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Abstract

Trace gas and aerosol data collected in the tropical tropopause layer (TTL) between 12–18.5 km by the M55 Geophysica aircraft as part of the SCOUT-AMMA campaign over West Africa during the summer monsoon in August 2006 have been analysed in terms of their air mass origins. Analysis of domain filling back trajectories arriving over West Africa, and in the specific region of the flights, showed that the M55 flights were generally representative of air masses arriving over West Africa during the first 2 weeks of August, 2006. Air originating from the mid-latitude lower stratosphere was under-sampled (in the mid-upper TTL) whilst air masses uplifted from central Africa (into the lower TTL) were over-sampled in the latter part of the campaign. Signatures of recent (previous 10 days) origins were superimposed on the large-scale westerly flow over West Africa. In the lower TTL, air masses were impacted by recent local deep convection over Africa at the level of main convective outflow (350 K, 200 hPa) and on certain days up to 370 K (100 hPa). Estimates of the fraction of air masses influenced by local convection vary from 10 to 50% depending on the method applied and from day to day during the campaign. The analysis shows that flights on 7, 8 and 11 August were more influenced by local convection than on 4 and 13 August allowing separation of trace gas and aerosol measurements into “convective” and “non-convective” flights. Strong signatures, particularly in short-lived species like CO, NO and fine-mode aerosols were seen during flights most influenced by convection up to 350–365 K. Observed profiles were also constantly perturbed by uplift (as high as 39%) of air masses from the mid to lower troposphere over Asia, India, and oceanic regions resulting in import of clean oceanic (e.g., O₃-poor) or polluted air masses from Asia (high O₃, CO, CO₂) into West Africa. Thus, recent uplift of CO₂ over Asia may contribute to the observed positive CO₂ gradients in the TTL over West Africa. This suggests a more significant fraction of younger air masses in the TTL making it difficult to derive mean age of air from average gradients. Transport of air masses from the mid-latitude lower stratosphere had an impact from the mid-TTL upwards (20–40% above 370 K) during the campaign period

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importing air masses with high O_3 and NO_y . Ozone profiles show a less pronounced lower TTL minimum than observed previously by regular ozonesondes at other tropical locations. Concentrations are less than 100 ppbv in the lower TTL and vertical gradients less steep than in the upper TTL. The air mass origin analysis and simulations of in-situ net photochemical O_3 production, initialised with observations, suggest that the lower TTL is significantly impacted by uplift of O_3 precursors (over Africa and Asia) leading to positive production rates (up to 2 ppbv per day) in the lower and mid TTL even at moderate NO_x levels. Photochemical O_3 production increases with higher NO_x and H_2O in air masses with O_3 less than 150 ppbv.

1 Introduction

Understanding the processes that govern the chemical composition, microphysics and transport in the tropical upper troposphere (UT) and lower stratosphere (LS) is important for the quantification of transport of trace gases and aerosols into the stratosphere, where they can impact stratospheric O_3 depletion and the stratospheric water budget. This region, designated the tropical tropopause layer (TTL), represents a transition zone between the troposphere and stratosphere and has been defined in a number of different ways. A comprehensive review of the TTL was given by Fueglistaler et al. (2009) who, for stratospheric purposes, prefer to define the TTL as the region characterised by slow ascending motion driven by radiative heating and the large-scale Brewer-Dobson circulation extending from the region above main convective outflow (neutral buoyancy) at 355 K (150 hPa, 14 km) up to 425 K (70 hPa, 18.5 km). This region is bounded laterally by the position the sub-tropical jets (STJs) which limits meridional transport in the lower part of the TTL but not in the upper TTL (e.g., see Volk et al., 2000). Other definitions (e.g., Gettelman et al., 2004) have fixed the lower boundary at the level of main deep convective outflow (350 K, 200 hPa, 12.5 km) although Fueglistaler et al. (2009) point out that air is largely subsiding at these levels. However, the actual quantification of the flux of trace gases into the stratosphere will depend on

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the composition of the air masses that enter the TTL, their residence time in the TTL before upward transport as well as processes that can change their concentrations (chemistry, microphysics, in-mixing from mid-latitudes). Quantification of processes influencing the chemical composition of the TTL is also important for climate since several trace gases, namely H₂O, O₃ and CO₂, make significant contributions to the Earth's radiative budget at these altitudes. Tropospheric O₃, which can be produced downwind from convective regions from lightning NO_x or following uplift of ozone precursors, governs the global oxidizing capacity and the lifetime of several greenhouse gases such as CH₄, through the production of the hydroxyl radical (OH). Also, the TTL constitutes a major source of new particle formation probably contributing to the maintenance of the global, stratospheric Junge aerosol layer (Brock et al., 1995; Borrmann et al., 2010).

In this paper, we present an analysis of air mass origins influencing chemical composition of the TTL over West Africa based on aircraft data collected during the summer monsoon in 2006 as part of a joint AMMA (African Monsoon Multidisciplinary Analysis) and SCOUT-O3 (Stratospheric-Climatic Links with Emphasis on the Upper Troposphere and Lower Stratosphere) airborne campaign. We analyse flights made by the M55-Geophysica aircraft between 12 and 20 km during August 2006 as part of a multi-aircraft measurement deployment. The SCOUT-AMMA M55 aircraft campaign is one of the first campaigns to take place in the Northern Hemisphere summer in a region influenced by both the African and Asian monsoons. Previous tropical campaigns have taken place in other locations and other times of year and focused primarily on diagnosing contributions from local deep convection and advection of air from the mid-latitude lower stratosphere. These campaigns took place in Southern Hemisphere summer over Australia (Brunner et al., 2009) and Brazil (Pommereau et al., 2009; Huntrieser et al., 2008) and Northern Hemisphere winter over Costa Rica (e.g., Marcy et al., 2007). One of the main aims of the West African M55 campaign was to quantify the contribution from different air mass origins on TTL chemical composition during the summer monsoon. At this time of year the TTL over West Africa can be impacted (a) by

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advection of air masses from upwind regions, and b) by recent regional convective up-
lift. Air masses advected into the TTL over West Africa may have been influenced by
uplift from the lower troposphere over upwind regions (e.g., Asia), intrusion of air from
the mid-latitude lower stratosphere or cross-hemispheric transport of air masses from
the Southern Hemisphere.

The meteorological situation during the summer monsoon over West Africa is dis-
cussed in Sect. 2. A brief description of the M55 campaign is given in Sect. 3. The
fractions of air masses arriving from different upwind regions via large-scale transport
are estimated based on back trajectory calculations to quantify the influence of differ-
ent origins (Sect. 4). Comparison of results calculated for a West African domain with
those calculated in the region or along M55 flight tracks allows the flights to be put into
a broader context during August 2006. The influence of recent convective uplift over
central and western Africa, diagnosed using coincidences between back trajectories
and convective cloud tops is discussed in Sect. 5. Results are compared with results
from mesoscale model simulations (Fierli et al., 2010). The results are used to divide
the M55 flights into those more or less influenced by recent convective uplift over Africa.
The influence of large-scale transport and regional convective uplift on measured M55
trace gas and aerosol profiles is then discussed in Sect. 6. Conclusions are presented
in Sect. 7.

2 Meteorological situation

A detailed discussion about the characteristics of the summer monsoon over West
Africa and comparisons between 2006 and 1990–2005 can be found in Janicot et
al. (2008). A description of the general meteorological situation during July and August
2006 in the upper troposphere and lower stratosphere (UTLS) including relation to the
passage of convective systems is also given in Cairo et al. (2010). By way of a brief
summary, the circulation in the lower troposphere is governed by dry north-easterly
Harmattan winds circulating round the Saharan anticyclone, and moist south-westerly

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winds circulating round the southern Atlantic Santa Helena anticyclone which converge at the Inter-Tropical Convergence Zone (ITCZ) around 10° N. During August 2006, the surface Inter-Tropical Front reached as far north as 12° N. This convergence zone, and the presence of elevated terrain to the east, leads to the formation of large organized mesoscale convective systems (MCSs) which are modulated by the propagation of African easterly waves tracking from east to west and the position and strength of the African easterly jet (AEJ): a mid-level thermal wind between 10–15° N. In August 2006, convective activity and resulting precipitation were slightly enhanced compared to the 1990–2005 average (Janicot et al., 2008).

Trace gases and aerosols injected into the TTL are transported rapidly westward due to the existence of the Tropical Easterly Jet (TEJ) at 200 hPa induced by upper level anticyclones related to the African and Asian monsoons. This can also lead to transport of lower stratospheric air into the TTL over the Himalayas (Gettleman and Forster, 2002). Whilst the TEJ has maximum speeds over the Indian Ocean it extends into West Africa in July and August with a core evident around 10–30° W. The TEJ over West Africa was stronger compared to the mean in August 2006 (Janicot et al., 2008). This westerly flow dominates the large-scale circulation but chemical composition over West Africa can also be influenced by more recent uplift into the TTL. As noted in Fuglistaler et al. (2009) and references therein, the propagation or breaking of waves such as planetary scale, quasi-stationary Rossby waves (to the west) or Kelvin waves (to the east) can lead to transport of air from the stratosphere to the troposphere. Propagation of Kelvin waves can also modulate convection although Janicot et al. (2008) note that this activity was low in August 2006 over West Africa. Transport of air masses between the extra-tropical stratosphere and the tropics is also influenced by the phase of the Quasi-Biennial Oscillation (QBO). As discussed in Cairo et al. (2010), the QBO was in the process of switching from westerly to easterly winds in the lower stratosphere and thus from a westerly phase more likely to favour extra-tropical transport into the tropical lower stratosphere to an easterly phase where this is less likely (e.g., O’Sullivan and Dunkerton, 1997). Analysis of N₂O data collected during the M55 campaign over West

Africa estimated a contribution from in-mixing of aged extra-tropical air into the West African TTL varying between 10% at 350 K to 30% at 390 K (Homan et al., 2010).

