that the most substantial source of stratospheric NO<sub>x</sub> (and other  
 198 nitrogen compounds) is N<sub>2</sub>O, which is emitted at the surface of the earth and which is  
 199 essentially inert in the troposphere. Further, human activities have led to a significant  
 200 increase in atmospheric concentrations of N<sub>2</sub>O (Park et al., 2012). Concentrations of N<sub>2</sub>O  
 201 have risen from levels below 280 ppb<sup>2</sup> before 1800 to more than 334 ppb today (Fig. 6).  
 202 In the second half of the 21<sup>st</sup> century, when chlorine containing ozone depleting species  
 203 will be substantially reduced (see section 6.3), the temporal development of N<sub>2</sub>O (together  
 204 with CO<sub>2</sub> and CH<sub>4</sub>) will constitute the dominant driver of changes of stratospheric ozone  
 205 (Ravishankara et al., 2009; Portmann et al., 2012; Solomon, 2021b).

## 206 5 Tropospheric Chemistry

### 207 5.1 Tropospheric ozone

208 The tropospheric ozone budget is influenced both by a downward transport of ozone from  
 209 the stratosphere to the troposphere and by in situ tropospheric reactions. Paul Crutzen,  
 210 in 1978/1979, together with Jack Fishman and Susan Solomon (30), presented evidence  
 211 for large *in situ* photochemical sources and sinks of tropospheric ozone (12, 14). And  
 212 tropospheric ozone is not only important as a key component of both pristine and polluted  
 213 air; it is also a major greenhouse gas (16).

Paul contributed significantly to scientific questions regarding tropospheric ozone (e.g.,  
 12, 14, 16, 24). He turned his attention to tropospheric chemistry, a topic which he had  
 hitherto not focused on (30) following the publication of the important paper by Levy  
 (1971). In the troposphere the chain of reactions



214 is the basic mechanism of chemical ozone formation (here RO<sub>2</sub> = HO<sub>2</sub>, CH<sub>3</sub>O<sub>2</sub>, or another  
 215 organoperoxy radical and M is a collision partner, see reaction 4).

216 Starting with the paper by Levy (1971), the importance of the photolysis of O<sub>3</sub> by solar  
 217 ultraviolet radiation was realised, which splits O<sub>3</sub> into an O atom and an O<sub>2</sub> molecule. The  
 218 O atom is formed in an (electronically) excited state, O(<sup>1</sup>D), and carries enough energy to  
 219 allow it to abstract hydrogen from H<sub>2</sub>O yielding two OH radicals:



220 a process energetically unfavourable for ground-state oxygen atoms (O(<sup>3</sup>P)). O(<sup>1</sup>D) can  
 221 react similarly with other hydrogen-bearing molecules, such as hydrocarbons. Levy  
 222 (1971) proposed that O<sub>3</sub> photolysis followed by reaction (11) should be a source of OH  
 223 radicals in the troposphere, which are responsible for the tropospheric oxidation of CH<sub>4</sub>

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<sup>2</sup>Parts per billion, a measure of the molar mixing ratio.

224 and CO. The different roles of CH<sub>4</sub> and CO in tropospheric chemistry and further details  
225 of Paul's discoveries in this field are described by Fishman et al. (2022). OH radicals in  
226 the troposphere are also responsible for reaction with a variety of other compounds, thus  
227 initiating the oxidation of these species and their removal from the atmosphere.

228 In a model study in 1991, Paul together with Peter H. Zimmermann (24), calculated  
229 the OH concentrations for pre-industrial and present conditions; they reported that tropo-  
230 spheric OH reacts rapidly with the majority of gases that are emitted to the atmosphere  
231 because of natural processes and human activities. Through the reaction with OH, these  
232 gases are removed from the atmosphere. This approach allowed changes in O<sub>3</sub>, CO, and  
233 OH concentrations to be determined that have occurred as a result of human activity (24).  
234 The fact that the highest tropospheric OH concentrations are found in the tropics (30),  
235 points to the importance of the tropics in atmospheric chemistry (see also section 5.3).

## 236 5.2 Tropospheric pollution

237 Paul also contributed to research on the impact of HNO<sub>3</sub> and H<sub>2</sub>SO<sub>4</sub> to acidic precipita-  
238 tion (18) and initiated major measurement campaigns on tropospheric pollution in Asia  
239 and in the Mediterranean (39, 43). The Indian Ocean Experiment found surprisingly  
240 high pollution levels over the entire Northern Indian ocean region during the dry season  
241 January–March 1999. In this time period, South and Southeast Asian emissions of pol-  
242 lutants resulted in poor air quality over a huge area (39). The Mediterranean Intensive  
243 Oxidant Study (MINOS), in the summer of 2001, uncovered air pollution layers in the  
244 Mediterranean from the surface to an altitude of 15 kilometres. Pollution was found to  
245 be transported to the Mediterranean from European, Asian and (to a lesser extent) North  
246 American sources (43).

## 247 5.3 Biomass-burning

248 First estimates by Paul and a number of colleagues in 1979 (15) showed a major impact  
249 of a number of trace gases originating from biomass-burning (e.g., CO, H<sub>2</sub>, NO, CH<sub>3</sub>Cl  
250 and COS) on the atmosphere. These estimates were based on measurements at the NCAR  
251 laboratories in Boulder (15, 30) and demonstrated that emissions of these gases from  
252 biomass-burning could constitute a significant fraction of their total global emissions (15).  
253 A year earlier, Pat Zimmerman and colleagues (including Paul) had already pointed out  
254 that the release of hydrocarbons (isoprene and terpenes) from biomass burning could  
255 contribute in an important way to the atmospheric sources of CO and H<sub>2</sub> (13).

256 These first results stimulated a considerable international research effort including  
257 several measurement campaigns in the tropics (e.g. 39). It was found that large quan-  
258 tities of ozone are formed as a result of biomass-burning emissions (22, 30) and that  
259 biomass-burning emissions in South and Southeast Asia strongly enhance tropospheric  
260 CO concentrations (39). Moreover, biomass-burning is an important source of (different  
261 types of) aerosol particles. Together with Andi Andreae, Paul showed that such particles  
262 may serve as cloud condensation nuclei and thus substantially influence cloud microphys-  
263 ical and optical properties (22). This has consequences for the radiation budget and the  
264 hydrological cycle of the atmosphere (22).

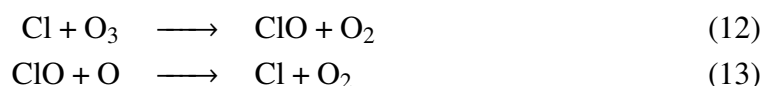
265 Today, the importance of biomass-burning for the atmosphere is well established.

266 Biomass-burning is widespread, especially in the tropics; it is a source of a great number  
 267 of photochemically and radiatively active trace gases, such as NO<sub>x</sub>, CO, CH<sub>4</sub>, reactive  
 268 hydrocarbons, H<sub>2</sub>, N<sub>2</sub>O, OCS, CH<sub>3</sub>Cl, and HCN. Biomass-burning is not restricted to  
 269 tropical deforestation, but is also a common activity related to agriculture (particularly in  
 270 developing countries) and fuel wood combustion; it may also happen through smouldering  
 271 fires in peatlands.

## 272 **6 Anthropogenic chlorine in the stratosphere and the** 273 **Antarctic ozone hole**

### 274 **6.1 Global ozone loss**

275 In 1973, Paul Crutzen received the preprint of a paper entitled “Stratospheric sink for  
 276 chlorofluoromethanes: chlorine atom-catalysed destruction of ozone” by M. J. Molina  
 277 and F. S. Rowland, neither of whom he had heard of before (30, 53). Their paper (Molina  
 278 and Rowland, 1974) pointed out that the chain of reactions



279 could destroy ozone in the stratosphere. The stratospheric chlorine, in turn, would originate  
 280 from the photolysis of industrially produced chlorofluorocarbons (CFCs) at stratospheric  
 281 altitudes. Further, their paper made the point that the reduction in stratospheric ozone  
 282 would result in an increase in the amount of UV reaching the earth’s surface, with potential  
 283 severe environmental consequences (Harris, 2020).

284 Paul immediately understood the importance of this paper and about two months after  
 285 the publication by Molina and Rowland (1974) presented a model analysis (7) on the  
 286 potential ozone depletion resulting from continued use of CFCs; the paper indicated the  
 287 possibility of up to ~ 40% ozone depletion near 40 km altitude as a result of continued  
 288 use of CFCs at 1974 rates.

289 In the following years the paper by Molina and Rowland (1974) stimulated the making  
 290 of a great deal of new measurements in the laboratory and in the atmosphere as well as  
 291 major advances in atmospheric modelling (Harris, 2020). Meanwhile, in 1974, Paul had  
 292 moved with his family to Boulder, Colorado, where he assumed two part-time positions;  
 293 one as a consultant at the Aeronomy Laboratory of the National Oceanic and Atmospheric  
 294 Administration (NOAA) and one at the Upper Atmosphere Project of the National Center  
 295 for Atmospheric Research (NCAR). Later, in 1977, he took up the directorship of the Air  
 296 Quality Division of NCAR. In 1980 he moved back to Europe to take up a directorship at  
 297 the “Max-Planck-Institut (MPI) für Chemie” in Mainz as the successor of Christian Junge.  
 298 Figure 7 shows him at his desk in his office at the “MPI für Chemie”. In Mainz you could  
 299 see him walk to work with his little rucksack in the morning; he did not own a car.

### 300 **6.2 Polar ozone loss**

301 In the early 1980s, Paul was heavily involved in various studies on the potential envi-  
 302 ronmental impacts of a major nuclear war (see section 7). But in 1985 came another

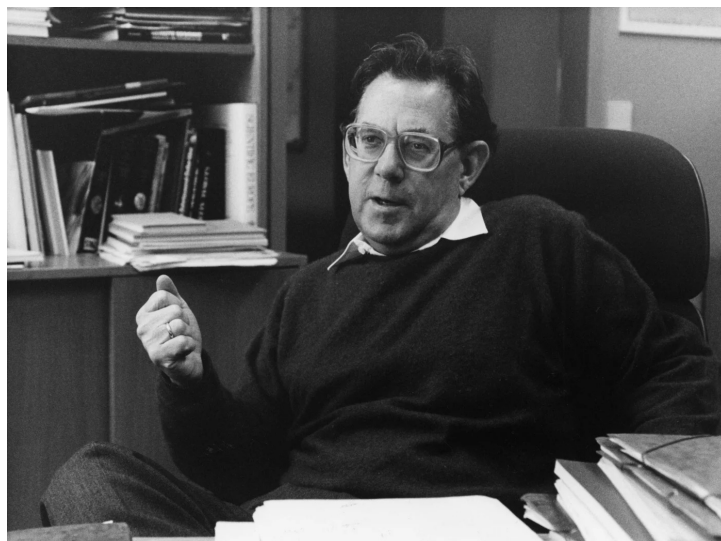


Figure 7: Paul Crutzen in his office at the “Max-Planck-Institut (MPI) für Chemie” in Mainz in 1994. (From the archives of the “Max-Planck-Gesellschaft”, Berlin-Dahlem).