Transport of pollutants into the upper troposphere over Asia by deep convection followed by westerly transport has been identified as a source of carbon monoxide over West Africa (Barret et al., 2008). Analysis of Microwave Limbs Sounder (MLS) data and global model results showed an influence from this source above 150 hPa over West Africa with maximum concentrations around 25° N. Convective uplift of moist air masses over the Bay of Bengal and China Sea followed by dehydration during transport towards North-west India (around the upper level anticyclone associated with the Asian monsoon) has also been used to explain the water vapour maximum observed in MLS water vapour data at 100 hPa over the Himalayas (James et al., 2008). Analysis of water vapour data collected on the M55 during SCOUT-AMMA in August 2006 have been linked to this convective uplift followed by westerly transport leading to enhanced water vapour concentrations over West Africa (Schiller et al., 2009).

At lower altitudes in the upper troposphere (200 hPa), CO maxima were observed further south in satellite and MOZAIC CO data during summer 2006, especially over central Africa with a maximum at the Equator. This has been explained by convective uplift of biomass burning pollutants (located south of the Equator at this time of year) and recirculation southwards by the large-scale Hadley circulation (e.g., Barret et al., 2008). Westerly transport of these emissions by the TEJ over southern West Africa may also explain higher CO concentrations in the upper troposphere over this region. Mari et al. (2008) examined this cross-hemispheric transport in detail showing that it is episodic and dependant on the strength of lower and mid-tropospheric transport from central Africa into the Gulf of Guinea. During periods when this transport is less active (e.g. 3–9 August 2006), pollutants remain trapped over the continent before being either transported directly to West Africa in the mid-troposphere or to the north-east where they can be convectively uplifted into the upper troposphere. During AMMA, polluted plumes originating from central African biomass burning were clearly observed in the mid-troposphere over the Gulf of Guinea (Reeves et al., 2010) as a result of

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direct transport in the lower/mid-troposphere. Real et al. (2010) also found evidence for transport of biomass burning plumes into the lower TTL over West Africa (M55 flight on 13 August, 2006).

Deep convection over western and central Africa also plays an important role in influencing trace gas concentrations in the TTL. During the monsoon season, large organized MCSs develop on a regular basis resulting in transport of air into the upper troposphere albeit with large-scale subsidence in surrounding regions at the level of convective outflow (around 200 hPa). Their impact on trace gas concentrations in the TTL is complex and will depend on where air masses were uplifted, on mixing with air in the upper troposphere, as well as on possible liquid and ice phase cloud processing. Convective systems are also a significant source of lightning NO_x over West Africa (Höller et al., 2009) which can lead to photochemical O₃ production downwind (Sauvage et al., 2007a; Barret et al., 2010).

In summary, the chemical composition of the TTL over West Africa is influenced by transport of air masses into the region primarily from the east. Chemical composition can also be influenced by additional inputs from the lower stratosphere, by cross-hemispheric transport of air masses from central Africa, or by local deep convection over Africa. The contribution of these different sources is examined in Sects. 4 and 5.

3 The M55 SCOUT-AMMA airborne campaign

The M55-Geophysica campaign, which took place from 1 to 16 August 2006, was part of a larger AMMA chemistry and aerosol field intensive designed to characterize the processes influencing chemical composition over West Africa. The British BAe-146 (B146), French Falcon F-20 (FF-20) and ATR-42 were based in Niamey, Niger (13.5° N 2.2° E) and the M55 and DLR-Falcon (DF-20) were based in Ouagadougou, Burkina Faso (12.4° N, -1.5 E). Details about the coordinated aircraft deployment and main findings can be found in Reeves et al. (2010). In this paper, we focus on analysis of M55 data, which sampled the TTL and LS between 12 and 20 km. The M55 made 5

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structure of the TTL over West Africa. Figure 2 (left panel) shows the average temperature profile as a function of potential temperature acquired south of 17.5°N . The average cold point tropopause can clearly be seen at around 375 K although it varied by $\pm 5\text{ K}$ between different flights. Analysis of tropical radiosonde data showed similar results during boreal summer (Fueglistaler et al., 2009). The total number of observations at a given theta level (Fig. 2 central panel) shows that the lower TTL was the most heavily sampled during the campaign. The right panel shows the fraction of observations in cloudy/hazy air masses derived from volume backscatter ratios greater than 1.2 measured with MAS (Adriani et al., 1999). There is a bimodal distribution with peaks at 320 K and 350 K representing main convective outflow altitudes. Above this level there is a small percentage (around 0.01%) of cloud / haze observations up to about 410 K suggesting that tropospheric influence may not have extended much above this level during the campaign over West Africa although some evidence for convective overshooting has been found (Khaykin et al., 2009). A thin aerosol layer was also observed in the LS at 475 K related to the presence of volcanic aerosols (Borrmann et al., 2010).

For the purposes of this analysis we use the following definitions: top of TTL at 410 K (70 hPa, 18.5 km) above which air is essentially stratospheric, the mid-TTL as the boundary between the upper and lower TTL at 370 K (100 hPa, 16.5 km) just below where the cold point tropopause is located, and the lower TTL between the mid-TTL and the lower boundary defined by the altitude of maximum of convective outflow at 350 K (200 hPa, 12.5 km). The layer above the altitude of maximum convective outflow (355 K, 150 hPa, 14 km) is also of interest.

4 Large-scale transport

In order to identify the impact of recent air mass origins on measured chemical composition, 10 day reverse domain filling trajectories from a 0.5 degree grid covering West Africa (10°E to 40°W ; 5 to 25°N) were calculated at 4 pressure levels from 1 to 16 August, 2006. Results based on 20-day back trajectories give very similar results

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5 impact on altitudes in the mid and upper TTL over West Africa with a mean influence over 16 days of 41% and 15% at 100 hPa and 70 hPa, respectively. These air masses have been transported around the Tibetan High, especially at 100 hPa. However, there is significant daily variability, ranging from 13 to 73% at 100 hPa, for example. The average values at 100 hPa are in reasonable agreement with Konopka et al. (2009) who estimated the fraction of air masses entering the TTL from the LS associated with the Asian summer monsoon upper level anticyclone. A significant fraction of air masses are also influenced by uplift of tropospheric air over Asia, especially at 150 hPa (39% on average) and 100 hPa (21% on average). Uplift of air masses over Africa (north of the Equator) has the most impact at 200 hPa (52% on average). In this analysis, based on trajectories calculated using ECMWF analyses, uplift from below 400 hPa is considered. Further analysis (not shown) shows that most trajectories originate from the mid-troposphere. Whilst such trajectories are able to capture large-scale uplift they are likely to underestimate uplift associated local deep convection. In this case they may underestimate the number of convective events as well as vertical velocities and the altitude to which convection has an impact (e.g., see Fierli et al., 2010). Therefore, the results presented in Fig. 4 may represent a lower limit. Air masses originating from SH central Africa (all altitudes) have their main influence in the latter half of the campaign period leading to a lower average influence (6%). Transport through this region is associated with southerly deviations in the position of the TEJ. Interestingly, at the end of the campaign period, these air masses also influence 150 hPa and 200 hPa (up to 23%) related to the development of large MCSs over Chad/Sudan (Real et al., 2010).

25 We now address whether, and to what extent, the sampling of the M55 was representative of the air masses arriving over West Africa during the campaign. The fraction of trajectories passing through the different regions before arriving in the sub-domain where the flights took place is shown in Fig. 4b. In general, the M55 flight region experienced similar fractions of air masses originating from the free troposphere over Africa at 200 hPa (52% on average), Asia at 150 hPa (39%), and 100 hPa (21%). At

100 hPa, the M55 region saw less air arriving from NH mid-latitudes (21% compared to 41% at 100 hPa) and more from SH central Africa (13% compared to 6% at 200 hPa) compared to the larger West African domain. These biases need to be borne in mind when analysing M55 flight data.

5 We now focus on discussing the origins of air masses sampled by the M55 in more detail. Figure 5 shows the starting points of 10-day back trajectories initialised along the M55 flight tracks (on average 200 points per flight with 1 min time interval). The results are colour-coded in terms of pressure variations along the back trajectory (left panel), and pressure at the ending point on the M55 flight path (right panel). As noted
10 already, air in the upper TTL has mainly resided in the TTL before arriving over West Africa and shows only small theta deviations over the last 10 days. A noteworthy fraction of upper TTL (>100 hPa) air originates from further north and west over Japan and eastern China/Korea consistent with the NH mid-latitude origin shown in Fig. 4. Figure 5 also shows that the origin of air in the mid TTL (around 150 hPa) is more
15 widespread with larger pressure variations, and air masses uplifted from Asia, India and, to a lesser extent, the Middle East.