major (and surprising) discovery – J. Farman, B. Gardiner, and J. Shanklin of the British Antarctic Survey reported that the spring values of total ozone in Antarctica since the late 1970s had fallen considerably (Farman et al., 1985); this phenomenon soon became known as the “ozone hole” (Stolarski et al., 1986; Newman et al., 2022). Paul attended a workshop on the subject in Boulder in early 1986, where he learned about the theories put forward to explain the phenomenon, most notably the idea by Solomon et al. (1986) of chlorine activation on the surface of stratospheric ice particles through the reaction



Already on the flight back to Germany and together with F. Arnold after his return, he developed the idea that in the cold Antarctic stratosphere, heteromolecular condensation of  $\text{H}_2\text{O}$  and  $\text{HNO}_3$  occurs (well above the threshold temperature for the formation of pure ice particles), which leads to the formation of  $\text{HNO}_3\text{—H}_2\text{O}$  aerosol particles in the stratosphere and, consequently, to a strong depletion of  $\text{HNO}_3$  (and thus  $\text{NO}_x$ ) in the gas-phase. (The idea of heteromolecular condensation of  $\text{H}_2\text{O}$  and  $\text{HNO}_3$ , as well as the removal of  $\text{NO}_x$  from the gas-phase had also been independently developed by Toon et al. (1986).) The condensation of  $\text{HNO}_3$  further leads to an increase in aerosol surface area and thus to an accelerated rate of heterogeneous reactions like reaction (14). The paper on this mechanism was published at the end of 1986 (21).

The ideas of Crutzen and Arnold (21) were an important contribution to the explanation of the ozone hole (albeit not the only one, see reference 30). Today, much more is known, for example that further heterogeneous reactions are important (like  $\text{HCl} + \text{HOCl} \longrightarrow \text{Cl}_2 + \text{H}_2\text{O}$ , reference 25). It has further been established that supercooled  $\text{H}_2\text{SO}_4/\text{HNO}_3/\text{H}_2\text{O}$  droplets exist in the stratosphere with important consequences for the microphysics of polar stratospheric clouds and heterogeneous reactions (see e.g. references 28, 33 and Hanson et al., 1994; Solomon, 1999; Shi et al., 2001).

Even in the absence of low temperatures,  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  particles always exist in the lower stratosphere (Junge et al., 1961); these particles are formed from stratospheric  $\text{SO}_2$ . Direct

injections by large volcanic eruptions are one (intermittent) source of stratospheric SO<sub>2</sub>. The other source is oxidation of OCS, which is produced at the earth's surface but reaches the stratosphere; this mechanism was also proposed by Paul (10) and soon confirmed by measurements that showed atmospheric concentrations of OCS of approximately 500 ppt<sup>3</sup> (Sandalls and Penkett, 1977).

### 6.3 The Montreal Protocol

A large number of scientists worked on the impact of chlorine released to the atmosphere as a result of human activities on stratospheric ozone and in particular on the Antarctic ozone hole. This research led eventually to the Montreal Protocol in 1987 and its adjustments and amendments in the following decades (WMO, 2018). Paul made substantial contributions to this research (e.g., references 7, 21, 25, 28, 33). The Montreal protocol is one of the most important environmental treaties agreed on so far and is today ratified by all member states of the United Nations; it is expected to lead to the eventual phase-out of all ozone depleting substances worldwide and it led to the slow decline of the atmospheric halogen loading that is observed today (WMO, 2018).

In a model study Newman et al. (2009) investigated what might have happened had the Montreal Protocol never been signed. In this case, a continued increase in ozone depleting substances in the atmosphere would have occurred at a rate of 3% per year. In such a “world avoided” scenario, large ozone depletions in the polar region become year-round phenomena (rather than just seasonal phenomena as is currently observed) and tropical lower stratospheric ozone collapses to near zero by about 2060. In response to such ozone changes, ultraviolet radiation at the ground also increases substantially (Newman et al., 2009). Paul's research forms part of the backbone of the science by an active group of researchers that fostered policy action such as the Montreal protocol (Solomon, 2021b); it therefore helped avoiding the massive ozone losses described by Newman et al. (2009) in their “world avoided” scenario. Instead today the emergence of a healing of the Antarctic ozone layer is observed (e.g., Solomon et al., 2016; WMO, 2018).

And Paul emphasised that things could have been much worse; he wrote in 1996: “Noting that nobody had worried about the atmospheric consequences of the release of Cl or Br before 1974, I can only conclude that we have been extremely lucky, which shows that we should always be on our guard for the potential consequences of the release of new products into the environment” (from reference 30).

## 7 “Nuclear winter”

The involvement of Paul Crutzen in “nuclear winter” studies may be explained (30) by his research interests both into the effects of NO<sub>x</sub> on stratospheric ozone (section 3) and in biomass-burning (section 5.3). His work on “nuclear winter” started in 1981, when he was asked by the editor of *Ambio* to contribute to a special issue on the environmental consequences of a major nuclear war. His initial thought was that he would make an update on model predictions of the destruction of ozone by the NO<sub>x</sub> that would be produced and carried up into the stratosphere by large nuclear explosions (an issue already discussed

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<sup>3</sup>Parts per trillion, a measure of the molar mixing ratio.

much earlier, 8). In the late seventies and early eighties, atmospheric models had evolved into two-dimensional zonally averaged models, which covered the atmosphere from pole to pole and where transport was represented by a combination of advection terms and a two-dimensional matrix of eddy diffusion coefficients – the zonally averaged transport of such models approximated a diabatic circulation (e.g., Brasseur, 2019).

John Birks (who spent a sabbatical at MPI Mainz at that time) joined him in this study (see also Fishman et al., 2022). In their first set of calculations, to their surprise, no stratospheric ozone depletion resulted from the nuclear explosions. The reason was that the first scenario they used implied that all explosions were less than one megaton (in equivalent weight of conventional explosive), which, as they found, do not reach the stratosphere. When the scenario was widened, to also include explosions of a few megatons each, substantial stratospheric ozone depletion was indeed predicted.

However, Paul and John Birks soon realised that the direct impacts of possible nuclear explosions were much stronger (19). Such possible nuclear explosions should cause large amounts of smoke emissions (as well as  $\text{NO}_x$  and hydrocarbon emissions) from fires, which would reach the middle and higher troposphere. The resulting absorption of sunlight by the black smoke could lead to darkness and strong cooling at the earth's surface. Further, the black smoke leads to a heating of the atmosphere at higher elevations, thus creating rather atypical meteorological and climatic conditions that have substantial environmental consequences (references 19, 20).

The term “nuclear winter” to describe the phenomenon was coined by the TTAPS (Turco, Toon, Ackerman, Pollack and Sagan) group in their subsequent paper that expanded on the climate effects of the injections of large amounts of soot and dust into the atmosphere as the result of a nuclear war (Turco et al., 1983). Numerous subsequent studies have confirmed that a large-scale nuclear war would cause severe climate effects, including a study by the U.S. National Academy of Sciences (1985); severe climate effects would also be caused by regional nuclear conflicts (Mills et al., 2008).

Although Paul did not count the nuclear winter idea among his greatest scientific achievements, he was convinced “that, from a political point of view, it is by far the most important” (30). He stated that this idea “highlights the dangers of a nuclear war and convinces me that in the long run mankind can only escape such horrific consequences if nuclear weapons are totally abolished by international agreement” (from reference 30).

## 8 Atmospheric chemistry and climate change

Atmospheric chemistry and climate change are linked;  $\text{CH}_4$  has an important chemical sink in the troposphere (reaction with OH, section 5.1) and the main atmospheric sink of  $\text{N}_2\text{O}$  is photolysis in the stratosphere (section 4.3). Likewise, ozone concentrations throughout the atmosphere are determined by atmospheric chemistry and are influenced by human activities. Finally, atmospheric chemistry affects the nature and the concentrations of cloud-condensation nuclei and thus their impact on cloud properties.

These facts led Paul Crutzen to conduct studies on reactions taking place in and on cloud droplets and tropospheric aerosol particles, which result in the removal of  $\text{NO}_x$  and reduce OH and  $\text{O}_3$  in air masses influenced by clouds (23, 26). Another topic of the work of Paul was the impact of (natural) halogens on the formation of cloud-condensation nuclei and on ozone in the marine boundary layer (31).



413 The most important atmospheric greenhouse gases are CO<sub>2</sub>, CH<sub>4</sub>, and N<sub>2</sub>O, the  
414 effective radiative forcing in 2019 (relative to the year 1750) for CO<sub>2</sub> is 2.16 W m<sup>-2</sup>, for  
415 CH<sub>4</sub> is 0.54 W m<sup>-2</sup> and for N<sub>2</sub>O is 0.21 W m<sup>-2</sup> (Arias et al., 2021). Further, tropospheric  
416 ozone is a significant greenhouse gas (16), which interacts with terrestrial infrared radiation  
417 in the 10 μm region (an atmospheric window region); the effective radiative forcing in  
418 2019 (relative to the year 1750) of ozone (where the forcing is dominated by tropospheric  
419 ozone) is 0.47 [0.24 to 0.71] W m<sup>-2</sup> (Arias et al., 2021).

420 Of greatest importance to radiative forcing are any changes in ozone in the tropopause  
421 region and in the tropics and subtropics (Riese et al., 2012). The latter point underlines the  
422 importance of observations of tropospheric ozone in the tropics and subtropics, where the  
423 atmospheric oxidation efficiency is highest, where high ozone concentrations are observed  
424 over the polluted regions of the continents and where observations are particularly sparse  
425 (30).

426 Moreover, emissions of CO<sub>2</sub> and N<sub>2</sub>O caused by human activities are not independent.  
427 The conversion of biomass to biofuel will allow some emissions of CO<sub>2</sub> (caused by the  
428 use of fossil fuel) to be saved, but this biofuel production will lead to some additional  
429 N<sub>2</sub>O emissions caused by the use of fertilisers (see also section 4.3). Paul conducted an  
430 analysis of these compensating effects (49) finding that the production of commonly used  
431 biofuels may contribute as much or more to global warming through these additional N<sub>2</sub>O  
432 emissions than the cooling by the fossil fuel savings.