To explore this further, Fig. 6 shows scatter plots of the difference in pressure versus the difference in latitude along the back trajectories coloured with O₃ (left panel) and H₂O (right panel) observed along the M55 flights. Ozone concentrations in air
20 masses originating from the mid-lower troposphere are generally below 125 ppbv and sometimes lower than 50 ppbv. This variability can be due to uplift of O₃ poor (marine) or rich (polluted) air masses and mixing with air masses already residing in the TTL. Cross-hemispheric transport of air masses via the TEJ with high O₃ and low H₂O are also evident (negative latitude changes and small pressure changes). The variability in
25 observed O₃ concentrations is discussed further in Sect. 6. Concerning H₂O, Schiller et al. (2009) noted that the TTL over West Africa during summer 2006 was moister compared to other tropical campaigns. Their analysis pointed to convective uplift of moist air masses associated with the Asian monsoon followed by dehydration during westerly transport around the upper level anticyclone leading to drying, but nevertheless moister

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air masses arriving over West Africa. This analysis is also in agreement with the study of James et al. (2008) who investigated the moist H_2O anomaly associated with the summer Asian monsoon seen in MLS data. This is confirmed by Fig. 6 which shows that a significant number of air masses sampled by the M55 were uplifted. At the point of measurement, although mainly low H_2O was observed, higher H_2O concentrations were also observed. Figure 7 shows where 10 day back trajectories arriving along the M55 flights left the lower troposphere, also coded in terms of O_3 (left panel) and H_2O (right panel). In this case a more strict criteria was used to define uplift as being from below 800 hPa. The points on the maps, although scattered, show a band extending from eastern Africa to south-east Asia. Interestingly, for H_2O there appears to be an east–west gradient in concentrations with air masses associated with more recent uplift over the Arabic peninsula having higher H_2O concentrations upon arrival over West Africa than those originating from further east over Asia. Recent convective influence could also explain higher H_2O concentrations in the TTL over West Africa.

This analysis shows that, in general, air in the upper TTL over West Africa, and in the region of the M55 flights, was influenced by a variety of more recent (10 day) air mass origins superimposed on the large-scale westerly flow. The mid and upper TTL was influenced by recent import of lower stratospheric air from NH mid-latitudes. The mid-TTL also shows a significant influence from air masses uplifted from the lower and mid troposphere over Asia, and India whereas the lower TTL was mainly impacted by air recently uplifted from the troposphere over Africa.

5 Recent regional convection

The influence of local convective uplift occurring over western and central Africa is examined further in this section. Coincidences were identified between ECMWF back trajectories over the whole West African domain (as described in the previous section) and convective cloud tops derived from Meteosat Second Generation (MSG) Infrared Channel ($10.6\ \mu\text{m}$) observations. A convective cloud was identified when cloud top

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radiance temperatures were below 200 K. The time resolution of the MSG images was 30 min and horizontal resolution ranged from 5 to 10 km. Since the images covered the longitudinal band 30° W to 40° E, the convection was local and relatively recent, i.e. given the average westward speed of the MCS, not older than 3–4 days. Figure 8 shows daily maps of this diagnosed convective impact at 200 hPa from 1 to 16 August.

The results show that large regions over West Africa were constantly influenced by recent local convection during the timeframe of the M55 campaign, making it difficult to define an “unperturbed” state in the lower TTL with respect to convective influence at this altitude. The fraction of convectively processed air (%) is also shown for each day in Fig. 8. This represents the fraction of trajectories that have passed over a convective system over Africa, as identified by MSG. At least 5% of trajectories need to have encountered convection in order to be included in the calculation of the average impact for each day. This provides an estimate about the regional influence of convection on different days during the campaign at the main convective outflow level (200 hPa). The flights on 8, 11 and 12 August were more influenced by convection compared to other days with fractions reaching 12–15%. These results can be compared with estimates of local convective impact along the flight tracks also based on coincidences between back trajectories and MSG clouds (Fig. 9). The average percentage of trajectories arriving between 350–360 K are shown as a vertical bar since it is not possible to discriminate differences in altitude based on coincidences with MSG convective clouds. The percentages are larger than those calculated for the whole domain at 200 hPa and can be explained by the location of particular flights and the fact that Fig. 9 also includes coincidences up to 360 K. In general, the along flight estimates show similar variations between the different days as seen for the whole region (Fig. 8) except for 7 August which shows more convective influence in the along flight analysis. The results in Fig. 9 confirm that the flights on 8 and 11 August were most impacted by local convection. The flight on 11 August flew in the outflow of a particularly large MCS system, which developed over Niger and Nigeria. The flight on 8 August was planned as a CALIPSO validation flight but was nevertheless downwind of a large MCS which had developed

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further east. On 7 August, the aircraft flew in the outflow of a small MCS that tracked across Niger, Burkina Faso and Mali (Cairo et al., 2010). The fact that this was a rather small system explains the lower along flight fractions in Fig. 9 and the low average regional convective impact in Fig. 8. The flights on 4 and 13 August, when the M55 flew to south to the Gulf of Guinea, show less convective influence, particularly on 4 August. On 13 August the southern part of the flight at 200 hPa was influenced by downwind transport of convectively uplifted air masses further east over central Africa (Real et al., 2010). In the case of the flight on the 4 August, it was possibly influenced by local sporadic convective near Ouagadougou, at the beginning and end of the flight.

These results can also be compared to a more detailed study described in a companion paper, Fierli et al. (2010), where a mesoscale model was used to calculate trajectories for M55 flights on 7, 8 and 11 August 2006. The BOLAM simulations, covering the latitude band 15° W to 45° E, were nudged with MSG brightness temperatures leading to a more realistic representation of convective systems during the campaign period (see also Orlandi et al., 2010). Their estimates of convective fraction (derived using model calculated back trajectories originating below 800 hPa) varied a lot between flights and with altitude. In general, Fierli et al. (2010) estimated significant convective impact over West Africa extending to higher altitudes although the results for different days are broadly consistent with the analysis presented in this paper. On 7 August up to 20% of the sampled air masses were influenced by convection up to 360 K. The influence was higher on 8 August (70% up to 360 K) and higher still on 11 August (over 80% up to 370 K). The flights on 7 and 8 August also had some convective influence at higher altitudes.

Figure 9 also compares local convective influence, based on coincidence with convective clouds, with the impact of large-scale uplift diagnosed from ECMWF trajectories along the flights. The large scale uplift is diagnosed as the percentage of 10-day back trajectories arriving in 4 K altitude bins between 340 K to 374 K along the M55 flights which have experienced irreversible uplift above 650 hPa (indicative of planetary boundary layer top in ECMWF analyses). To discriminate between uplift over Africa and

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and roughly $1\ \mu\text{m}$, denoted N_{14} , it can be seen that during August 2006, more variable aerosol concentrations were observed at 350 K in the “convective” flights compared to the “non-convective” flights. There also appears to be a minimum in concentrations around 355 K and some air masses have low aerosol below 360 K possibly due to washout in convective systems. In both “convective” and “non-convective” flights, concentrations decrease towards the top of the TTL (420 K). In the lower TTL, the origin of the aerosols is unknown but may be related to uplift of precursor emissions or recent nucleation. Figure 11 shows the correlation between fine-mode aerosols and CO for the 7 and 8 August with air masses influenced by recent nucleation removed. In this case, N_6 (6 nm to $<1\ \mu\text{m}$) is plotted and the difference between N_6 and N_{14} was used to remove nucleation events associated with recent convection (see also Borrmann et al., 2010). Figure 11 shows a clear positive correlation suggesting a common origin. This could either be from biomass burning over Asia or central Africa or from anthropogenic sources. However, Borrmann et al. (2010) show a minimum in the fraction of non-volatile particles between 340–370 K. This suggests the presence of aerosols, like sulphate, which could have been transported from anthropogenic sources in Asia, rather than a biomass burning source (see also CO discussion). Interestingly, in the upper TTL, where aerosol concentrations are lower, there is a maximum (up to 50%) in the fraction of non-volatile particles indicating the presence of non-sulphate aerosol. These air masses are more aged (see discussion about CO_2) and may represent a residual layer of older aerosols that have not been removed by washout at lower altitudes.

CO was only measured during the “convective” flights on 7 and 8 August. The observations show considerable variability (45 to 100 ppbv) up to 370 K and generally decreasing concentrations above (see Fig. 10a). The range of observed CO, which has a lifetime of 2–3 months, reflects a combination of recent inputs into the TTL and mixing with more aged air masses. Comparison with data collected in lower troposphere can be used as an indicator of convective uplift. Over West Africa during late July/August 2006, lower tropospheric CO varied from 145 ppbv over southern tropical forests to 80–90 ppbv north of 14°N (Sauniois et al., 2009; Reeves et al., 2010). An

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analysis of all aircraft vertical profile data collected during July–August 2006 showed a maximum in CO concentrations in the upper troposphere (200 hPa) (Bouarar et al., 2010) and that flights influenced by local convection had more variable CO ranging from 80 to 200 ppbv compared to “non-convective” flights (100–150 ppbv). The M55 data is at the lower end of this range (80–100 ppbv at 200 hPa). CO concentrations ranged between 70–90 ppbv in air masses where aerosol nucleation, likely to be associated with recent convection, was diagnosed to have taken place.

As already mentioned in Sect. 3, analysis of MLS CO data over the Tibetan plateau shows the existence of a positive CO anomaly up to 100 hPa (Randel et al., 2007), with concentrations in the range 80–100 ppbv attributed to deep convective uplift of anthropogenic and biomass burning emissions, particularly over India and South-east Asia (Park et al., 2009). Surface CO concentrations measured over Asia are highly variable and depend on whether a site is influenced by marine or polluted air masses. For example, data collected in southern Thailand shows concentrations varying from 50 ppbv to several hundreds of ppbv (Pochanart et al., 2003). However, it does appear that uplift in the Asian summer monsoon, followed by slow ascending transport around the upper level Tibetan anticyclone leads to enhanced CO up to 370 K at this time of year. Some of this pollution is transported westward by the TEJ and contributes to enhanced CO over southern West Africa as seen in satellite, MOZAIC and aircraft data (Barret et al., 2008; Reeves et al., 2010). As well as CO, Park et al. (2008) also showed enhancements up to 20 km in trace gases measured by the ACE satellite that are markers of surface pollution such as HCN (an indicator of biomass burning) and hydrocarbons over a region extending from 0–120° E and 10–40° N during June to August 2004–2006. However, MODIS fire count data show that biomass burning was not very prevalent during the latter half of July, or the beginning of August in 2006 (see <http://rapidfire.sci.gsfc.nasa.gov/firemaps/>) suggesting that anthropogenic sources were more important. This is also supported by the fine-mode aerosol analysis discussed previously. Therefore, since CO concentrations in air masses originating either from local African convection or upwind convection can have similar concentrations

in the TTL, it is difficult to separate these two origins using CO alone. The measurements of aerosols, NO_y and as discussed later, CO₂, support the hypothesis that the lower-mid TTL was strongly influenced by uplift of lower tropospheric air masses.

One exception is an event with very low CO concentrations (<50 ppbv) observed between 350 and 360 K that was also subject to uplift from the lower troposphere. These air masses had low H₂O suggesting dehydration during transport. However, surface measurements of lower tropospheric CO over Asia reported by the US National Oceanic and Atmospheric Administration (NOAA) Carbon Cycle Greenhouse Gas (CCGG) monitoring network (<http://www.esrl.noaa.gov/gmd/ccgg/iadv/>) do not show values below 50 ppbv. This, and the fact that O₃ concentrations were not particularly low (50–60 ppbv), suggests that uplift from an oceanic region followed by mixing with air of lower stratospheric origin had taken place, even in the lower part of the TTL. The analysis presented in Sect. 4 suggests a lower stratospheric influence of 10–20% at 350–360 K in the M55 flights. However, this contribution may have been underestimated in this case.

CO₂ was measured during “convective” flights influenced by primarily by African convection on 11 August and uplift over India on 8 August. Below 360 K, air masses that were uplifted show variable CO₂ concentrations between 374 to 377 ppmv. Flights of the DLR-Falcon (up to 200 hPa) found slightly enhanced CO₂ concentrations (379 to 380 ppmv) in convectively uplifted air masses compared to more typical upper tropospheric values (374 to 376 ppmv) (Andrés-Hernández et al., 2009). Low concentrations observed up to 368 K on 11 August have been attributed to active local convective uplift on that day (Fierli et al., 2010; Homan et al., 2010) transporting boundary layer CO₂ with low CO₂ concentrations into the TTL. These air masses were associated with higher H₂O, some enhancement in aerosols and O₃ concentrations around 60 ppbv. It is likely that daytime drawn down of CO₂ over tropical forests led to lower concentrations being uplifted (Williams et al., 2001). Apart from this low CO₂ event, a positive gradient in CO₂ is seen in the lower and mid-TTL in these “convective” flights. Above this altitude CO₂ starts to decline below the top of the TTL (425 K) reflecting an

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increasing contribution from lower stratospheric air masses: 20–30% based on Fig. 4 and N₂O data analysis presented by Homan et al. (2010). The CO₂ profiles measured on 4 and 13 August also show CO₂ increasing with altitude above 350 K except for one low CO₂ event between 360–368 K. An event with higher CO₂ (379–380 ppmv at 200 hPa) has been linked to convective uplift of central African biomass burning emissions on 13 August (black points in Fig. 10b) (see Real et al., 2010). This event, which is not diagnosed in the ECMWF trajectories, was accompanied by enhanced concentrations of O₃, NO, NO_y and aerosols. A similar event was observed on 4 August in the aerosol data.

The constant perturbation of the CO₂ profile by deep convection over Africa or by convective uplift to the east resulted in concentrations as high as 382 ppmv being observed in the TTL over West Africa. Note that concentrations observed during summer 2006 at NOAA CCGG surface sites in Asia varied from 373 to 382 ppmv (see <http://www.esrl.noaa.gov/gmd/ccgg/iadv/>). Measured TTL CO₂ during the 8 August flight that was particularly influenced by uplift over India (see Fig. 9) was in this range. Given the long lifetime of CO₂, these recent signatures are superimposed on the large-scale picture. The M55 profiles are partly in accordance with the notion of slow ascent of air masses injected several months earlier during Northern Hemisphere spring when surface CO₂ concentrations are higher (e.g., Boering et al., 1996, and supported by Homan et al., 2010). However, measured profiles of CO₂, CO etc. up to 370 K also reflect more recent origins including local (last 3–4 days) and upwind (last 10–20 days) convection. This suggests that the age spectrum of air masses in the lower and mid-TTL may include a higher component of younger air masses than previously assumed, for example in studies estimating the age of air in the mid-TTL at the cold point tropopause (e.g., Park et al., 2007; Gettelman et al., 2009). Age of air depends on location and season. In the Northern Hemisphere summer, the Asian monsoon is an effective transport mechanism for transporting trace gases rapidly up to 370 K. The African monsoon is also effective at transporting air masses up to at least 360–365 K.

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The interpretation of the measured O_3 profiles over West Africa during the M55 campaign is complex because O_3 has competing sources in the stratosphere and the lower troposphere. Analysis of tropical ozonesonde data from the SHADOZ network (e.g., Folkins et al., 2002; Thompson et al., 2007) shows the presence of an O_3 minimum, varying between 30 and 60 ppbv, in the region of convective outflow (200 hPa, 350 K). This is often explained by the uplift of O_3 -poor air from clean marine regions. Concentrations increase rapidly with altitude to the top of the TTL. This is attributed to an increasing contribution from import of LS air masses particularly above the cold point tropopause (e.g., Konopka et al., 2010). Previous studies have also shown significant longitudinal variations in upper tropospheric O_3 related to the large scale circulation in the troposphere and lower stratosphere, leading to, for example, a maximum over the Tibetan Plateau linked to the formation of the summer Asian monsoon (Randel et al., 2007) and, the formation of a semi-permanent maximum in the UT over the southern Atlantic Ocean due to the recirculation of air masses rich in O_3 from central African biomass burning regions by the Hadley circulation (e.g., Sauvage et al., 2006; Sauvage et al., 2007b). Therefore, it is clear that O_3 concentrations in the TTL are variable and depend on location and time of year. In addition, O_3 and its precursors such as CO or NO_x are highly variable in the troposphere. This makes it difficult to justify the use of seasonal or annual average vertical profiles to estimate the impact of deep convection and transport into the stratosphere, as was the case in recent studies (Konopka et al., 2010; Gettelman et al., 2009).

In the lower TTL, as part of AMMA, analysis of MOZAIC and MLS O_3 data at 200 hPa over western and central Africa during the summer 2006 monsoon, showed a maximum at 10° S, a minimum at 5° N, and increasing concentrations further north (Barret et al., 2010) giving rise to a pronounced latitudinal gradient varying from 45–50 ppbv over the southern coast of West Africa to more than 70–80 ppbv over the Sahara. The minimum over southern West Africa is related to convective uplift of O_3 -poor air masses from forested regions (high O_3 dry deposition, Sauniois et al., 2009) or oceanic regions, and coincides with the observed CO maximum. Reeves et al. (2010) also noted latitudinal

differences in ozonesonde data collected in July and August 2006 at Cotonou, Benin (6° N, 2° E) and Niamey, Niger (13° N, 2° E), further north. Although minimum concentrations at 200 hPa were slightly lower over Cotonou, a pronounced upper tropospheric maximum is present over Niamey between 450–250 hPa most likely due to photochemical production of O₃ from lightning NO_x emissions or from uplift of a mixture of soil NO_x and biogenic hydrocarbons (advected northwards in the lower troposphere, Saunois et al., 2009).

The measured M55 O₃ profiles reflect the different combination of air mass origins as discussed already for other trace species. Whilst it is difficult to generalise, the “convective profiles” (Fig. 10a), which were primarily collected north of 10° N, generally have higher O₃ concentrations at 350–360 K than the “non-convective” profiles (Fig. 10b) which include descents to 350 K (200 hPa) at 5° N and have lower concentrations. The M55 data are within the range measured by the DLR Falcon during the same period (Andrés-Hernández et al., 2009) up to 150 hPa. They also found evidence for higher O₃ concentrations in regions of convective outflow coincident with elevated levels of O₃ precursors from lightning NO_x emissions and surface uplift as mentioned previously.

Figure 12 shows the average O₃ profile as a function of potential temperature, using data from all local flights, together with the fraction of data points lying three standard deviations outside the mean. As discussed above, the lower part of the ozone profile can be understood in terms of local recent deep convection with a local maximum at 340 K and a weak minimum at 350 K. In the lower TTL, O₃ concentrations increase but at a slower rate than above 370 K (roughly 2 ppbv per K below 370 K and 6 ppbv per K above 370 K). This is due to the important contribution from uplift of tropospheric air masses as confirmed by the pressure variations along the back trajectories (rightmost panel in Fig. 12) as well as enhanced variability in the O₃ data up to 370 K. Average concentrations in the lower TTL are also below 100 ppbv, again indicating the importance of uplift of tropospheric air masses. Above 370 K, O₃ concentrations are higher at this time of year compared to measurements made at other times of year at other tropical locations due to seasonal variations in lower stratospheric temperatures and

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stratospheric up-welling (Randel et al., 2007; Homan et al., 2010) and the important contribution from LS import via the Asian monsoon anticyclone as shown in this study and by others (e.g., Konopka et al., 2010).