433 A lot of Paul's work on the link between atmospheric chemistry and climate change  
434 was done in collaboration with his friend Veerabhadran Ramanathan (e.g., 16, 39, 40),  
435 who works at Scripps Institution of Oceanography, University of California San Diego in  
436 La Jolla, USA. Paul was a regular visitor at Scripps and between 1992 and 2008 served as  
437 distinguished professor (part-time) at Scripps in La Jolla.

## 438 9 Geoengineering

439 In 2006 Paul Crutzen wrote an essay (47) entitled “Albedo enhancements by stratospheric  
440 sulfur injections: a contribution to resolve a policy dilemma?” In this essay he stated  
441 that “the usefulness of artificially enhancing earth's albedo and thereby cooling climate by  
442 adding sunlight reflecting aerosol in the stratosphere [...] might again be explored and  
443 debated . . .” (47). This topic was not entirely new (e.g., Budyko, 1977; Dickinson, 1996)  
444 and even Paul himself had used the term “geoengineering” earlier (in 2002, reference  
445 41). Further, this was not a single essay, but rather a special section in *Climatic Change*  
446 consisting of six papers in total (56). The essay by Paul had a profound influence; it  
447 initiated a large number of studies on geoengineering with many different facets.

448 Today it is accepted that there are two fundamentally different kinds of geoengineering:  
449 (a) actively removing CO<sub>2</sub> from the atmosphere, and (b) cooling the earth by either reflect-  
450 ing more sunlight into space (the original proposal in reference 47) or by modifying cloud  
451 properties. The number of scientific publications on all aspects of climate engineering  
452 increased very strongly after the publication of the special section of *Climatic Change* in  
453 2006 (56). In an interview in 2013, Paul supported continuing research on geoengineering,  
454 but doubted that sunlight-reflecting aerosol would ever be injected into the stratosphere  
455 (Möllers et al., 2015). His position (like on other occasions) was that predictions on the  
456 effect of geoengineering and the possible environmental consequences should be reliable,

457 before application can be considered.

458 Soon after the publication of Paul’s essay, there was an assessment by the Royal  
459 Society (Shepherd, 2009), and many other international assessment reports and research  
460 initiatives followed (reference 56 and references therein); the most recent assessment is by  
461 the National Academies of Sciences, Engineering, and Medicine (2021). Geoengineering  
462 has also been addressed in the frame of the IPCC, and one should particularly emphasise  
463 the geoengineering model intercomparison project (GeoMIP, Kravitz et al., 2011).

464 Clearly, geoengineering would have a lot of unintended side-effects, one important  
465 example is that an intentional injection of sulphur into the stratosphere might enhance  
466 stratospheric ozone depletion in general and lead to a slower recovery of the Antarctic  
467 ozone hole in particular (e.g. Tilmes et al., 2021). Indeed, the most recent (2022) “WMO  
468 Scientific assessment of ozone depletion” contains for the first time a chapter on the  
469 question of the impact of stratospheric sulphur injections on the stratospheric ozone layer.

470 The final two sentences of Paul’s essay in 2006 should not be forgotten: “. . . I repeat:  
471 the very best would be if emissions of the greenhouse gases could be reduced so much  
472 that the stratospheric sulfur release experiment would not need to take place. Currently,  
473 this looks like a pious wish” (47).

## 474 **10 The Anthropocene**

475 Paul Crutzen coined the term “Anthropocene”; it was a spontaneous thought at a sym-  
476 posium of the International Geosphere-Biosphere Program (IGBP) in February 2000 at  
477 Cuernavaca, Mexico. He recalls the event: “the chairman spoke of the holocene period  
478 the whole time. Suddenly it occurred to me that we were no longer in the holocene. In  
479 the holocene period human influence was limited. So I then said: ‘We are no longer in  
480 the holocene, we are in the anthropocene’. [. . .] during the coffee break it was the topic  
481 of conversation” (from reference 54).

482 But he soon discovered that E. F. Stoermer (whom he never met) had come up with the  
483 term independently and had already used it informally in the 1980s (Möller et al., 2015;  
484 Trischler, 2016). Paul and E. F. Stoermer introduced the concept of the Anthropocene  
485 together in a newsletter of the IGBP in the year 2000 (37). Putting forward the hypothesis  
486 of a new geological epoch, the Anthropocene, is based on the fact that humanity has  
487 become a powerful geological force and will continue to be such a force for thousands of  
488 years (41). Paul also suggested that the “Anthropocene could be said to have started in the  
489 late eighteenth century, when analyses of air trapped in polar ice showed the beginning of  
490 growing global concentrations of carbon dioxide and methane” (reference 41).

491 The substantial increase of carbon dioxide in the atmosphere, which is driven by the  
492 burning of fossil fuel by humans and which started with the industrial revolution is the  
493 most notable example of the impact of humanity on nature. But it is clearly not the only  
494 one. During the past three centuries, human population has increased tenfold, which has  
495 resulted in a huge growth in cattle population (and in an concomitant rise in atmospheric  
496 methane concentrations – the increase is more than 100% today compared to preindustrial  
497 concentrations). Fossil fuel burning also causes a release of SO<sub>2</sub> to the atmosphere that is  
498 now exceeding the sum of all natural emissions. More nitrogen is now fixed synthetically  
499 and applied as fertilisers in agriculture than fixed naturally in all terrestrial ecosystems (see  
500 also section 4.3 and reference 11); this has led to an increase of about 20% in atmospheric

N<sub>2</sub>O (Fig. 6). Furthermore, humanity has released a type of gas (chlorofluorocarbons) to the atmosphere that has no natural analogues. The release of chlorofluorocarbons to the atmosphere would have caused the destruction of the ozone layer as we know it had no international regulatory measures been taken (see sections 6.2 and 6.3).