The data presented here show that the M55 observed non-negligible concentrations of NO and CO up to 370 K suggesting that the lower TTL is potentially more photochemically active than previously assumed, and that the Asian and African monsoons provide an important mechanism for transporting pollutants into this region. Real et al. (2010) used a photochemical trajectory model initialised with AMMA data (for the biomass burning plume case on 13 August 2006) to calculate rates of in-situ net O₃ photochemical production, NPO₃, downwind from West Africa in the upper troposphere. They found that slow but positive NPO₃ (in runs with and without mixing with surrounding air masses) can be maintained even in air masses with NO concentrations of 200 pptv. Here, we performed further runs of the CiTTyCAT photochemical model taking initial conditions of O₃ precursors such as VOCs and NO_y species from the background conditions in the upper tropospheric case of Real et al. (2010). CO was initialised at 90 ppbv based on M55 data. The model was run for 4 days at 150 hPa and average NPO₃ calculated over the last 24 h. Mixing was not taken into account in the simulations. A series of sensitivity tests were carried out varying initial concentrations of O₃, H₂O and O₃ precursors (NO, CO and VOCs). As already discussed in Real et al. (2010), NPO₃ was quite insensitive to initial CO and VOC concentrations. The sensitivity of NPO₃ to initial NO, O₃ and H₂O concentration is shown in Fig. 13. NO concentrations were varied but concentrations of other NO_y species were kept the same totalling around 300 pptv (not including NO).

At O₃ concentrations below 150 ppbv NPO₃ is always positive and varies from about 0.5 ppbv per day to more than 2 ppbv per day with higher NPO₃ at higher NO and higher H₂O. H₂O and O¹D from O₃ photolysis react to produce OH, and therefore HO₂ which, in the presence of NO, leads to O₃ production. At higher O₃ concentrations of around 150 ppbv and low NO, O₃ destruction starts to dominate due to increased loss via H₂O plus O¹D. This leads to net O₃ destruction at higher O₃ and low NO but still net

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O_3 production at higher NO concentrations. In the lower TTL, air masses influenced by convection have higher NO than used here (800 pptv) which will lead to higher NPO_3 (up to several ppbv per day). In the upper TTL, H_2O is low but NO is also around 200 to 400 pptv (see Fig. 10b) and O_3 is increasing leading to small but positive NPO_3 of up to 0.5 ppbv per day. Thus, it is possible that much of the TTL over West Africa, and especially the lower TTL, at least in the lower TTL, is photochemically active and in a net O_3 production regime.

Therefore, in contrast to certain studies, such as Konopka et al. (2010), which assume that tropospheric photochemistry is not important in the TTL, the results presented here, and those previously reported (e.g., Folkens et al., 2002), show that photochemistry does play an important role in governing the O_3 budget, especially in the lower to mid TTL. The O_3 gradient seen the M55 observations (Fig. 12) below 370 K can be explained by uplift of precursors into this region, and maintenance of O_3 concentrations by in-situ photochemical O_3 production. Ideally, budget calculations need to be performed with global models including both tropospheric and stratospheric chemistry. Previous studies of the tropospheric O_3 budget have already shown that results are very sensitive to the amount of O_3 transported into the troposphere from the stratosphere (Wu et al., 2007), and often gradients of trace species are smeared out. If global models overestimate O_3 in the TTL, they are likely to underestimate NPO_3 in this region, especially if H_2O is also overestimated (as shown by Fig. 13). Lightning NO_x , on the other hand, may be underestimated in global models as shown by Barret et al. (2010) over West Africa.

In summary, trace gas and aerosol profiles measured on the M55 during August 2006 show distinct signatures of recent local convection up to 350–365 K, particularly in shorter-lived species like NO, CO and aerosols. Signatures due to uplift over upwind regions are also evident up to 370 K and superimposed on features related to longer timescale variations. This is particularly true for longer-lived species like O_3 , NO_y and CO_2 .

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The SCOUT-AMMA M55 aircraft campaign was one of the first campaigns to take place in the Northern Hemisphere summer in a region influenced by both the African and Asian monsoons. In the analysis presented here, reverse domain-filling trajectories over West Africa and in the region of the flights, as well as back trajectories along the M55 flights, have been used to diagnose the origin of air masses influencing the TTL during first half of August 2006. Most of the air arriving over West Africa was already residing in the tropical upper troposphere and lower stratosphere. However, superimposed upon this were important contributions from local recent convection over Africa, uplift of tropospheric air masses over Asia, India, and oceanic regions as well transport of air masses from the lower stratosphere during the 10 days before the measurements were made.

Local convective influence was found to be important in the lower TTL, and to have an influence on measured trace gas and aerosol profiles up to an altitude of 360–365 K, with the main influence at 350 K (200 hPa, 12.5 km) at convective outflow. Estimates based on reverse domain filling back trajectories originating from 1000–400 hPa in the African lower and mid troposphere suggest an impact of around 52% during the campaign period. Convective fractions estimated using trajectories along flight tracks show significant variability between flights with the 8 and 11 August most influenced by convection. Estimates derived from coincidences with MSG cloud tops show similar variability but fractions are lower (10–15%) because they are averages over a region and reflecting the fact that certain flights targeted convective systems. The higher estimates, based on back trajectories are generally consistent with results from Fierli et al. (2010) who found contributions as high as 60% below 355 K and, during certain flights, as high as 50% up to 370 K (100 hPa, 16.5 km). Whilst, these results highlight the need to use a variety of approaches to estimate convective impacts on the TTL, they do show that the M55 flights on 7, 8 and 11 August were more influenced by local African convection. This allowed separation of chemical and aerosol data into

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“convective” and “non-convective” flights. The “convective” flights exhibited more variability and, in the case of certain species, pronounced features in the lower TTL linked to local deep convection such as enhanced NO due to lightning emissions or uplift of soil NO_x emissions.

5 Air masses in the mid-TTL over West Africa (150–100 hPa, 14–16.5 km or 355–370 K) were also significantly impacted by uplift of tropospheric air masses with differing trace gas signatures ranging from clean air (low O₃, low CO, high H₂O) to air masses likely to have been influenced by pollution over Asia or India containing higher O₃, CO, CO₂ and NO_y. Around 40% of air masses at 355 K and 20% of air masses at
10 370 K may have been affected by such injection from the mid and lower troposphere. At 370 K, there was also an important contribution from air masses advected from the mid-latitude lower stratosphere (45%) although the M55 appears to have under-sampled this source (21%) suggesting that the lower stratospheric fractions estimated by Homan et al. (2010) may be a lower limit. The estimates of Indian and Asian uplift
15 in this study may also be a lower limit due the use of large-scale analyses to calculate back trajectories in regions influenced by deep convection associated with the monsoon. This may also lead to an underestimation of the altitude to which deep convection can have an impact.

Overall, the lower TTL over West Africa appears to have been significantly impacted
20 by the upward transport of tropospheric air masses related to the Asian and African summer monsoons. In the case of Asian uplift this results in the transport of trace gases, and possibly aerosols, directly into the lower and mid-TTL. This suggests that air in the mid-TTL may have a larger component of younger air than previously estimated. The mean O₃ profile appears to be affected by these transport processes leading to
25 a less steep gradient in the lower-mid TTL and higher concentrations than seen at tropical locations away from the influence of the Asian or African monsoons. Analysis of the observed profiles together with photochemical model calculations suggests that the lower TTL is photochemically active even at background levels of NO_x leading to positive net photochemical O₃ production. In particular, in the lower TTL higher

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H₂O and NO_x related to convection can lead to enhanced O₃ production. It remains a challenge for global models to reproduce the vertical structure of trace gases and aerosols observed in the TTL.

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Table 1. Latitudes, longitudes and pressures used to define regions of TTL air mass origins (see text for details).

Region	Latitudes	Longitudes	Pressure range (hPa)
Northern Hemisphere mid-latitudes	35–50° N	25° W–150° E	50–1000
Southern Hemisphere Central Africa	0–15° S	5° E–50° E	50–1000
Asian troposphere	0–30° N	60° E–160° E	400–1000
African troposphere	5–25° N	10° W–50° E	400–1000

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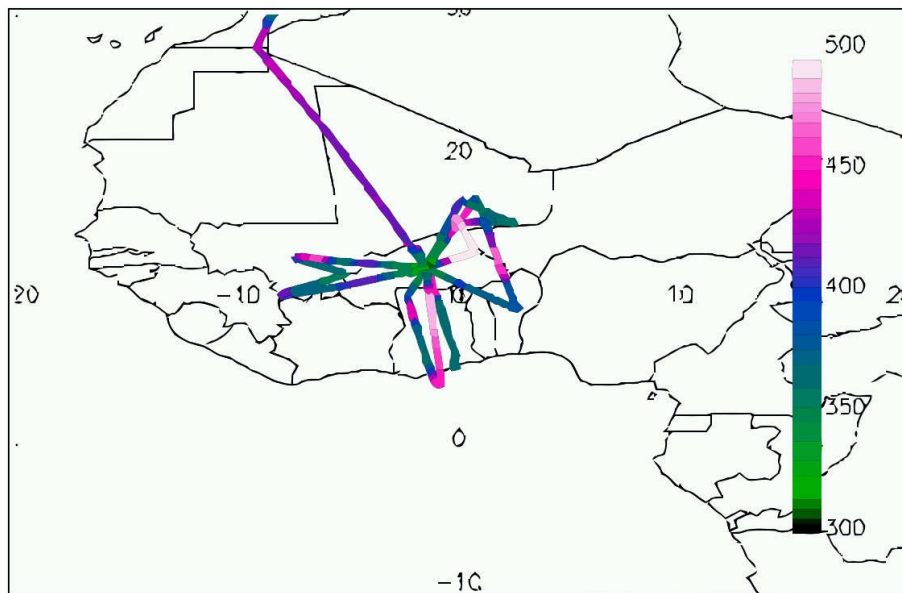


Fig. 1. M55 flight tracks over West Africa, colour coded in potential temperature (K) along the flights.

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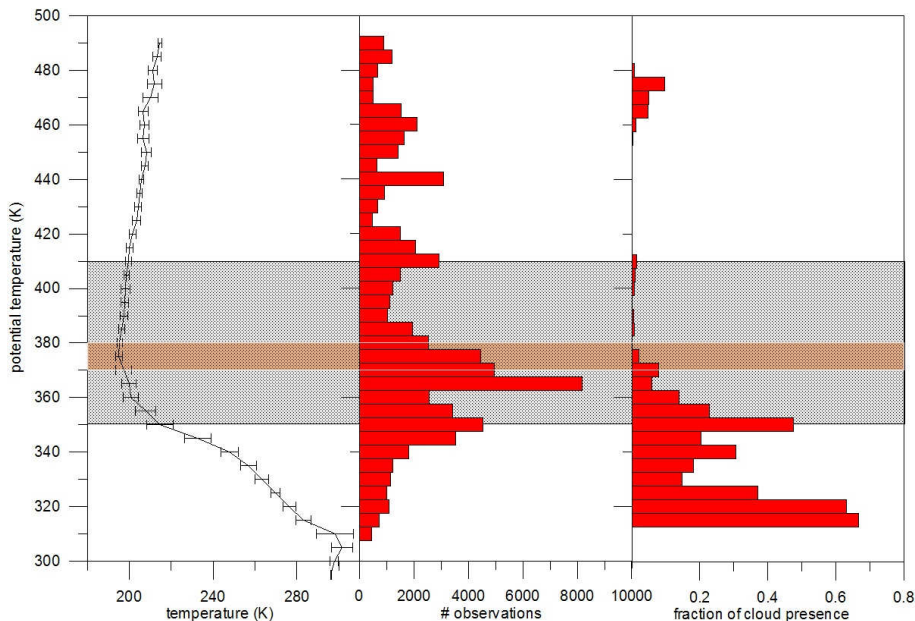


Fig. 2. Left panel: temperature versus potential temperature as measured along the M55 flights between 1–16 August 2006. Only data southward of 17.5° N were used to calculate averages. Bars denote one standard deviation. Middle panel: total number of observations in seconds spent at a given theta level versus potential temperature. Right panel: fraction of observation in cloudy (hazy) air, defined as air masses with Volume Backscatter Ratio >1.2 . Grey shading denotes TTL region and brown shading the tropopause. See text for details.

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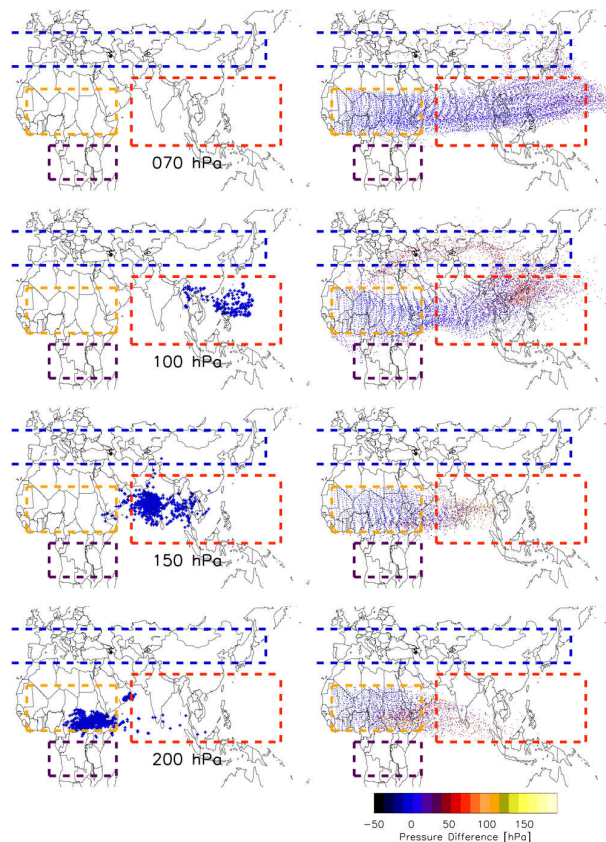


Fig. 3. Right panels: endpoints of 10 day reverse domain filling trajectories over West Africa, arriving at 200 hPa, 150 hPa, 100 hPa and 70 h Pa on 7 August 2006. Dashed boxes represent main regions of origin (see Table 1). Colours show the pressure difference along each back trajectory with respect to the arrival point. Left panels: Location of trajectories crossing 1000–400 hPa during transport to West Africa.

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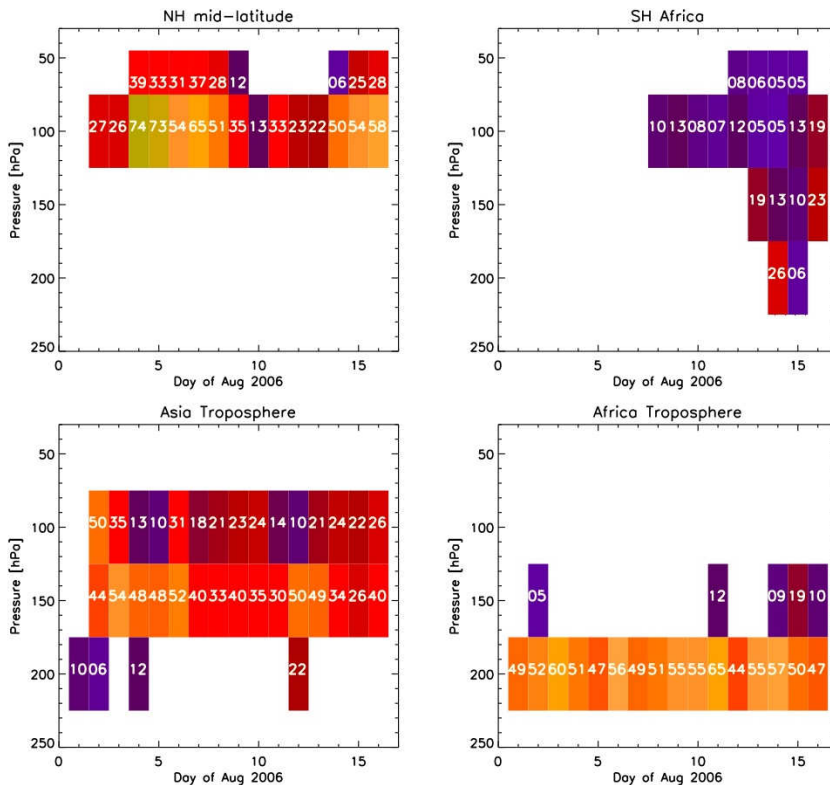


Fig. 4a. Relative contribution of different air mass origin regions (defined in Table 1) based on 10 day back trajectories to air masses arriving over the West African domain (10° E to 40° W and 5° to 25° N) between 1–16 August 2006. Colour codes the percentage of regional influence in terms of fraction of trajectories passing through a particular region (purple=low, yellow=medium, red=high). This value is also expressed numerically in each coloured cell. Note that trajectories can pass through more than one region leading to total fractions of more than 100%.

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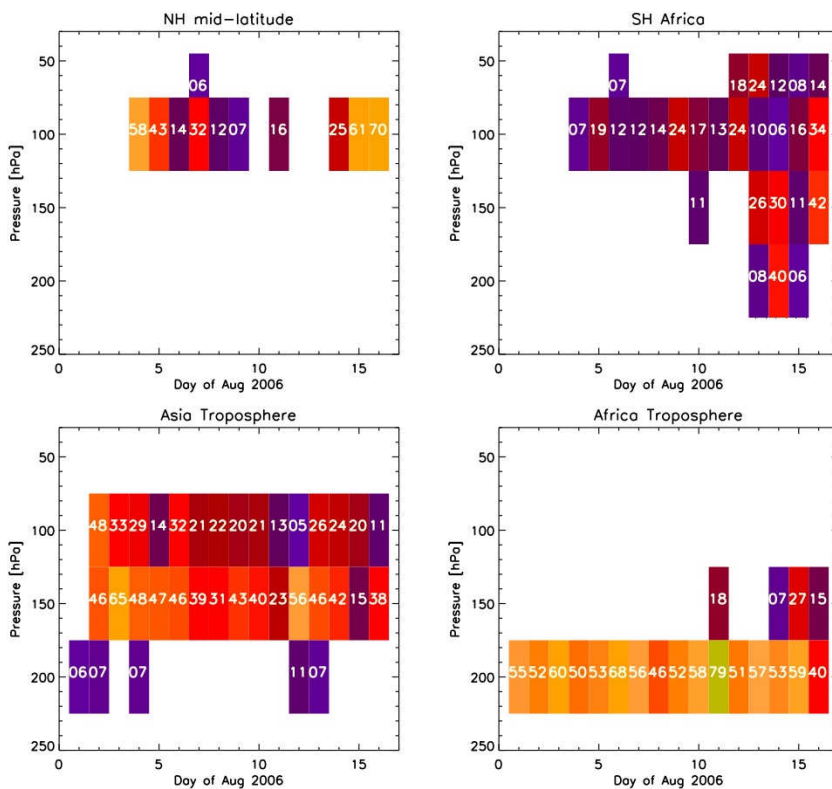


Fig. 4b. Same as Fig. 4a but based on 10 day back trajectories traced back from a sub-domain including the M55 flight tracks (8° E to 5° W and 5 to 15° N).