These are important examples, but there are many other major and still growing impacts of humanity on the earth system (37, 41, 48, 50, 57, 58). In Paul's own words: "In the Anthropocene, climate change is tightly linked to many of the other grand challenges that we face – including the improvement of air quality, public health, sustainable development, and human well-being" (Al-Delaimy et al., 2020). This underlines the urgency of achieving an effective planetary stewardship (50) in spite of uncertainties in our understanding of the earth system. "In the Anthropocene some risks may be overestimated, while others are underestimated by the best available science at any given time, and human society is constantly forced to make decisions while recognizing that science is incomplete" (Solomon, 2021b).

## 11 Leadership

Paul Crutzen was always a scientist and never became purely a manager or an administrator. But he was nonetheless responsible for the setting up of many programmes to study and understand chemical phenomena in the atmosphere. He was also available to serve as an advisor of environmental politics; for example he was an advisor of the Enquete commission of the German parliament "zum Schutz der Erdatmosphäre"<sup>4</sup> from 1987 to 1990. And later, as a member of the Pontifical Academy of Sciences, he contributed substantially to formulating Pope Francis' encyclical on climate change (Fishman et al., 2022).

Of particular significance regarding the programmes initiated by Paul was the International Global Atmospheric Chemistry (IGAC) programme of the IGBP (International Geosphere-Biosphere Program) which was set up in 1988 at Dookie college, a small college in Australia, in northern Victoria (see Fig. 8). Following the Dookie meeting, Paul served as the first Vice Chair of the IGAC steering committee from 1988 to 1995 (Chair, R. Prinn) playing a leading role in the establishment of biospheric processes as an essential component of IGAC (Prinn, 1994).

Paul was also instrumental in raising the profile of atmospheric chemistry within earth system science, showing its relevance to the stability of life on earth. This became particularly important following the discovery of the ozone hole and the realisation that large-scale ozone loss in the stratosphere was a definite possibility (see section 6.2).

Although Paul was a theoretician, he always emphasised the importance of observations (in particular atmospheric observations), he once said "observations always come first". There were major observational platforms set up under the influence of Paul; one is the CARIBIC experiment (which is still in operation) where a commercial airliner is used to transport a container equipped with a lot of sensors for measuring individual molecules (42). A large number of molecules including hydrocarbons, halocarbons, nitrogen oxides, CO, ozone, water vapour, etc. is measured on a global scale in the upper troposphere/lower stratosphere in order to improve the understanding of chemical processes and chemical

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<sup>4</sup>In English: "for the protection of earth's atmosphere".



Figure 8: The founding fathers of the International Global Atmospheric Chemistry (IGAC) programme at Dookie college in 1988. From left to right: E. Sanhueza, D. Albritton, R. Duce, I. Galbally, P. Crutzen, D. Ehhalt, P. Buat-Menard, H. Rodhe. (Picture courtesy of Henning Rodhe.)

543 composition worldwide. He also initiated trace gas measurements on the Trans-Siberian  
544 Railroad covering very remote locations (35).

545 Paul played a key role in the development of the global observing system from space. In  
546 the early 1980s he created a group in his department to exploit optical sensors for the mea-  
547 surement of trace gases. Building on the development of Differential Optical Absorption  
548 Spectroscopy (DOAS) this group led initially by the late Dieter Perner and John Burrows  
549 began in 1984 to develop concepts for passive remote sensing of the atmosphere using solar  
550 radiation. In July 1988 this led to the proposal of the SCIAMACHY (SCanning Imaging  
551 Absorption spectroMeter for Atmospheric CHartographY), which was subsequently se-  
552 lected for the ESA (European Space Agency) Envisat mission. SCIAMACHY had a wide  
553 set of sensors and revealed sources of many important molecules, like methane, sulphur  
554 dioxide (SO<sub>2</sub>), and nitrogen dioxide (NO<sub>2</sub>) (e.g., reference 46). It was typical of Paul  
555 that he would contribute to and support such an ambitious pioneering earth observation  
556 research mission.

557 A smaller version of SCIAMACHY, called SCIA-mini, of which Paul was co-Principal  
558 Investigator with John Burrows, was proposed for a launch opportunity in late 1988. After  
559 descopeing to being nadir viewing only, the concept was was selected for launch in 1990  
560 and renamed GOME (Global Ozone Measurement Experiment). A continuing series of  
561 instruments resulted from this initiative. GOME, which as the name suggests, had a focus  
562 on ozone also delivered measurements of key trace gases; the first global observations of  
563 nitrogen dioxide, sulphur dioxide and formaldehyde (HCHO) (34 and references therein).  
564 The objective of the GOME instruments was and is to observe the global recovery of O<sub>3</sub>  
565 as a result of the measures, enacted by the Montreal protocol (see section 6.3). Paul was  
566 also an influential member of the science team of the Halogen Occultation Experiment  
567 (HALOE) onboard the Upper Atmosphere Research Satellite (UARS), which measured the

vertical distributions of ozone and key trace gases that affect the global ozone distribution (references 27, 29, 36, 38).

One area where Paul had a big impact, but perhaps not so well recognised, was being a founding member of the European Research Council (ERC). He was also deeply involved in the setting up of the Academia Europea. Academia Europea had the intention of highlighting scholastic activities, while the ERC had the intention of funding various research programmes. Paul was instrumental in the launch of the ERC and a highly esteemed member from 2005 to 2008. Paul was always determined to support scientific activities in Europe with a particular focus on cooperation of scientists from different European countries.

Paul was also a key figure in the pioneering work of establishing the new concept of interactive open access publishing; together with Uli Pöschl and Arne Richter he helped to create the journal “Atmospheric Chemistry and Physics (ACP)”, which features a public peer-review and an interactive public discussion of submitted manuscripts (51).

In an e-mail on 18 September 2000, A. Richter said that a “meeting of the ‘younger and wilder’ atmospheric scientists under the lead of Ulrich Pöschl and Paul Crutzen regarding the launch of a new EGS [now EGU] journal on atmospheric chemistry took place on 15 September in Mainz”; this meeting was the birthplace of the new journal ACP (51). The journal ACP was founded in 2000 with Paul as a member of the advisory board. Today, ACP is a very well established scientific journal. In Paul’s words: “It has been an amazing journey: over a short period of merely a decade, a novel idea originating from Uli Pöschl and developed by an enthusiastic group of hundreds of scientists, created a new way of scientific publishing and communication, initially covering the fields of atmospheric chemistry and physics. The example has since been followed by many successors in other disciplines, with more to come.” (51).

Paul did not shy away from controversies in scientific or even public debate. He was involved in a number of issues that were intensively discussed among scientists and in the public, in particular: supersonic transport (section 4.1), “nuclear winter” (section 7), the ozone hole and the Montreal Protocol (sections 6.2 and 6.3), geoengineering (section 9), and the concept of the Anthropocene (section 10). In all these cases, there was public dispute and intellectual fights between many scientists, but Paul was always on the side of those who eventually prevailed.

## 12 Overall assessment

### The person

Paul Crutzen was a warmhearted, friendly and humorous person; he was always open to everyone. At a conference, or at similar occasions one would have very enjoyable and pleasant evenings with him — and if you were lucky, Terttu would be present as well. He was concerned about the welfare of his fellow human beings; he once said “private things are always important”. As busy Paul was with his scientific work, he always had time for his family. All of the Crutzen family have wonderful memories of family gatherings, weekends, holidays, and vacations.

Paul was a very cosmopolitan person, he spoke many languages in addition to his native Dutch. His wife Terttu is from Finland and Swedish became the family language

611 which was spoken at their home. On the occasion of a visit to Jülich (on 15 February 1996)  
612 Paul had a conversation with Martin Böhmermann from the Netherlands, who worked in  
613 Jülich as a glassblower; of course in Dutch. Böhmermann immediately realised that Paul  
614 was from Amsterdam – and not from the richest part of town. Of course Paul’s command  
615 of English was very good. And the same was true for German. The conversations at the  
616 MPI in Mainz were in English, Dutch, German and sometimes in Swedish (depending on  
617 the person he talked to). He also spoke French and learned a bit of Finnish.

618 Paul was travelled a lot and his office was often empty. But he was always accessible  
619 – there was always a fax number, where he could be reached. And one could count on a  
620 rapid reaction, typically within a day or two, with well thought through comments on any  
621 scientific request (be it a paper draft or the chapter of a thesis).

622 When Paul’s mind was set on science, he might forget everything else. It would  
623 happen that he asks students or postdocs to come to his home on a Saturday afternoon;  
624 after finding them at MPI. At his home there were intense discussions, checking of numbers  
625 and examining model results. Paul might even forget to offer tea or coffee – of course  
626 Terttu would not. He might also call you in the middle of the night (from the US, the  
627 fax machine at MPI was out of paper) asking you to collect several fax messages from  
628 the fax at the Meteorology department at the University of Mainz and distribute them to  
629 colleagues at MPI. First thing in the morning; these messages were important.

630 In his scientific work, Paul was very careful with a lot of attention to details. But at  
631 the same time he could be rather impatient. He always wanted to see the new results, the  
632 new plots or the new simulations immediately. He worked very hard and in a dedicated  
633 way; he said once (after discussing extensively the reviewer’s comments on the paper of  
634 reference 29) “. . . it is hard work to write a paper – really hard work. . . Some people claim  
635 that they can write a paper in a day – I could not do that”.

636 Paul Crutzen was a very unpretentious person. Susan Solomon remembers a story  
637 when as a graduate student, she visited the MPI in Mainz around 1980, and she was  
638 at a meeting with a dozen of other scientists where she addressed him as Paul and also  
639 disagreed slightly with some technical point he had made. Susan could literally hear the  
640 sharp intake of breath from others because she had disagreed. He was oblivious to titles,  
641 just smiled and said “ja, you’re right”. The opposite happened as well; lots of discussions  
642 among scientists on a particular issue resulting finally in the conclusion “Paul is right  
643 here” – where people had not thought of Paul’s point earlier in the discussion.

## 644 **The scientist**

645 Paul Crutzen was a brilliant scientist – what is particularly remarkable is the breadth of  
646 his work. Many great scientists have achieved great advances, but what is particularly  
647 impressive about Paul is the number of different topics in atmospheric science to which  
648 he made fundamental contributions – these contributions are described above and shall  
649 not all be repeated here.