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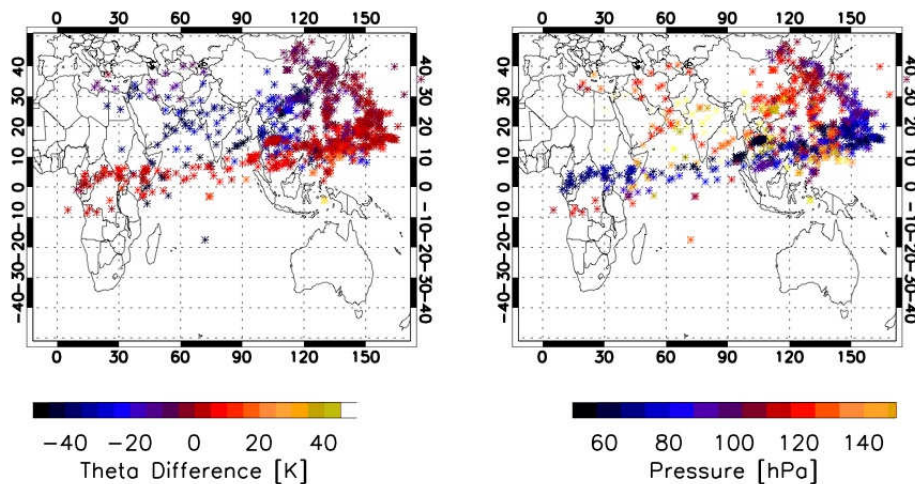


Fig. 5. Starting points of 10-day back trajectories initialised along the M55 flight tracks south of 18° N colour coded in terms of theta variations (K) along the back trajectories (left panel) and pressure at the starting point of the back trajectory, i.e. along the M55 flight path (right panel).

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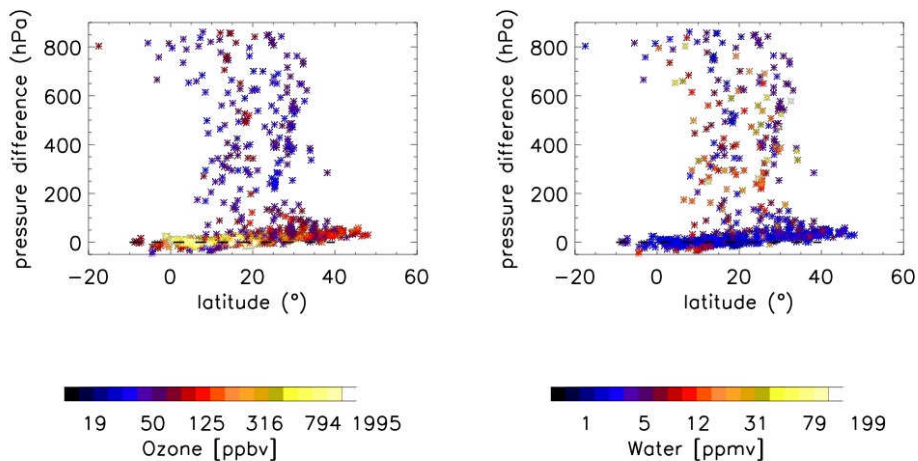


Fig. 6. Scatter plot of the pressure differential versus latitude differential along the back trajectories from M55 flight tracks, colour coded in terms of ozone in ppbv (left panel) and water vapour in ppmv (right panel) observed along the flights.

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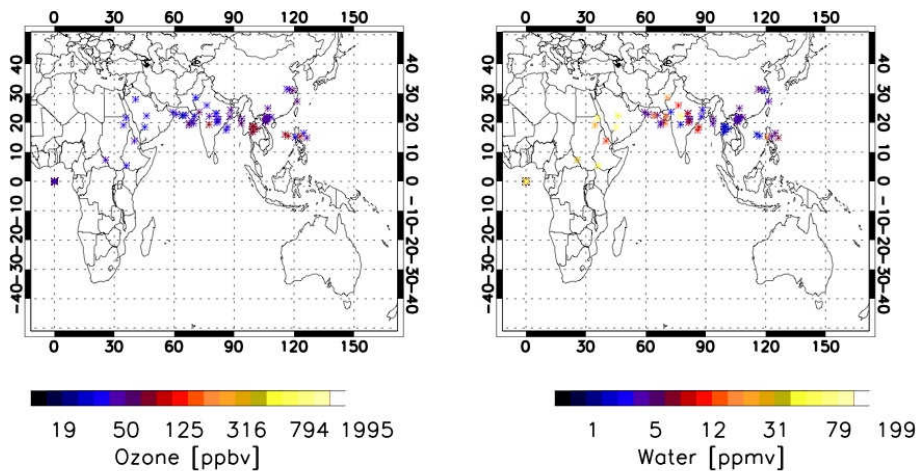


Fig. 7. Origins of M55 flight track back trajectories from below 800 hPa, colour coded in terms of ozone in ppbv (left panel) and water vapour in ppmv (right panel) observed along the flights.

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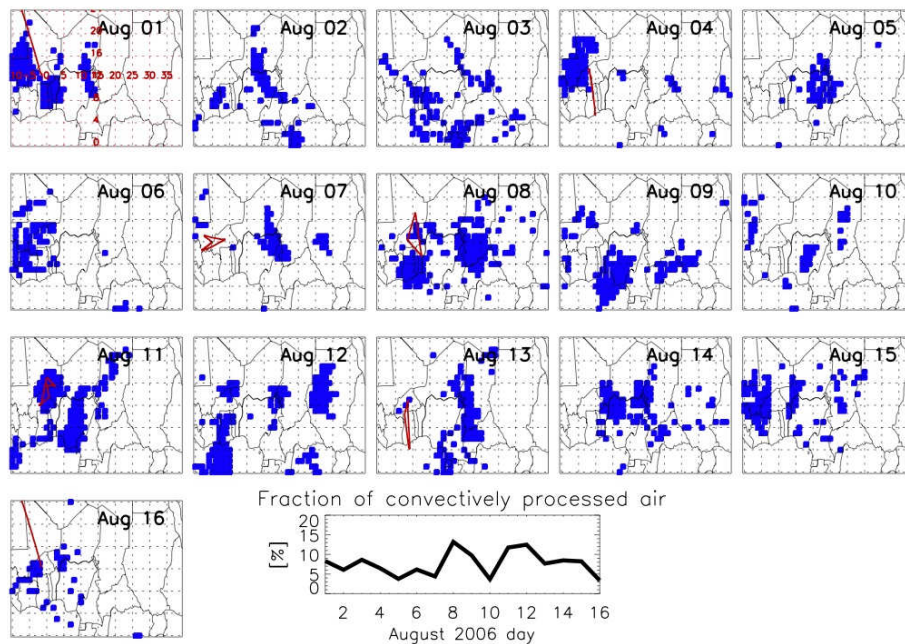


Fig. 8. Daily maps of convective impact at 200 hPa level derived from the percentage of trajectories passing over convective systems diagnosed from MSG cloud tops (blue points, see text for details). M55 flight paths are displayed for the days flights took place. Also shown (inset, bottom right) is an estimate of the fraction of convectively processed air for the West African region. Only grid points where at least 5% of trajectories have been “processed” by convection were used in the calculation of the average.

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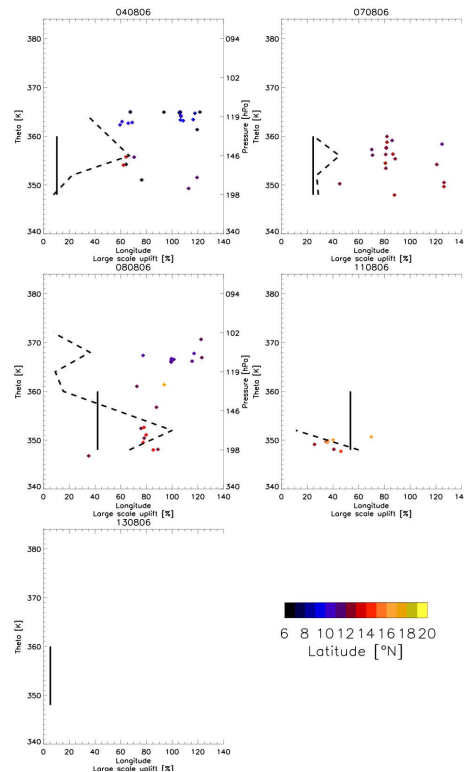


Fig. 9. Fraction of air masses influenced by large-scale convective uplift (%) calculated using 10 day back trajectories arriving along M55 flights in 4 K bins between 340 K and 375 K over West Africa for each local M55 flight (dashed lines). Also shown are average longitudes where air masses arriving along local flights were uplifted irreversibly above 800 hPa coloured by latitude where air masses were sampled by the M55. The average fraction of back trajectories from flight tracks between 350–360 K coinciding with MSG cloud tops (%) are shown as a vertical bar as an indicator of local convective impact (see text for details).

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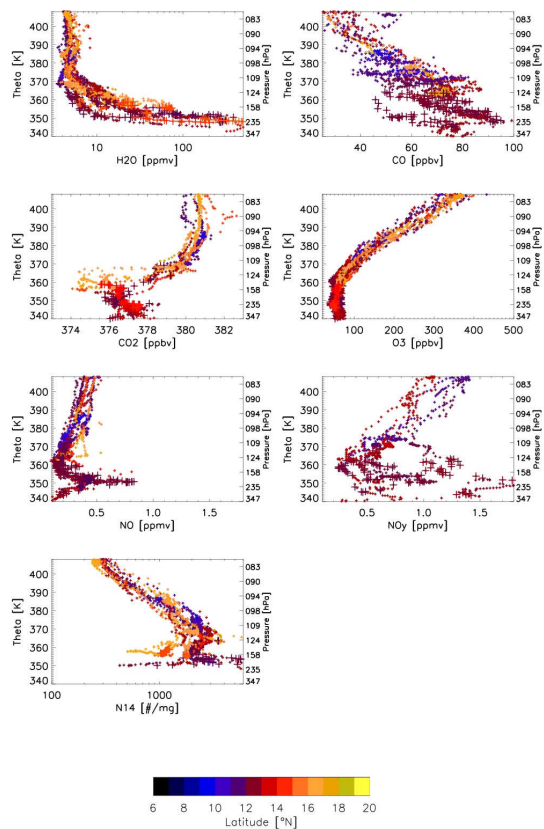


Fig. 10a. Vertical profiles of trace gas concentrations (volume mixing ratio in ppbv or ppmv as appropriate) and aerosol mass mixing ratios (particles per mg air), N_{14} (14 nm to $<1 \mu\text{m}$), for “convective” M55 flights on 7, 8 and 11 August 2006. Profiles are plotted as a function of theta (K) and coloured by latitude where the measurements were made. Crosses indicate points where back trajectories show uplift from below 800 hPa during 10 days before the flight.