650 His research interests were broad, encompassing topics in the mesosphere, the strato-  
651 sphere and the troposphere – with a particular emphasis on the issue of climate change  
652 and air quality (12, 14, 16, 32, 39, 40). Further, he initiated the discussion on the question  
653 of geoengineering (47) and coined the term “Anthropocene”, which he suggested as the  
654 geological epoch starting with the industrial revolution. Stating that humanity since that

time had had a vast impact on the planet (37, 41).

Paul's work not only had a profound influence on the scientific world, but his work also had an influence on many aspects of environmental politics throughout many countries – to name but a few, the issue of supersonic transport in the stratosphere (section 4.1), the Montreal protocol, (section 6.3), the concept of “nuclear winter” (section 7), and many aspects of air pollution caused by human activities.

Paul's scientific works, especially on the impact of human activities on the atmosphere and climate are influential today and they will continue their influence into the future. They will show the way forward for future work on many scientific questions and likewise for the necessary future policy action to halt the warming of climate.

## Acknowledgements

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Stuart Penkett and Henning Rodhe were both mentioned in Paul Crutzen's papers at the Royal Society, they read the draft of the memoir several times and provided very useful suggestions and help. The text in section 11 goes back to a draft by Stuart Penkett. Further shorter pieces were contributed by John P. Burrows, Peter S. Liss and Richard P. Wayne. I also thank Richard P. Wayne for a thorough review of the entire memoir.

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There is a lot of biographical material on Paul Crutzen available (Möllers et al., 2015; Lelieveld, 2021; Moortgat et al., 2021; Rodhe, 2021; Solomon, 2021a; Zalasiewicz et al., 2021; Zetzsch, 2021; Fishman et al., 2022, see also references 30, 45, 55); these publications were consulted for the memoir and are recommended for further reading<sup>5</sup>.

Data shown in Fig. 6 are courtesy of E. Dlugokencky, D. M. Etheridge, R. Prinn, M. Rubino, R. Wang, and R. Weiss; Law Dome Ice Core data for N<sub>2</sub>O may be obtained here <https://data.csiro.au/collections/collection/CIcsiro:37077v1> and data on the recent N<sub>2</sub>O atmospheric mixing ratios come from station networks (NOAA/GML: [www.esrl.noaa.gov/gmd/ccgg/trends\\_n2o/](http://www.esrl.noaa.gov/gmd/ccgg/trends_n2o/) and AGAGE: <https://agage.mit.edu/data/agage-data>).

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<sup>5</sup>See also the information on the web: <https://www.mpic.de/3864489/paul-crutzen>

## AWARDS AND RECOGNITION

Paul J. Crutzen, Mario J. Molina, and F. Sherwood Rowland were awarded the Nobel Prize in Chemistry in 1995 for their ground-breaking work “in atmospheric chemistry, particularly concerning the formation and decomposition of ozone”; further major awards and forms of recognition of Paul J. Crutzen are given below.

Rolex-Discover Scientist of the Year (1984)

Leó Szilárd Award for “Physics in the Publics Interest” of the American Physical Society (1985)

Fellow of the American Geophysical Union (1986)

Tyler Prize for Environmental Achievement (1989)

Corresponding Member of the Royal Netherlands Academy of Arts and Sciences (1990)

Volvo Environment Prize (1991)

Member of the Royal Swedish Academy of Sciences (1992)

Member of the Royal Swedish Academy of Engineering Sciences (Kungl. Ingenjörsvetenskapsakademien IVA) (1992)

Member of the Leopoldina (1992)

Group Achievement Award for outstanding contributions to the success of the HALOE/UARS satellite experiment by NASA Langley Research Center (1994)

German prize for the environment (Bundestiftung Umwelt) (1994)

Member of the National Academy of Sciences (United States of America) (1994)

Global Ozone Award for “Outstanding Contribution for the Protection of the Ozone Layer” by the United Nations Environment Programme (1995)

Member of the “Pontificia Academia Scientiarum” (Pontifical Academy of Sciences) (1996)

“Commandeur in de Orde van de Nederlandse Leeuw” (1996)

Recipient of the “Bundesverdienstkreuz mit Stern” (1996)

Member of the “Accademia Nazionale dei Lincei” (Italy) (1997)

Foreign Member of the Russian Academy of Sciences (1999)

Golden Medal by the Academy of Athens (2003)

Recipient of the UNEP/WMO Vienna Convention Award (2005)

Recipient of the Jawaharlal Nehru Birth Centenary Medal 2006, Indian National Science Academy (2006)

Foreign Member of the Royal Society (ForMemRS) (2006)

International Member of the American Philosophical Society (2007)

Recipient of the “Verdienstorden des Landes Rheinland-Pfalz” Mainz, Germany (2013)



733 Honorary Member of the “Nationale Akademie der Wissenschaften Leopoldina”  
734 (2014)  
735 Honorary Member of the Royal Netherlands Chemical Society (2017)  
736 Lomonosov Gold Medal (2019)

## Author Profile

**Rolf Müller** worked with Paul Crutzen at the “Max-Planck-Institut für Chemie” in Mainz from 1988 to 1995. He gained his PhD under the supervision of Karin Labitzke and Paul Crutzen in 1994 from the Free University of Berlin. He works since 1995 at the Institute of Energy and Climate Research of the Forschungszentrum Jülich. He is also a lecturer at the University of Wuppertal.

*No colour photographs of Rolf Müller yet – could be added*

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