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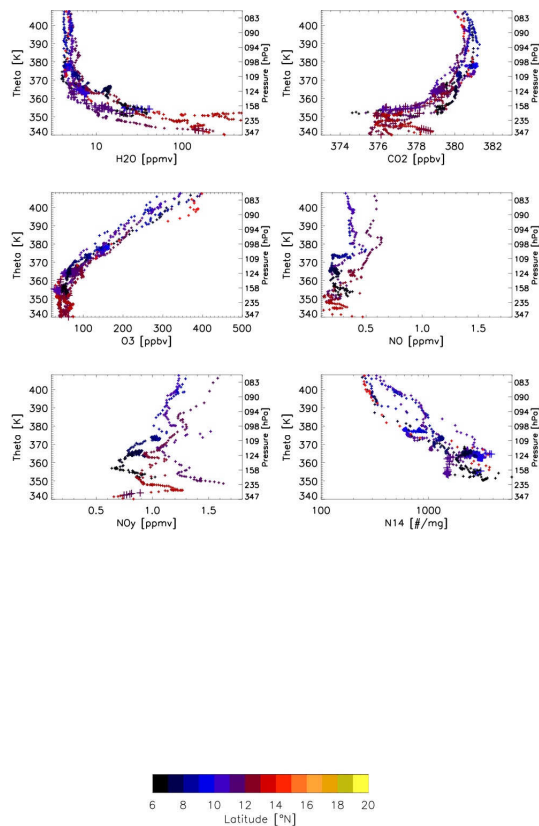


Fig. 10b. Vertical profiles of trace gas concentrations (volume mixing ratio in ppbv or ppmv as appropriate) and aerosol mass mixing ratio (particles per mg air), N_{14} (14 nm to $<1\ \mu\text{m}$), for “non-convective” M55 flights on 4 and 13 August 2006. Profiles are plotted as a function of theta (K) and coloured by latitude where the measurements were made. Crosses indicate points where back trajectories show uplift from below 800 hPa during 10 days before the flight.

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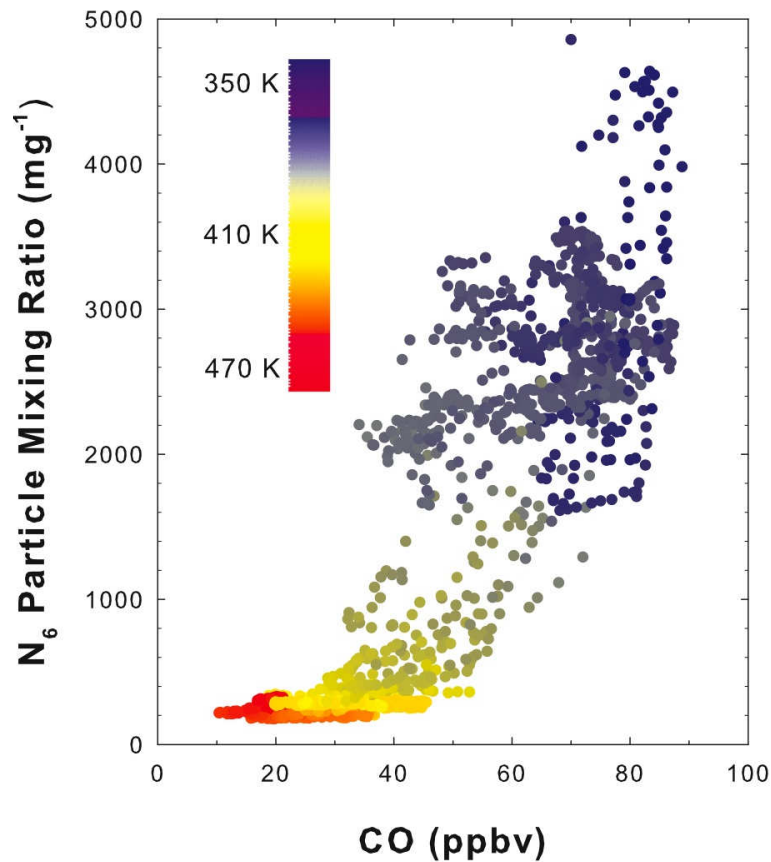


Fig. 11. Scatter plot of CO (ppbv) versus aerosol mass mixing ratio (greater than 6 nm, less than 1 μm) (particles per milligram of air) for all the local M55 flights over West Africa coloured as a function of potential temperature (K). Aerosol nucleation events have been removed (see text for details).

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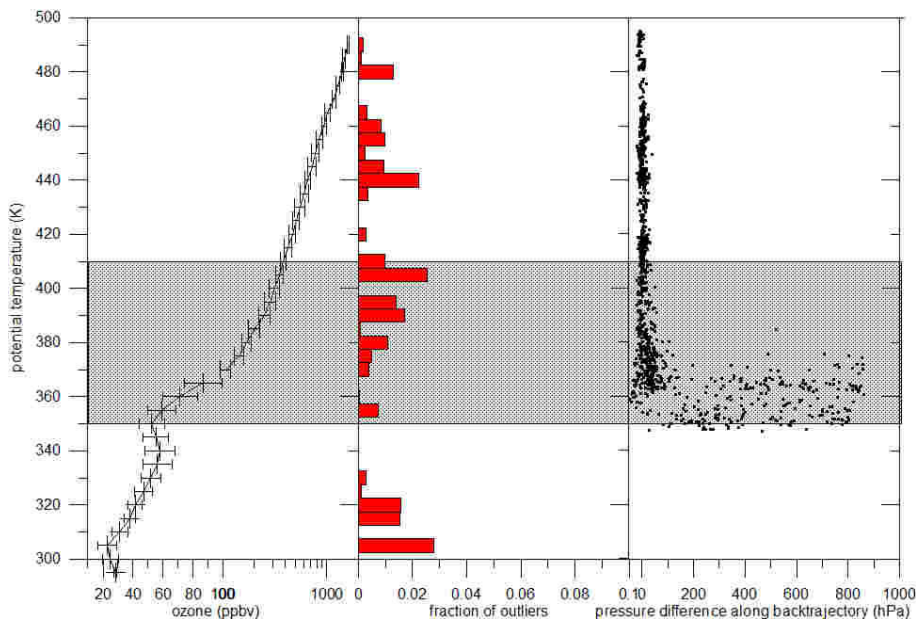


Fig. 12. Leftmost profile, average ozone concentrations (ppbv) using M55 data collected south of 17.5° N versus potential temperature (K), bars denote one standard deviation; note the linear (<100 ppbv)/log(>100 ppbv) horizontal axis for ozone. Middle bar chart, fraction of ozone data points lying outside three standard deviations from the mean, versus potential temperature; Rightmost panel, pressure difference (hPa) along 10-day back trajectories arriving along M55 flights (south of 17.5° N) versus potential temperature (K). Grey shading denotes the TTL region.

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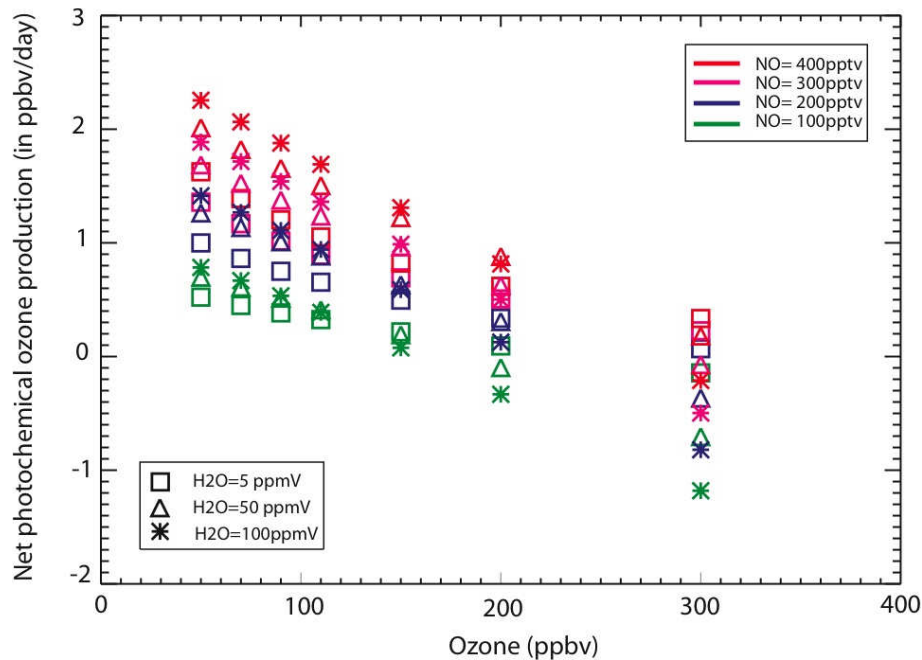


Fig. 13. Net photochemical O₃ production rates (ppbv per day) versus initial ozone concentrations (ppbv) as a function of initial H₂O concentrations (ppmv, symbols – see inset) and NO concentrations (pptv, colours as indicated). Results from photochemical model simulations run for 4 days at 150 hPa and initialised with concentrations taken from aircraft observations in the upper troposphere. Net photochemical O₃ production rates are averages over the last 24 h. See text for further details.